Interactive comment on “Decoupling the effects of clear atmosphere and clouds to simplify calculations of the broadband solar irradiance at ground level” by A. Oumbe et al.

A. Oumbe et al.
lucien.wald@mines-paristech.fr

Received and published: 23 June 2014

ANSWERS TO REVIEWER 2

Comment 1. The authors do a nice job of ensuring realistic combinations of model parameters. Nevertheless I wonder if 20 cases of $P_c$ are enough, considering that each $P_c$ is comprised of 7 parameters. Especially since statistics over these 20 cases are the key measure of whether the separability assumption is satisfied (the 95th percentile will separate a single case from all of the rest). If computational burden is an issue, perhaps fewer steps could have been used in some of the other parameters. Nevertheless the quality of the results, such as shown by their consistency between solar
zenith angles, demonstrates that the set of cases studied is sufficient to support the conclusions made.

We fully agree. Actually, we made the computations twice independently at two different institutes by different persons, plus another one with the STREAMER model. Because they are randomly selected, the 20 cases of Pc were different in each study. Conclusions were similar, which is in agreement with the last sentence of the comment by the reviewer. Practically, it was difficult to concatenate all computations to get 60 random selected cases of Pc. We have opted to present the results of one of these three computation series. In addition, numerous additional runs were made to better understand the results. For example, new computations were made to better answer comments.

Comment 2. P2011, L26: A fixed value is used for the cloud droplet effective radius \(r_{\text{eff}}\) for each of water and ice cloud. What is the sensitivity to variations in \(r_{\text{eff}}\)? For instance, the near-infrared cloud reflectance would be expected to change with \(r_{\text{eff}}\), and Nakajima and King (1990, Journal of the Atmospheric Sciences, 47, 1878-1893) studied the retrieval of \(r_{\text{eff}}\) over a range of at least 2 to 32 micrometres.

We have used the values of 10 \(\mu\)m for water cloud, and 20 \(\mu\)m for ice cloud for effective radius of droplets. In a preliminary study (PhD of A. Oumbe 2009), the influence of the changes in effective radius, from 3 to 50 \(\mu\)m was found negligible for ice clouds. For water clouds, the smaller the radius, the greater the influence, though this influence is still negligible with respect to other variables. We acknowledge that the study could be completed by taking into account changes in effective radius.


Text has been revised accordingly: "Default values in libRadtran for cloud liquid content
and droplet effective radius are used: 1.0 g m\(^{-3}\) and 10 \(\mu\)m for water cloud, and 0.005 g m\(^{-3}\) and 20 \(\mu\)m for ice cloud. In a preliminary study (Oumbe, 2009, Fig. 4.6, p. 53), the influence of the changes in effective radius, from 3 to 50 \(\mu\)m was found negligible for ice clouds. For water clouds, the smaller the radius, the greater the influence, though this influence is still negligible with respect to other variables."

Comment 3. P2019, L2: The authors might consider noting the maximum albedo expected for desert regions (no more than around 0.5 I think), since many users of downstream products will be interested in the performance in snow-free arid regions.

We fully agree. Northern Africa and Arabia are desert areas and exhibit large ground albedo up to approximately 0.5 (Tsvetsinskaya et al., 2002; Wendler and Eaton, 1983). Text has been revised accordingly, page 2015, lines 16-17. The text is now: "The diffuse irradiance \(D\) and therefore \(G\) are strongly influenced by rho\(_g\). The influence of changes in \(P_c\) on \(K_c\) increases with rho\(_g\). Deserts such as Northern Africa and Arabia exhibit large ground albedo up to approximately 0.5 (Tsvetsinskaya et al., 2002; Wendler and Eaton, 1983); the error (P95) on \(G\) is of order of 10 W m\(^{-2}\). Fresh snow-covered or ice-covered areas may exhibit very large rho\(_g\). For rho\(_g\)=0.9, the error on \(G\) can be large for small \(\tau_S\), i.e. 30 W m\(^{-2}\). One has to be cautious in using Eq. 3 in such extreme cases."


Comment 4. Figures 3 and 4 plot results in relative terms. Consider plotting the results in absolute terms as well, since as you point out the absolute errors are important to consider in practice. Perhaps a second panel could be added to Figure 4 showing the
same results as the current Figure 4 plot but in absolute terms.
Done for Figs 3 and 4.

Comment 5. P2016, L11-12: It is true that the uncertainty contributed by the assumption that changes in Pc can be neglected is within the WMO criterion of high quality measurements. However, in application there will be other contributions such as uncertainties in cloud property retrievals, aerosol amount and type, satellite calibration, radiative transfer model approximations, etc. The total uncertainty of output from any system using this assumption will be higher and probably outside the WMO high quality threshold.

Text has been changed and is now: These results match the WMO requirements for high quality measurements. However, in applications as discussed in the following section, there will be other sources of uncertainties, and the total uncertainty of any model using Eq. 3 will be greater and probably exceeding these WMO requirements.

Technical correction 6. P2019, L4: Change “pyranometer” to “pyrheliometer”. Done

Technical correction 7. P2019, L28: The authors intend to mean a reduction for each of points (i), (ii) and (iii). Therefore, “reducing” should be moved before “(i)” to say this. Done

Interactive comment on Geosci. Model Dev. Discuss., 7, 2007, 2014.