Abstract. The ability of the chemical transport model (CTM) PMCAMx to reproduce aerosol optical depth (AOD) measurements by the Aerosol Robotic Network (AERONET) and the Moderate Resolution Imaging Spectroradiometer (MODIS) over Europe during a photochemically active period is evaluated. Periods with high dust levels are excluded so the analysis focuses on the ability of the model to simulate the mostly secondary aerosol and its interactions with water. PMCAMx reproduces the monthly mean MODIS and AERONET AOD values over the Iberian Peninsula, the British Isles, central Europe, and Russia with fractional bias less than 15% and fractional error less than 30%. However, the model overestimates the AOD over northern Europe probably due to an overestimation of organic aerosol and sulfates. On the other end, PMCAMx underestimates the monthly mean MODIS AOD over the Balkans, the Mediterranean, and the South Atlantic. These errors are probably due to an underestimation of sulfates. Sensitivity tests indicate that the evaluation results of the monthly mean AODs are quite sensitive to the relative humidity (RH) fields used by PMCAMx, but are not sensitive to the simulated size distribution and the black carbon mixing state.
1 Introduction

Atmospheric aerosols are suspensions of solid and/or liquid particles in air that scatter and absorb light. The aerosol optical depth (AOD) is defined as the integrated extinction coefficient over the entire atmospheric column and is a measure of the total aerosol loading (King et al., 1999; Kokhanovsky, 2008; Vijayarachavan et al., 2008; Hidy et al., 2009). Calculations of AOD require knowledge of the aerosol vertical profile, including the particulate matter size distribution, chemical composition, and microphysical state (Seinfeld and Pandis, 2006).

Aerosol properties can be retrieved from ground-based measurements as well as from satellite earth observations (Holben et al., 1998; Levy et al., 2007a, b; Kokhanovsky, 2008; Levy et al., 2010; Duncan et al., 2014; Hu et al., 2014). Global observations of high spatial coverage are provided by satellites (King et al., 1999; Vijayarachavan et al., 2008; Hidy et al., 2009) and more limited spatial coverage by ground-based stations. Regarding temporal coverage, satellite observations are sparse when compared against ground measurements. Ground-based measurements of AOD are direct measurements while satellite AOD measurements are indirect, resulting from inversion procedures and exhibiting larger uncertainties. The magnitude of the satellite AOD uncertainties is higher over land where the surface reflectance cannot be neglected and it must be retrieved simultaneously with the aerosol properties (Levy et al., 2007a, 2007b, 2010). The satellite inversion procedure is simpler over water since the surface contribution is small and the detected signal is mostly due to aerosol reflectance (Shi et al., 2011; Anderson et al., 2013; Schutgens et al., 2013).

Chemical transport models (CTMs) are valuable tools for the study of the impact of pollutant emissions, the development of air quality improvement strategies, studies of aerosol radiative forcing, visibility, and global climate change. Uncertainties of the CTM’s input data, including meteorological fields, emission inventories, and boundary conditions as well as weaknesses in representation of atmospheric processes may lead to weak model performance (Kinne et al., 2003, 2006). CTMs have
been used in the past to provide AOD predictions either globally (Chin et al., 2002, 2004; Lee et al., 2010; Johnson et al., 2012; Meij et al., 2012; Pozzer et al., 2012; Yu et al., 2012) or over specific regions like Asia (Han et al., 2010; Park et al., 2011), United States (Roy et al., 2007), and Europe (Jeuken et al., 2001; Hodzic et al., 2006; Meij et al., 2007; Tombette et al., 2008; Myhre et al., 2009; Carnevale et al., 2011; Im et al., 2014). Model evaluation often relies on in-situ ground measurements but also measurements from airborne platforms. These in-situ measurements cover by necessity a limited part of the modeling domain. Comparisons against remote sensing data have been used to close that gap.

Jeuken et al. (2001) compared the TM3 CTM AOD predictions with the ATSR-2 radiometer AOD retrievals during a 1997 summer episode over Europe. Model errors (neglecting organics and mineral aerosol) in the vertical distributions of sulfate, ammonium, and nitrate, in the hygroscopic growth, and in the optical parameters led to an average AOD (at 550 nm) underestimation by 0.17-0.19. Hodzic et al. (2006) used the CHIMERE model to simulate AOD at 865 nm over Europe during August 2003. The model generally reproduced AOD within a factor of 2 and with correlation coefficients ranging from 0.4 to 0.6 in comparison with POLDER and AERONET. Sporadic aerosol emissions due to forest fires or dust events led to regional AOD underestimations. Meij et al. (2007) used the mesoscale TAPOM model to investigate AOD over Milan, Italy during June 2001. Simulated and observed AODs by AERONET, MODIS and MISR (Multi-angle Imaging Spectroradiometer) differed by a factor of 2 or 3 in days with cirrus clouds and Saharan dust but showed good agreement in clear sky days. A finer model resolution gave a more detailed AOD distribution pattern and improved by 15% the agreement with the AOD observations. Tombette et al. (2008) compared the Polair3D estimated AOD against AERONET measurements over Europe for 2001. The black carbon (BC) mixing state had almost no effect on the estimated single scattering albedo (SSA) but the aerosol water content influenced significantly both the SSA and the AOD. Myhre et al. (2009) used the global Oslo CTM2 to
predict AODs at 550 nm focusing on specific European regions (Adriatic Sea, Black Sea, and Po Valley). Comparisons against AOD measurements from AERONET, MODIS, and MISR were presented for a short period during late summer-early autumn of 2004. The model underestimated AOD around Venice against AERONET because of organic carbon underestimation. Carnevale et al. (2011) implemented the TCAM CTM to simulate AODs during 2004 over Italy. In general, TCAM was found to underestimate MODIS AODs. Analysis of the extinction coefficient showed that the submicron inorganic aerosol played a key role. Im et al. (2014) simulated air pollution over Europe using the WRF-CMAQ modeling system for 2008. The model underestimated AERONET AOD measurements by 3-22% on average. AOD underestimations were attributed to underestimation of either the anthropogenic emissions or the natural and re-suspended dust emissions.

The PM$_1$ composition predictions of PMCAMx have been evaluated over Europe in May 2008 (Fountoukis et al., 2011). PMCAMx performance against airborne measurements was as good as its performance against the hourly ground measurements. More than 94% of the organic aerosol (OA) hourly values and more than 82% of the sulfate ones were reproduced within a factor of 2. However, the evaluation was limited in space. The model performance was only evaluated against ground measurements which were taken at 4 stations located in the Netherlands, Greece, Ireland, and Germany as well as against airborne measurements from 15 flights in North-Western Europe.

One of the limitations of the previous evaluation exercises is that errors in dust emissions, transport, and removal often dominate the overall results. In the present work MODIS and AERONET AODs are filtered to exclude periods with high dust levels and to focus on the rest of the anthropogenic and biogenic aerosol components. A period with high photochemical activity is selected so that the emphasis is on secondary aerosol components.

In the present study we provide a first time evaluation of the ability of PMCAMx (Murphy and Pandis, 2009; Fountoukis et al., 2011) to reproduce AOD observations over Europe. The objective of
this work is to identify weaknesses and strengths of PMCAMx and its inputs, by taking advantage of the wide spatial coverage of MODIS and the temporal coverage of AERONET.

2 PMCAMx description

PMCAMx-2015 is a three-dimensional CTM that employs the framework of CAMx (Environ, 2003) simulating the processes of horizontal and vertical advection, horizontal and vertical dispersion, wet and dry deposition as well as gas, aqueous, and aerosol-phase chemistry. Three detailed aerosol models are employed: inorganic aerosol growth (Gaydos et al., 2003; Koo et al., 2003), aqueous-phase chemistry (Fahey and Pandis, 2001) as well as OA formation and chemical aging (Murphy and Pandis, 2009). The specific modules utilize a sectional approach that dynamically models the evolution of the aerosol size distribution. Ten size sections covering particle diameters from 40 nm to 40 μm are used. PMCAMx treats both primary and secondary organic aerosol as semivolatile and photochemically reactive employing the volatility basis set (Murphy and Pandis, 2009).

The PMCAMx European modeling domain in this application is a region of 5,400x5,832 km² with 36x36 km² grid resolution and 14 vertical layers extending up to approximately 6 km. Simulations were performed on a polar stereographic map projection. Horizontal wind components, vertical diffusivity, temperature, pressure, water vapor, clouds, and rainfall were provided by the Weather Research and Forecasting meteorological model (WRF) (Skamarock et al., 2008). Anthropogenic gas emissions are from the European emissions database GEMS while elemental carbon and organic carbon emissions are from the Pan European Carbonaceous Aerosol Inventory (Kulmala et al., 2011). Biogenic emissions were calculated by the MEGAN v2.04 model (Guenther et al., 2006). The marine aerosol emission model developed by O' Dowd et al. (2008) was employed for the estimation of mass fluxes for both accumulation and coarse mode, including the organic aerosol fraction. Emissions from wildfires are taken from IS4FIRES (Sofiev et al., 2009).
To limit the effect of the initial conditions on the results, the first six days of each simulation were excluded from the analysis. Concentrations of the major PM$_{2.5}$ aerosol components at the boundaries of the domain (Table S1 in Supplementary Information) are based on measurements of typical background concentrations in sites close to the domain boundaries (Zhang et al., 2007; Seinfeld and Pandis, 2006). All concentrations given here are under ambient temperature and pressure conditions.

2.1 AOD prediction by PMCAMx

The size and chemically resolved concentrations of aerosol particles are simulated by PMCAMx for every computational cell. Inorganic aerosol water concentration is calculated online by the thermodynamic equilibrium model ISORROPIA (Nenes et al., 1998). Taking into account all the vertical layers, we calculate the PMCAMx AOD at 550 nm as the sum of the extinction coefficients at each layer:

$$\text{AOD} = \sum_{i=1}^{14} b_{\text{ext},i} \Delta z_i$$

(1)

where $b_{\text{ext},i}$ is the extinction coefficient of layer $i$ and $\Delta z_i$ is the corresponding layer thickness. Assuming that the particles are homogeneous spheres and that all particles in each size bin have the same composition (internal mixture), the aerosol extinction coefficient ($b_{\text{ext},i}$) for layer $i$ is:

$$b_{\text{ext},i} = \sum_{j=1}^{10} \frac{\pi D_j^2}{4} N_j Q_{\text{ext},j} (m_j, D_j)$$

(2)

where $D_j$ is the mean diameter of size bin $j$, $N_j$ is the aerosol number concentration for bin $j$ and $Q_{\text{ext},j}$ is the extinction efficiency of a single particle having a complex refractive index $m_j$. The extinction efficiency for bin $j$ is estimated as the sum of the scattering, $Q_{\text{scat},j}$, and absorption, $Q_{\text{abs},j}$, efficiencies:
\[ Q_{\text{ext},j} = Q_{\text{scat},j} + Q_{\text{abs},j} \tag{3} \]

Aerosol scattering and absorption efficiencies \((Q_{\text{scat},j}, Q_{\text{abs},j})\) are calculated using Mie theory (Seinfeld and Pandis, 2006) and mass concentrations provided for each size bin by PMCAMx, including the concentrations of particulate water. The complex refractive, \(m_j\), index of a homogeneous sphere is estimated using the volume weighted average of the individual refractive indices (Pilinis and Pandis, 1995). Sulfate and ammonium are assumed to have a real refractive index of 1.53, which is the value of ammonium sulfate (GEISA, 2011; NASA, 2006). Nitrate is assumed to have a real refractive index of 1.56, similar to the value of ammonium nitrate (NASA, 2006). Sodium and chloride have a real refractive index of 1.5 (GEISA, 2011; NASA, 2006). Dust is assumed to have a complex refractive index of 1.53-0.0055i (GEISA, 2011). OA is assumed to be non-absorbing with a refractive index of 1.5 (Nessler et al., 2005; Fierz Schmidhauser et al., 2010). Biomass burning was minimal during the period of interest (Crippa et al., 2014), so this simplifying assumption regarding the OA absorptivity has little effect on the predicted AOD. The black carbon refractive index has the largest uncertainties (Bond and Bergstrom, 2005) and we use a value of 1.75-0.44i (GEISA, 2011). In the base case BC is assumed to be internally mixed with the other components in each size range. The sensitivity of the model predictions to this assumption is discussed in a subsequent section.

3 MODIS and AERONET data

The cloud screened and quality assured Level 2 AERONET direct AOD measurements are used for the PMCAMx evaluation. AERONET applies the Beer Lambert Bouguer law to measure AOD from direct sun observations (Holben et al., 1998) therefore it is considered to be the ground truth with AOD uncertainties of 0.01 – 0.02 (Eck et al., 1999). In this work only the AOD values corresponding to Angstrom Exponent values greater than 0.9 are employed in an effort to exclude periods with high dust levels (Schuster et al., 2006). This filter rejected 29% of AODs over land and 28% over water. The geographical distribution of the corresponding AERONET stations is depicted in Fig. 1 and the number...
of stations in each region is shown in Table 1. Some AERONET stations in the domain of interest did not have available Level 2 AOD data for the period of interest while all data from three stations (OHP_OBSERVATOIRE in South France, FORTH_CRETE in Crete, Greece, and ATHENS_NOA in Athens, Greece) have been excluded after the dust coarse particle rejection filtering.

The polar-orbiting MODIS monitors global aerosol properties from two satellites: Terra and Aqua (Salomonson et al., 1989). MODIS employs 36 channels from 0.412 to 14.2 μm, has a wide swath of 2,330 km, and observes every part of the globe at least once daily. The default resolution for aerosol retrieval is 10x10 km² (Levy et al., 2009). Each data set retrieved by MODIS is associated with a Quality Assurance Confidence (QAC) flag which ranges from 0 (no confidence) to 3 (highest confidence). For increased spatial coverage we use the union of Terra and Aqua MODIS AOD retrievals with QAC ≥ 1. We employ the MODIS Level 2 Collection 5.1 aerosol datasets. AOD retrievals are provided at seven wavelengths (470, 550, 660, 870, 1,200, 1,600, 2,100 nm) over water surface and four wavelengths (470, 550, 660, 2,100 nm) over land. In this study we focus on the 550 nm values. Figure 2 presents the geographical distribution of the available MODIS AOD measurements during the period of interest (1-29 May 2008) over Europe. The average number of retrievals is 12±9. The maximum number of retrievals is 65 in areas in the North Atlantic.

Dust emissions from the Sahara are not included in the PMCAMx emissions used here and the focus of this study is on periods and regions in which Saharan dust does not contribute significantly to the AOD. To exclude periods with high dust levels and to focus on the rest of the anthropogenic and biogenic aerosol components, MODIS AODs are filtered. Over water we employ the dust coarse particle rejection filter of Barnaba and Gobbi (2004). According to this filter, AOD values greater than 0.3 also corresponding to coarse mode fraction higher than 0.3 are assumed to be dust-influenced periods. Over land we only use the AOD values which correspond to Angstrom Exponent values...
exceeding 0.9 (Schuster et al., 2006). The above filters discard 16% of MODIS AOD values over land and 0.4% over water.

The evaluation of the MODIS AODs at 550 nm for the land algorithm was performed following the approach of Remer et al. (2005) and Levy et al. (2007b). The collocated data were sorted according to the AERONET measurements. The resulting data were partitioned into bins of 100 points and then averaged. At higher optical depths since the data became sparser we employed 25 points for each bin. The regression results of the collocated AODs, prior to binning had a slope of 1.05. 73% of the 8,331 collocated points fall within the expected error envelope denoted by the dashed lines in Fig. S1. These results indicate that the mean MODIS AOD over land in the region and period of interest was retrieved with the expected accuracy. The highest quality flag QAC = 3 provides the closest match but including the QAC = 2 and 1 retrievals results in only a minor reduction of accuracy while increasing significantly the size of the dataset (Table S2).

Previous studies have shown that MODIS AOD retrievals have an expected error of \( \pm (0.05 + 0.15\text{AOD}_{\text{AERONET}}) \) over land and \( \pm (0.03 + 0.05\text{AOD}_{\text{AERONET}}) \) over water (Chu et al., 2002; Remer et al., 2005; Levy et al., 2007a,b, 2010; Anderson et al., 2013). Table S3 summarizes the values of the expected MODIS AOD uncertainties for the various regions in our modeling domain during May 2008, based on the monthly mean values of AERONET AOD. The MODIS-AERONET AOD differences for this period are consistent with the expected uncertainty of the MODIS retrievals (Fig. S1).

4 Evaluation of PMCAMx AOD predictions

The dust-screened monthly mean AODs for Europe during May 2008 retrieved by MODIS and predicted by PMCAMx are shown in Fig. 3. The PMCAMx AODs have been calculated for exactly the same periods as the MODIS retrievals to allow the direct comparison of the two.

The MODIS retrievals show high AOD values (> 0.25) over England, South Ireland, North Italy, South Poland, East Romania, Greece, and North Atlantic. Low AOD values (< 0.1) were retrieved over
East France, Belgium, Sweden, and North Russia. PMCAMx predicts high AODs over England, South Ireland, North Italy, and central Atlantic and low AODs over North Sweden, East Russia, North and South Atlantic.

### 4.1 Overall evaluation

The difference between PMCAMx and MODIS monthly mean AODs is depicted in Fig. 4. PMCAMx AODs are higher than those of MODIS over England, Ireland, France, Germany, central and South Italy, North and East Europe, central, North and West Russia, West Balkans, and central Atlantic. On the other hand PMCAMx predicts lower AODs than MODIS over parts of Russia, North Italy, central and South Balkans, South Poland, North and South Atlantic, and the African coast of the Mediterranean. On a domain average basis PMCAMx predicts an AOD equal to 0.14 while MODIS retrieved 0.16. Detailed comparisons for each region can be found in Table 2. 94% of the monthly mean AOD values fall inside the expected MODIS error envelope over land (Fig. 5a). Over the whole domain the PMCAMx monthly mean AODs have a mean error of 0.05 and a fractional bias of -16% compared to the MODIS monthly mean AODs (Table 2).

PMCAMx AODs were also compared against the AERONET values for the simulation period. Once more the comparisons were done for the grid cells of the AERONET stations and corresponding measurement periods. The PMCAMx monthly mean AODs had a mean error of 0.03 and a fractional bias of 4% compared to the AERONET monthly mean AODs (Table 1). The comparison of the PMCAMx with AERONET monthly mean AODs is summarized in Fig. 5b for the 50 AERONET stations which are employed in the present study.
4.2 Regional evaluation

Spain and Portugal: The relatively low AOD levels (0.11 for the 8 AERONET stations and 0.14 for MODIS) are reproduced well by PMCAMx (0.12 for the AERONET sites and 0.12 for the periods of the MODIS retrievals). The monthly mean PMCAMx AOD predictions have a mean error of 0.02 (AERONET) and 0.04 (MODIS) (Tables 1 and 2). The model shows little bias (5%) compared to the AERONET stations and a small tendency towards underprediction (-15%) compared to MODIS. 83% of the monthly mean PMCAMx AODs are within the expected MODIS error envelope. Sulfate and organic aerosol are the major components of dry fine PM in Spain and Portugal (Table 3).

Russia, Belarus, and Ukraine: PMCAMx reproduces well (0.14 predicted versus 0.15 measured) the average AOD observations at the 5 AERONET stations in this region (2 in West Russia, 1 in Belarus, 1 in Ukraine, and 1 in Crimea) (Table 1). The model has a similar good performance against the MODIS retrievals (0.12 predicted versus 0.13 retrieved) (Table 2). As a result, the monthly mean PMCAMx AOD predictions have a low mean error of 0.02 (AERONET) and 0.04 (MODIS). PMCAMx shows a slight tendency towards underprediction (-11%) compared to AERONET and no bias (<1%) compared to MODIS. 92% of the monthly mean PMCAMx AODs are within the expected MODIS error envelope. Sulfates and organic aerosol predominate in this region and it appears that PMCAMx performs reasonably well in this ground-level measurement poor region.

United Kingdom (UK) and Ireland: This area was relatively polluted during the simulation period with high levels of nitrates, sulfates, and organic aerosol (Table 3). PMCAMx reproduces the relatively high average MODIS (0.23 predicted versus 0.21 retrieved) and AERONET (0.24 predicted versus 0.25 measured in the station of Chibolton). The monthly mean PMCAMx AODs have a mean error of 0.04 compared to MODIS with a small tendency towards overprediction (14%). 90% of the monthly mean PMCAMx AODs fall within the expected MODIS error envelope.
**Balkans:** The Balkans according to PMCAMx had some of the highest sulfate levels in the domain during the simulation period (Table 3). The model underpredicts the AOD both against MODIS (0.14 predicted versus 0.19 retrieved) and the two AERONET stations (0.15 predicted versus 0.21 measured). The corresponding fractional biases are -24% against MODIS and -33% against AERONET. However, 80% of the monthly mean PMCAMx AODs fall within the expected MODIS error envelope. Given that most of the predicted AOD is due to the sulfate these results suggest that the PMCAMx underprediction is probably due to their underestimation.

**Central Europe:** PMCAMx showed a small tendency towards overprediction of the moderate AODs in this region compared to both AERONET (12%) and MODIS (13%). For example, overpredictions were evident over France and Germany (Fig. 4). The corresponding fractional errors on a monthly average basis were 22% against AERONET and 30% against MODIS. Sulfate and organic aerosol are the major fine PM components in central Europe and they are probably slightly overestimated by PMCAMx in this region during the simulation period. Errors in the relatively humidity fields could also explain parts of these discrepancies.

**East Europe:** PMCAMx slightly overpredicted the AODs in this region compared to both AERONET (24%) and MODIS (25%). 82% of the monthly mean PMCAMx AODs fell within the expected MODIS error envelope. PMCAMx predicts more frequently AODs > 0.1 than measured by AERONET during the corresponding period of measurements probably because of an overestimation of sulfates and organic aerosol.

**North Europe:** PMCAMx reproduces the low pollution levels in this area with a mean AOD of 0.12 compared to 0.08 by the 4 AERONET stations. The absolute monthly mean errors are low: 0.04 (AERONET) and 0.06 (MODIS). However, there is significant fractional positive bias compared to
both AERONET (36%) and MODIS (47%). 54% of the monthly mean PMCAMx AODs fall outside the expected MODIS error envelope. 53% of the hourly PMCAMx AODs are greater than 0.1 while only 18% of the AERONET values are greater than 0.1. Sulfates and organic aerosol are the dominant fine PM components in this region and at least one of them is overestimated by the model.

Turkey and Northern Africa: There are only two AERONET stations in this area and PMCAMx underpredicts by 30% the corresponding moderate AOD measurements. However, the model performance appears to be much better against the MODIS retrievals covering a much bigger area and the underprediction drops to 13%. 79% of the PMCAMx AODs fall inside the expected MODIS error envelope. Sulfates and organic aerosol are the major fine PM components in these regions and they are probably slightly underestimated by the model.

Mediterranean Sea: PMCAMx exhibits a tendency towards underprediction (-24%) against MODIS and 58% of the monthly mean PMCAMx AODs fall inside the expected MODIS error envelope. The major discrepancies are evident in the southern part of the Mediterranean especially close to the African coast. These suggest that dust may be partially responsible for the errors even after the filtering of the data. The model performance is better in the eastern Mediterranean (Fig. 4). Sulfates dominated the AOD in the Mediterranean during the simulation period according to PMCAMx.

South Atlantic: The PMCAMx AOD predictions are significantly lower (-45%) compared to the MODIS retrievals in this region with 74% of the monthly mean PMCAMx AODs falling outside the expected MODIS error envelope. Sulfate and sea-salt dominated the predicted AOD in this region in May 2008 and there is evidence that they may be underpredicted. However, errors in relative humidity or cloud contamination could be also responsible for these discrepancies (Anderson et al., 2013).
North Atlantic: The model performance is much better in the North than in the South Atlantic. The  
mean AOD error is 0.04 compared to MODIS with a tendency towards underprediction (-21%). 53% of 
the monthly mean PMCAMx AODs fall inside the expected MODIS error envelope. There is one 
AERONET station in this area (in Helgoland around 50 km from the coast of Germany) and PMCAMx 
predicts an average AOD equal to 0.16 compared to the 0.11 measured. Sulfates, organic aerosol, and 
sea-salt were the major fine PM components in North Atlantic during May 2008 (Table 3).  

Black Sea: PMCAMx exhibits a tendency towards underprediction (-18%) versus MODIS in this 
relatively polluted region. 66% of the PMCAMx AODs fall within the expected MODIS error envelope. 
Sulfates were the major fine PM components in the Black Sea during the simulation period.  
The results of the PMCAMx-MODIS comparison for the various regions are summarized in Fig. 6. 
These results suggest that the variability of the MODIS retrievals exceeds that of the PMCAMx 
predictions for almost all areas. This could be due to the spatial resolution of the model inputs including 
the emissions, but also to other reasons including missing short-term air pollution sources in the 
inventory, potential cloud contamination of the retrievals, etc.  

5 Sensitivity analysis of the predicted AODs  
There are various possible sources of bias in the PMCAMx predictions of AOD other than the 
concentration and composition of aerosol. We explore here the role of the relative humidity calculated 
by the WRF model, the role of the mixing state of BC, as well as that of the predicted aerosol size 
distribution.  

In the first test the absolute humidity was increased uniformly by 5%, while maintaining the 
maximum relative humidity in cloud-free regions at 99%. The PMCAMx monthly mean AOD increased 
on average by 13% (Fig. S2). The increases ranged from 7% in Turkey and Northern Africa to 31% in
the North Atlantic. This AOD change can explain a significant part of the base case discrepancies which cause a fractional error of PMCAMx 22% versus AERONET and 33% versus MODIS.

In another test the diameter of all particles was increased by 20%. 72% of the PMCAMx monthly mean AOD values changed by less than 0.01. The average increase of the monthly mean AOD was 1% (ranging from 0.3% in the Black Sea to 4% in the UK and Ireland).

In a third sensitivity test we assumed that BC was always externally mixed with the other components in each size range, forming pure BC spheres. 73% of the PMCAMx monthly mean AOD values changed by less than 0.01 in this test. The average change of the monthly mean AOD was negligible (< 0.5%).

6 Conclusions

Previous evaluations of the ability of the 3-D CTM PMCAMx to reproduce the aerosol levels in Europe, the US, and Mexico City have been based on comprehensive chemical composition measurements at a few ground sites and limited data from a few flights. In this study we expand these efforts by using the MODIS retrievals of AOD over Europe during a photochemically active period (May 2008). Given the uncertainty of the MODIS AOD retrievals we compliment the evaluation using the corresponding AERONET observations in the same region which are more accurate but cover only specific areas. We exclude periods during which the different areas are strongly affected by dust (mainly from the Sahara) in an effort to focus on the other primary and secondary anthropogenic and biogenic aerosol components.

The details of the evaluation results differ, as expected, depending on the use of either the MODIS or the AERONET results. These differences are due to the different spatial coverage of the datasets, but also to the MODIS AOD retrieval uncertainties. The major conclusion is that PMCAMx can reproduce the observed AODs for this period with little bias (-16% for MODIS and +4% for
AERONET). The corresponding fractional errors are 33% against MODIS and 22% against AERONET. These results are consistent with those of Fountoukis et al. (2001) who compared the PMCAMx predictions for the same period against ground measurements of fine PM composition in four sites and airborne measurements from several flights over central and northern Europe.

The AOD performance of PMCAMx against the MODIS retrievals is excellent (absolute fractional bias less than 15% and fractional error less than 35%) in the Iberian Peninsula, UK/Ireland, central Europe, Russia-Belarus-Ukraine, Turkey-northern Africa. It is good (absolute fractional bias less than 25% and fractional error less than 35%) in East Europe, the Balkans, and over the Mediterranean, the North Atlantic, and the Black Sea. Finally, its performance is average (absolute fractional bias less than 50% and fractional error less than 55%) in the relatively clean area of North Europe and the South Atlantic. The performance is more or less similar against AERONET with the exception of a few areas with only one or two AERONET stations. The average performance against the AERONET measurements is considered using the above criteria excellent and against MODIS it is the borderline between good and excellent.

The above PMCAMx performance suggests that overall the model does a good job in reproducing the fine aerosol sulfate, organics, nitrate, and sea-salt levels over Europe during the evaluation period. Its major weaknesses appear to be potential overpredictions of sulfate and/or organics over North and East Europe, underprediction of sulfate over the Balkans, and underprediction of fine sodium chloride, sulfates, or organics in the southern Mediterranean and South Atlantic. However, these discrepancies are quite sensitive to the relative humidity fields predicted by WRF. In a sensitivity test the average predicted AOD increased by 13% (ranging from 7 to 31% depending on the area) for a uniform 5% change in RH. On the other hand, the details of the fine PM size distribution and the black carbon mixing state have a very small effect on the AOD predictions.
**Code availability**

PMCAMx is the research version of the publicly available CAMx (www.camx.org). The Fortran source code of CAMx (Version 6.20 was posted on March 23, 2015) and a User’s Guide both prepared by ENVIRON can be downloaded through the above website. The PMCAMx code is used as testbed for testing of different hypotheses, algorithms, etc. The version used in this paper (PMCAMx-2015) as well as the most current version can be obtained upon request by contacting Prof. S. Pandis (spyros@chemeng.upatras.gr).

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**References**


GEISA (Gestion et Etude des Informations Spectroscopiques Atmosphériques: Management and Study of Atmospheric Spectroscopic Information) 2011: http://www.pole-ether.fr/


Table 1. Error metrics for the evaluation of PMCAMx against AERONET monthly mean AODs.

<table>
<thead>
<tr>
<th>Region</th>
<th>Number of AERONET stations</th>
<th>Mean AERONET AOD</th>
<th>Mean PMCAMx AOD</th>
<th>Mean Error</th>
<th>Mean Bias</th>
<th>Fractional Error</th>
<th>Fractional Bias</th>
</tr>
</thead>
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<td>UK/Ireland</td>
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<td>0.25</td>
<td>0.24</td>
<td>0.01</td>
<td>-0.01</td>
<td>0.04</td>
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<td>Central Europe</td>
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<td>0.16</td>
<td>0.17</td>
<td>0.03</td>
<td>0.02</td>
<td>0.22</td>
<td>0.12</td>
</tr>
<tr>
<td>North Europe</td>
<td>4</td>
<td>0.08</td>
<td>0.12</td>
<td>0.04</td>
<td>0.04</td>
<td>0.36</td>
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<td>0.05</td>
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<tr>
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<td>0.24</td>
</tr>
<tr>
<td>Balkans</td>
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</tr>
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<td>0.16</td>
<td>-0.11</td>
</tr>
<tr>
<td>Turkey and Northern Africa</td>
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<td>0.17</td>
<td>0.12</td>
<td>0.05</td>
<td>-0.05</td>
<td>0.30</td>
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</tr>
<tr>
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<td>-</td>
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<td>-</td>
<td>-</td>
</tr>
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<td>0.05</td>
<td>0.05</td>
<td>0.37</td>
<td>0.37</td>
</tr>
<tr>
<td>South Atlantic Ocean</td>
<td>-</td>
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<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>-</td>
</tr>
<tr>
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<td>-</td>
<td>-</td>
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<tr>
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<td>0.03</td>
<td>0.001</td>
<td>0.22</td>
<td>0.04</td>
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</tbody>
</table>

Mean Error = \( \frac{1}{N} \sum_{i=1}^{N} |P_i - O_i| \)  
Mean Bias = \( \frac{1}{N} \sum_{i=1}^{N} (P_i - O_i) \)  
Fractional Error = \( \frac{2}{N} \sum_{i=1}^{N} \left| \frac{P_i - O_i}{P_i + O_i} \right| \)  
Fractional bias = \( \frac{2}{N} \sum_{i=1}^{N} \left( \frac{P_i - O_i}{P_i + O_i} \right) \)

where \( P_i \) are predicted values by PMCAMx, \( O_i \) the AERONET retrievals and \( N \) the number of stations.
Table 2. Error metrics for the evaluation of PMCAMx against MODIS monthly mean AODs.

<table>
<thead>
<tr>
<th>Region</th>
<th>Mean MODIS AOD</th>
<th>Mean PMCAMx AOD</th>
<th>Mean Error</th>
<th>Mean Bias</th>
<th>Fractional Error</th>
<th>Fractional Bias</th>
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</thead>
<tbody>
<tr>
<td>UK and Ireland</td>
<td>0.21</td>
<td>0.23</td>
<td>0.04</td>
<td>0.02</td>
<td>0.22</td>
<td>0.14</td>
</tr>
<tr>
<td>Central Europe</td>
<td>0.16</td>
<td>0.17</td>
<td>0.05</td>
<td>0.01</td>
<td>0.30</td>
<td>0.13</td>
</tr>
<tr>
<td>North Europe</td>
<td>0.09</td>
<td>0.14</td>
<td>0.06</td>
<td>0.05</td>
<td>0.53</td>
<td>0.47</td>
</tr>
<tr>
<td>Spain and Portugal</td>
<td>0.14</td>
<td>0.12</td>
<td>0.04</td>
<td>-0.03</td>
<td>0.28</td>
<td>-0.15</td>
</tr>
<tr>
<td>East Europe</td>
<td>0.13</td>
<td>0.15</td>
<td>0.05</td>
<td>0.03</td>
<td>0.35</td>
<td>0.25</td>
</tr>
<tr>
<td>Balkans</td>
<td>0.19</td>
<td>0.14</td>
<td>0.05</td>
<td>-0.04</td>
<td>0.28</td>
<td>-0.24</td>
</tr>
<tr>
<td>Russia, Belarus, and Ukraine</td>
<td>0.13</td>
<td>0.12</td>
<td>0.04</td>
<td>0.01</td>
<td>0.30</td>
<td>-0.01</td>
</tr>
<tr>
<td>Turkey and Northern Africa</td>
<td>0.16</td>
<td>0.14</td>
<td>0.05</td>
<td>-0.03</td>
<td>0.31</td>
<td>-0.13</td>
</tr>
<tr>
<td>Mediterranean Sea</td>
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<td>0.14</td>
<td>0.04</td>
<td>-0.04</td>
<td>0.25</td>
<td>-0.24</td>
</tr>
<tr>
<td>North Atlantic Ocean</td>
<td>0.17</td>
<td>0.14</td>
<td>0.04</td>
<td>-0.03</td>
<td>0.30</td>
<td>-0.21</td>
</tr>
<tr>
<td>South Atlantic Ocean</td>
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<td>0.06</td>
<td>-0.06</td>
<td>0.45</td>
<td>-0.45</td>
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<tr>
<td>Black Sea</td>
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<td>0.04</td>
<td>-0.03</td>
<td>0.22</td>
<td>-0.18</td>
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<td>0.14</td>
<td>0.05</td>
<td>-0.02</td>
<td>0.33</td>
<td>-0.16</td>
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Table 3. Monthly predicted mean ground-level concentration in $\mu g/m^3$ of the major PM$_{2.5}$ components.

<table>
<thead>
<tr>
<th>Region</th>
<th>SO$_4^{2-}$</th>
<th>OA</th>
<th>EC</th>
<th>Cl$^-$</th>
<th>Na$^+$</th>
<th>NH$_4^+$</th>
<th>NO$_3^-$</th>
<th>H$_2$O</th>
<th>Crustal</th>
</tr>
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<tr>
<td>UK and Ireland</td>
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<td>3.4</td>
<td>0.5</td>
<td>0.6</td>
<td>0.8</td>
<td>2.4</td>
<td>3.8</td>
<td>33.3</td>
<td>0.7</td>
</tr>
<tr>
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<td>3.0</td>
<td>3.4</td>
<td>0.5</td>
<td>0.2</td>
<td>0.4</td>
<td>1.4</td>
<td>1.3</td>
<td>10.1</td>
<td>0.6</td>
</tr>
<tr>
<td>North Europe</td>
<td>2.2</td>
<td>2.3</td>
<td>0.2</td>
<td>0.2</td>
<td>0.4</td>
<td>0.9</td>
<td>0.6</td>
<td>5.1</td>
<td>0.4</td>
</tr>
<tr>
<td>Spain and Portugal</td>
<td>1.6</td>
<td>1.2</td>
<td>0.2</td>
<td>0.1</td>
<td>0.2</td>
<td>0.7</td>
<td>0.5</td>
<td>11.7</td>
<td>0.3</td>
</tr>
<tr>
<td>East Europe</td>
<td>2.9</td>
<td>3.1</td>
<td>0.4</td>
<td>0.1</td>
<td>0.4</td>
<td>1.2</td>
<td>0.8</td>
<td>9.2</td>
<td>0.6</td>
</tr>
<tr>
<td>Balkans</td>
<td>3.9</td>
<td>2.7</td>
<td>0.3</td>
<td>0.1</td>
<td>0.3</td>
<td>1.4</td>
<td>0.3</td>
<td>4.9</td>
<td>0.6</td>
</tr>
<tr>
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<td>2.1</td>
<td>0.3</td>
<td>0.03</td>
<td>0.2</td>
<td>0.9</td>
<td>0.2</td>
<td>3.9</td>
<td>0.5</td>
</tr>
<tr>
<td>Turkey and</td>
<td>2.7</td>
<td>2.2</td>
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<td>0.2</td>
<td>0.4</td>
<td>1.0</td>
<td>0.6</td>
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</tr>
<tr>
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<td></td>
<td></td>
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<td></td>
<td></td>
</tr>
<tr>
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<td>1.3</td>
<td>1.4</td>
<td>1.2</td>
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<td>11.3</td>
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</tr>
<tr>
<td>North Atlantic Ocean</td>
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<td>1.0</td>
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<td>0.5</td>
<td>0.3</td>
<td>8.8</td>
<td>0.3</td>
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<tr>
<td>Black Sea</td>
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<td>2.8</td>
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<td>0.6</td>
<td>0.7</td>
<td>1.3</td>
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<td>0.6</td>
<td>0.9</td>
<td>0.5</td>
<td>10</td>
<td>0.5</td>
</tr>
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</table>
Figure 1. Geographical distribution of the 50 AERONET stations used in the present study.
Figure 2. Geographical distribution of the number of available AOD retrievals from MODIS over Europe during May 2008. White color denotes no retrievals. Land is partitioned into 8 regions including the United Kingdom and Ireland, central Europe, North Europe, Spain and Portugal, East Europe, Balkans, Russia/Belarus/Ukraine, Turkey, and Northern Africa. The sea is partitioned into 4 regions: the Mediterranean, North Atlantic, South Atlantic, and Black Sea.
Figure 3. Monthly mean AODs from PMCAMx and MODIS (QAC ≥ 1) during May 2008. White color denotes no AOD retrieval. A coarse particle rejection filter has been employed. The PMCAMx AODs correspond to the periods of the MODIS retrievals.
Figure 4. Difference of the PMCAMx from MODIS (QAC ≥ 1) monthly mean AODs during May 2008. Positive means that PMCAMx overpredicts AOD compared to MODIS. There were not enough dust-screened AOD retrievals for the model evaluation in the white areas.
Figure 5. a) Comparison of the PMCAMx predictions with MODIS (QAC ≥ 1) monthly mean AODs. The different colors indicate density. The dashed red lines denote the ocean expected error envelope and the dotted lines denote the land envelope which describes MODIS AOD uncertainties with respect to AERONET. The solid red line is the 1:1 line. b) Comparison of the PMCAMx with AERONET monthly mean AODs. The PMCAMx values correspond to the periods of measurement for the 50 AERONET stations.
Figure 6. a) Box plots of the PMCAMx and MODIS (QAC ≥ 1) monthly mean AODs for land. The central mark is the median, the edges of the box are the 25th and 75th percentiles, the whiskers extend to the extreme data points considered to be not outliers, and the outliers are plotted individually by the red marks. Points are drawn as outliers if they are larger than Q3+1.5(Q3-Q1) or smaller than Q1-1.5(Q3-Q1), where Q1 and Q3 are the first and third quartiles, respectively. b) Box plots of the PMCAMx and MODIS (QAC ≥ 1) monthly mean AODs for water.