Supporting Materials to "Modular Assessment of Rainfall-Runoff Models Toolbox (MARRMoT) v1.0: an open-source, extendable framework providing implementations of 46 conceptual hydrologic models as continuous space-state formulations"

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S1 Introduction

These Supporting Materials contain documentation for various parts of the MARRMoT software. Section S2 contains model descriptions for the 46 conceptual models included in MARRMoT. Section S3 shows how the constitutive functions of each model are translated into Matlab code, and which models use which of the resulting flux functions. Section S4 shows how 7 different Unit Hydrograph approaches are coded in MARRMoT and which models use these. Section S5 shows an overview of generalized parameter ranges for the 46 models.
S2 Model descriptions

This section contains mathematical descriptions of all models that are included in the Modular Assessment of Rainfall-Runoff Models Toolbox v1.0 (MARRMoT). All descriptions follow the same layout (see the example model at the end of this section):

- **Title:** gives an informal name for the model structure followed by a unique ID;
- **Introduction:** gives a brief description of the model, including one or more original reference(s), the number of stores and parameters, a list containing parameter names and occasionally note-worthy deviations from the original model;
- **Process list:** a brief overview of the main processes the model is intended to represent;
- **Figure:** a wiring diagram that shows the names of model stores and fluxes;
- **Matlab name section:** gives the name of the file that contains Matlab code for this model;
- **Model equations section:** a mathematical description of the model. This uses Ordinary Differential Equations (ODEs) to describe the changes in model storage(s) and constitutive functions that detail how individual fluxes operate.

MARRMoT models intend to stay close to the original models they are based on but differences are unavoidable. We strongly recommend users to read the original paper cited for each model as well as our interpretation given in this document. In many cases, more than one version of a model exists, but these are not always easily distinguishable. There is a certain degree of model name equifinality, where a single name is used to refer to various different version of the same base model. A good example is TOPMODEL, of which many variants exist based around the initial concept of topographic indices. MARRMoT models tend to be based on older rather than newer publications for any given model (to stay close to the "intended" model by the original author(s)) but our selection has been pragmatic to achieve greater variety in the available fluxes and model structures in MARRMoT. The description of each model lists the papers that form the basis of the MARRMoT version of that model.

MARRMoT is set up to work with arbitrary user-defined time step sizes for climate input data. For consistency of parameter values across different time step sizes, the internal dynamics of each model are specified using the base units \( \text{mm} \) and \( \text{d} \). The temporal resolution of climate data is converted to \( \text{mm/d} \) within each model, and model output is converted back to the user-specified time step size. Internal fluxes in each MARRMoT model use the base units and are in \( \text{mm/d} \) and parameter values are specified in the base or derived units (e.g. \( \text{d}^{-1} \) for time coefficients). These units are kept throughout this document.

The computational implementation of constitutive functions is given in section S3 and Unit Hydrographs are specified in section S4. Generalized parameter ranges for all models are given in section S5.
Example model (model ID: nn)

The Example model (fig. S1) is used in the MARRMoT User Manual to show how to create a new MARRMoT model from scratch (Knoben et al., 2018). It has 3 stores and 7 parameters ($UZ_{max}$, $c_{rate}$, $p_{rate}$, $k_lz$, $\alpha$, $k_g$, $d$). The model aims to represent:

- Saturation excess from the upper zone;
- Two-way interaction between upper and lower zone through percolation and capillary rise;
- A split between fast subsurface flow and groundwater recharge from the lower zone;
- Slow runoff from the groundwater;
- Triangular routing of combined surface and subsurface flows.

File names

Model: m_nn_example_7p_3s
Parameter ranges: m_nn_example_7p_3s_parameter ranges

Model equations

\[
\frac{dUZ}{dt} = P + q_c - E - q_{se} - q_p \tag{1}
\]
\[
E = E_p \times \frac{UZ}{UZ_{max}} \tag{2}
\]
\[
q_c = c_{rate} \left(1 - \frac{UZ}{UZ_{max}}\right) \tag{3}
\]
\[
q_{se} = \begin{cases} 
P, & \text{if } UZ = UZ_{max} \\
0, & \text{otherwise} 
\end{cases} \tag{4}
\]
\[
q_p = p_{rate} \tag{5}
\]

Where $UZ$ [mm] is the current storage in the upper zone, refilled by precipitation $P$ [mm/d] and capillary rise $q_c$ [mm/d] and drained by evaporation $E$ [mm/d], percolation $q_p$ [mm/d] and saturation excess $q_{se}$ [mm/d]. Evaporation occurs at the potential rate $E_p$ scaled by the current storage in $UZ$ compared to maximum storage $UZ_{max}$ [mm]. Capillary rise occurs at a maximum rate $c_{rate}$ [mm/d] if $UZ = 0$ and decreases linearly if not. Saturation excess flow only occurs when $UZ$ is at maximum capacity.

Figure S1: Structure of the Example model
Percolation occurs at a constant rate \( p_{rate} \) [mm/d].

\[
\frac{dLZ}{dt} = q_p - q_c - q_{lz} \tag{6}
\]

\[
q_{lz} = k_{lz} \times LZ \tag{7}
\]

\[
q_{lz} = k_{lz} \times LZ \tag{8}
\]

Where \( LZ \) [mm] is the current storage in the lower zone, refilled by percolation \( q_p \) [mm/d] and drained by capillary rise \( q_c \) [mm/d] and outflow \( q_{lz} \) [mm/d]. Outflow has a linear relation with storage through time parameter \( k_{lz} \) [d\(^{-1}\)].

\[
\frac{dG}{dt} = q_g - q_s \tag{9}
\]

\[
q_g = \alpha \times q_{lz} \tag{10}
\]

\[
q_s = k_g \times G \tag{11}
\]

Where \( G \) [mm] is the current groundwater storage, refilled by recharge \( q_g \) [mm/d] and drained by slow flow \( q_s \) [mm/d]. Recharge is a fraction \( \alpha \) [\(-\)] of outflow from the lower zone. Outflow has a linear relation with storage through time parameter \( k_g \) [d\(^{-1}\)]. Saturation excess \( q_{se} \), interflow \( q_I \) and slow flow \( q_s \) are combined and routed with a triangular Unit Hydrograph with time base \( d \) [d] to give outflow \( Q \) [mm/d].
S2.1 Collie River Basin 1 (model ID: 01)

The Collie River Basin 1 model (fig. S2) is part of a top-down modelling exercise and is originally applied at the annual scale (Jothityangkoon et al., 2001). This is a classic bucket model. It has 1 store and 1 parameter \( S_{\text{max}} \). The model aims to represent:

- Evaporation from soil moisture;
- Saturation excess surface runoff.

S2.1.1 File names

Model: m_01_collie1_ls_1p
Parameter ranges: m_01_collie1_ls_1p_parameter_ranges

S2.1.2 Model equations

\[
\frac{dS}{dt} = P - E_a - Q_{se} \tag{12}
\]
\[
E_a = \frac{S}{S_{\text{max}}} \cdot E_p \tag{13}
\]
\[
Q_{se} = \begin{cases} P, & \text{if } S > S_{\text{max}} \\ 0, & \text{otherwise} \end{cases} \tag{14}
\]

Where \( S [\text{mm}] \) is the current storage in the soil moisture and \( P \) the precipitation input \([\text{mm/d}]\). Actual evaporation \( E_a [\text{mm/d}] \) is estimated based on the current storage \( S \), the maximum soil moisture storage \( S_{\text{max}} [\text{mm}] \), and the potential evapotranspiration \( E_p [\text{mm/d}] \). \( Q_{se} [\text{mm/d}] \) is saturation excess overland flow.
S2.2 Wetland model (model ID: 02)

The Wetland model (fig. S3) is a conceptualization of the perceived dominant processes in a typical Western European wetland (Savenije, 2010). It belongs to a 3-part topography driven modelling exercise, together with a hillslope and plateau conceptualization. Each model is provided in isolation here, because they are well-suited for isolating specific model structure choices. It has 1 store and 4 parameters ($D_w$, $S_{w,max}$, $\beta_w$ and $K_w$). The model aims to represent:

- Stylized interception by vegetation;
- Evaporation;
- Saturation excess runoff generated from a distribution of soil depths;
- A linear relation between storage and slow runoff.

S2.2.1 File names

Model: m_02_wetland_4p_1s
Parameter ranges: m_02_wetland_4p_1s_parameter_ranges

S2.2.2 Model equations

\[
\frac{dS_w}{dt} = P_e - E_w - Q_{w,sof} - Q_{w,gw} \tag{15}
\]
\[
P_e = \max(P - D_w, 0) \tag{16}
\]
\[
E_w = \begin{cases} E_p, & \text{if } S_w > 0 \\ 0, & \text{otherwise} \end{cases} \tag{17}
\]
\[
Q_{w,sof} = \left(1 - \left(1 - \frac{S_w}{S_{w,max}}\right)^{\beta_w}\right) \times P_e \tag{18}
\]
\[
Q_{w,gw} = K_w \times S_w \tag{19}
\]

Where $S_w$ is the current soil water storage [mm]. Incoming precipitation $P$ [mm/d] is reduced by interception $D_w$ [mm/d], which is assumed to evaporate before the next precipitation event. Evaporation from soil moisture $E_w$ [mm/d] occurs at the potential rate $E_p$ whenever possible. Saturation excess surface runoff $Q_{w,sof}$ [mm/d] depends on the fraction of the catchment that is currently saturated, expressed through parameters $S_{w,max}$ [mm] and $\beta_w$ [-]. Groundwater flow $Q_{w,gw}$ [mm/d] depends linearly on current storage $S_w$ through parameter $K_w$ [d$^{-1}$]. Total flow:

\[
Q = Q_{w,sof} + Q_{w,gw} \tag{20}
\]

Figure S3: Structure of the Wetland model
S2.3 Collie River Basin 2 (model ID: 03)

The Collie River Basin 2 model (fig. S4) is part of a top-down modelling exercise and is originally applied at the monthly scale (Jothityangkoon et al., 2001). It has 1 store and 4 parameters ($S_{\text{max}}$, $S_{fc}$, $a$, $M$). The model aims to represent:

- Separate bare soil and vegetation evaporation;
- Saturation excess surface runoff;
- Subsurface runoff.

S2.3.1 File names

Model: m_03_collie2_4p_2s
Parameter ranges: m_03_collie2_4p_2s_parameter_ranges

S2.3.2 Model equations

\[
\frac{dS}{dt} = P - E_b - E_v - Q_{se} - Q_{ss} \quad \text{(21)}
\]

\[
E_b = \frac{S}{S_{\text{max}}} (1 - M) \cdot E_p \quad \text{(22)}
\]

\[
E_v = \begin{cases} 
M \cdot E_p, & \text{if } S > S_{fc} \\
\frac{S}{S_{fc}} M \cdot E_p, & \text{otherwise}
\end{cases} \quad \text{(23)}
\]

\[
Q_{se} = \begin{cases} 
P, & \text{if } S > S_{\text{max}} \\
0, & \text{otherwise}
\end{cases} \quad \text{(24)}
\]

\[
Q_{ss} = \begin{cases} 
a \cdot (S - S_{fc}), & \text{if } S > S_{fc} \\
0, & \text{otherwise}
\end{cases} \quad \text{(25)}
\]

Where $S$ [mm] is the current storage in the soil moisture and $P$ [mm/d] the precipitation input. Actual evaporation is split between bare soil evaporation $E_b$ [mm/d] and transpiration through vegetation $E_v$ [mm/d], controlled through the forest fraction $M$ [-]. The evaporation estimates are based on the current storage $S$, the potential evapotranspiration $E_p$ [mm/d], maximum soil moisture storage $S_{\text{max}}$ [mm] and field capacity $S_{fc}$ [mm] respectively. $Q_{se}$ [mm/d] is saturation excess overland flow. $Q_{ss}$ [mm/d] is subsurface flow regulated by runoff coefficient $a$ [d$^{-1}$]. Total flow:

\[
Q = Q_{se} + Q_{ss} \quad \text{(26)}
\]
S2.4 New Zealand model v1 (model ID: 04)

The New Zealand model v1 (fig. S5) is part of a top-down modelling exercise that focusses on several catchments in New Zealand (Atkinson et al., 2002). It has 1 store and 6 parameters ($S_{\text{max}}$, $S_{fc}$, $M$, $a$, $b$ and $t_{c,bf}$). The model aims to represent:

- Separate vegetation and bare soil evaporation;
- Saturation excess overland flow;
- Subsurface runoff when soil moisture exceeds field capacity;
- Baseflow.

S2.4.1 File names

Model: m_04_newzealand1_6p_1s
Parameter ranges: m_04_newzealand1_6p_1s_parameter_ranges

S2.4.2 Model equations

\[
\frac{dS_m}{dt} = P - E_{\text{veg}} - E_{bs} - Q_{se} - Q_{ss} - Q_{bf} \tag{27}
\]

\[
E_{\text{veg}} = \begin{cases} 
M \ast E_p, & \text{if } S > S_{fc} \\
\frac{S_m}{S_{fc}} \ast M \ast E_p, & \text{otherwise}
\end{cases} \tag{28}
\]

\[
E_{bs} = \frac{S}{S_{\text{max}}} (1 - M) \ast E_p \tag{29}
\]

\[
Q_{se} = \begin{cases} 
P, & \text{if } S \geq S_{\text{max}} \\
0, & \text{otherwise}
\end{cases} \tag{30}
\]

\[
Q_{ss} = \begin{cases} 
(a \ast (S - S_{fc}))^b, & \text{if } S \geq S_{fc} \\
0, & \text{otherwise}
\end{cases} \tag{31}
\]

\[
Q_{bf} = t_{c,bf} \ast S \tag{32}
\]

Where $S_m$ [mm] is the current soil moisture storage which gets replenished through precipitation $P$ [mm/d]. Evaporation through vegetation $E_{\text{veg}}$ [mm/d] depends on the forest fraction $M$ [-] and field capacity $S_{fc}$ [-]. $E_{bs}$ [mm/d] represents bare soil evaporation. When $S$ exceeds the maximum storage $S_{\text{max}}$ [mm], water leaves the model as saturation excess runoff $Q_{se}$. If S exceeds field capacity $S_{fc}$ [mm], subsurface runoff $Q_{ss}$ [mm/d] is generated controlled by time parameter $a$ [$d^{-1}$] and nonlinearity parameter $b$ [-]. $Q_{bf}$ represents baseflow controlled by time scale parameter $t_{c,bf}$ [$d^{-1}$]. Total runoff $Q_t$ [mm/d] is:

\[
Q_t = Q_{se} + Q_{ss} + Q_{bf} \tag{33}
\]
S2.5 IHACRES (model ID: 05)

The IHACRES model (fig. S6) as implemented here is a modification of the original equations (Littlewood et al., 1997; Ye et al., 1997; Croke and Jakeman, 2004), which explicitly account for the various fluxes in a step-wise order. Furthermore, IHACRES usually uses temperature as a proxy for potential evapotranspiration \((E_p)\). Here it uses estimated \(E_p\) directly to be consistent with other models. The equations for \(E_a\) and \(U\) are set up following Croke and Jakeman (2004), with the non-linearity in \(U\) based on Ye et al. (1997). This version thus uses a catchment moisture deficit formulation, rather than a catchment wetness index. Littlewood et al. (1997) recommends the two parallel routing functions. The model has 1 deficit store and 6 parameters \((lp, d, p, \alpha, \tau_q, \tau_s)\). The model aims to represent:

- Catchment deficit build-up
- Slow and fast routing of effective precipitation.

S2.5.1 File names

Model: m_05_ihacres_6p_1s
Parameter ranges: m_05_ihacres_6p_1s_parameter_ranges

S2.5.2 Model equations

\[
\frac{dCMD}{dt} = -P + E_a + U \quad (34)
\]

\[
E_a = E_p \times \min\left(1, e^{2(1-\frac{CMD}{lp})}\right) \quad (35)
\]

\[
U = P \left(1 - \min\left(1, \left(\frac{CMD}{d}\right)^p\right)\right) \quad (36)
\]

\[
U_q = \alpha \times U \quad (37)
\]

\[
U_s = (1 - \alpha) \times U \quad (38)
\]

Where \(CMD\) is the current moisture deficit \([\text{mm}]\), \(P\) \([\text{mm/d}]\) the incoming precipitation that reduces the deficit, \(E_a\) \([\text{mm/d}]\) evaporation that increases the deficit, and \(U\) \([\text{mm/d}]\) the effective precipitation that occurs when the deficit is below a threshold \(d\) \([\text{mm}]\).

Evaporation occurs at the potential rate \(E_p\) until the moisture deficit reaches wilting point \(lp\) \([\text{mm}]\), after which evaporation decreases exponentially with increasing deficit. Effective precipitation \(U\) equals incoming precipitation \(P\) when the deficit is zero, and decreases as a linear fraction of \(P\) until moisture deficit is larger than a threshold \(d\) \([\text{mm}]\), after which precipitation does not contribute to streamflow any longer. \(U\) is divided between fast and slow routing components based on fraction \(\alpha\).
Both routing schemes are exponentially decreasing over time with lags $\tau_q$ [d] and $\tau_s$ [d] respectively. Total flow:

$$Q = x_q + x_s$$  \hspace{1cm} (39)
**S2.6 Alpine model v1 (model ID: 06)**

The Alpine model v1 model (fig. S7) is part of a top-down modelling exercise and represents a monthly water balance model (Eder et al., 2003). It has 2 stores and 4 parameters ($T_t$, $ddf$, $S_{max}$, $t_c$). The model aims to represent:

- Snow accumulation and melt;
- Saturation excess overland flow;
- Linear subsurface runoff.

**S2.6.1 File names**

Model: m_06_alpine1_4p_2s

Parameter ranges: m_06_alpine1_4p_2s_parameter_ranges

**S2.6.2 Model equations**

\[
\frac{dS_n}{dt} = P_s - Q_N \tag{40}
\]

\[
P_s = \begin{cases} 
P, & \text{if } T \leq T_t \\ 
0, & \text{otherwise} \end{cases} \tag{41}
\]

\[
Q_N = \begin{cases} 
ddf \times (T - T_t), & \text{if } T \geq T_t \\ 
0, & \text{otherwise} \end{cases} \tag{42}
\]

Where $S_N$ is the current snow storage [mm], $P_s$ the precipitation that falls as snow [mm/d], $Q_N$ snow melt [mm/d] based on a degree-day factor ($ddf$, [mm/°C/d]) and threshold temperature for snowfall and snowmelt ($T_t$, [°C]).

Figure S7: Structure of the Alpine model v1
\[
\frac{dS_m}{dt} = P_r + Q_N - E_a - Q_{se} - Q_{ss} \quad (43)
\]

\[
P_r = \begin{cases} 
P, & \text{if } T > T_i \\ 0, & \text{otherwise} \end{cases} \quad (44)
\]

\[
E_a = \begin{cases} 
E_p, & \text{if } S > 0 \\ 0, & \text{otherwise} \end{cases} \quad (45)
\]

\[
Q_{se} = \begin{cases} 
P_r + Q_N, & \text{if } S_m \geq S_{max} \\ 0, & \text{otherwise} \end{cases} \quad (46)
\]

\[
Q_{ss} = t_c * S_m \quad (47)
\]

Where \( S_m \) [mm] is the current soil moisture storage, which is assumed to evaporate at the potential rate \( E_p \) [mm/d] when possible. When \( S_m \) exceeds the maximum storage \( S_{max} \) [mm], water leaves the model as saturation excess runoff \( Q_{se} \). \( Q_{ss} \) represents subsurface flow controlled by time scale parameter \( t_c \) [d\(^{-1}\)]. Total runoff \( Q_t \) [mm/d] is:

\[
Q_t = Q_{se} + Q_{ss} \quad (48)
\]
S2.7 GR4J (model ID: 07)

The GR4J model (fig. S8) is originally developed with an explicit (operator-splitting) time-stepping scheme (Perrin et al., 2003). Recently a new version has been released that works with an implicit time-stepping scheme (Santos et al., 2017). The implementation given here follows most of the equations from Santos et al. (2017), but uses the original Unit Hydrographs for flood routing given by Perrin et al. (2003). It has 2 stores and 4 parameters \((x_1, x_2, x_3, x_4)\). The model aims to represent:

- Implicit interception by vegetation, expressed as net precipitation or evaporation;
- Different time delays within the catchment expressed by two hydrographs;
- Water exchange with neighbouring catchments.

S2.7.1 File names

Model: m_07_gr4j_4p_2s
Parameter ranges: m_07_gr4j_4p_2s_parameter_ranges

S2.7.2 Model equations

\[
\frac{dS}{dt} = P_s - E_s - Perc \\ (49)
\]

\[
P_s = P_n * \left( 1 - \left( \frac{S}{x_1} \right)^2 \right) \\ (50)
\]

\[
P_n = \begin{cases} P - Ep, & \text{if } P \geq Ep \\ 0, & \text{otherwise} \end{cases} \\ (51)
\]

\[
E_s = E_n * \left( 2 \frac{S}{x_1} - \left( \frac{S}{x_1} \right)^2 \right) \\ (52)
\]

\[
E_n = \begin{cases} Ep - P, & \text{if } Ep > P \\ 0, & \text{otherwise} \end{cases} \\ (53)
\]

\[
Perc = \frac{x_1^{-4}}{4d} * \left( \frac{4}{5} \right)^{-4} S^5 \\ (54)
\]

Where \(S\) is the current soil moisture storage [mm], \(P_s\) [mm/d] is the fraction of net precipitation \(P_n\) [mm/d] redirected to soil moisture, \(E_s\) [mm/d] is the fraction of net evaporation \(E_n\) [mm/d] subtracted from soil moisture, and \(perc\) [mm/d] is percolation to deeper soil layers. Parameter \(x_1\) [mm] is the maximum soil moisture storage.
Percolation $\text{perc}$ and excess precipitation $P_n - P_r$ are divided into 90% groundwater flow, routed through a triangular routing scheme with time base $x_4 [d]$, and 10% direct runoff, routed through a triangular routing scheme with time base $2x_4 [d]$.

\[
\frac{dR}{dt} = Q_9 + F(x_2) - Q_r
\]

\[
F(x_2) = x_2 \times \left( \frac{R}{x_3} \right)^{3.5}
\]

\[
Q_r = \frac{x_3}{4d} R^5
\]

Where $R [\text{mm}]$ is the current storage in the routing store, $F(x_2) [\text{mm/d}]$ the catchment groundwater exchange, depending on exchange coefficient $x_2 [\text{mm/d}]$ and the maximum routing capacity $x_3 [\text{mm}]$, and $Q_r [\text{mm/d}]$ routed flow. Total runoff $Q_t [\text{mm/d}]$:

\[
Q_t = Q_r + \max(Q_1 + F(x_2), 0)
\]
S2.8 United States model (model ID: 08)

The United States model (fig. S9) is part of a multi-model comparison study using several catchments in the United States (Bai et al., 2009). It has 2 stores and 5 parameters ($\alpha_{ei}$, $M$, $S_{max}$, $f_c$, $\alpha_{ss}$). The model aims to represent:

- Interception as a percentage of precipitation;
- Separate unsaturated and saturated zones;
- Separate bare soil evaporation and vegetation transpiration;
- Saturation excess overland flow;
- Subsurface flow.

S2.8.1 File names

Model: m_08_us1_5p_2s
Parameter ranges: m_08_us1_5p_2s_parameter_ranges

S2.8.2 Model equations

\[
dS_{us} = P - E_{us,ei} - E_{us,veg} - E_{us,bs} - r_g \tag{59}
\]

$E_{us,ei} = \alpha_{ei} \times P \tag{60}$

\[
E_{us,veg} = \begin{cases} 
\frac{S_{us}}{S_{us} + S_{sat}} \times M \times E_p, & \text{if } S_{us} > S_{usfc} \\
\frac{S_{us}}{S_{us} + S_{sat}} \times M \times E_p \times \frac{S_{us}}{S_{usfc}}, & \text{otherwise} 
\end{cases} \tag{61}
\]

$E_{us,bs} = \frac{S_{us}}{S_{us} + S_{sat}} \times (1 - M) \times \frac{S_{us}}{S_{max} - S_{sat}} \times E_p \tag{62}$

$rg = \begin{cases} 
P, & \text{if } S_{us} > S_{usfc} \\
0, & \text{otherwise} \end{cases} \tag{63}$

$S_e = \begin{cases} 
S_{us} - S_{usfc}, & \text{if } S_{us} > S_{usfc} \\
0, & \text{otherwise} \end{cases} \tag{64}$

$S_{usfc} = f_c \times (S_{max} - S_{sat}) \tag{65}$

Where $S_{us}$ [mm] is the current storage in the unsaturated zone, $E_{us,ei}$ [mm/d] evaporation from interception, $E_{us,veg}$ [mm/d] transpiration through vegetation, $E_{us,bs}$ [mm/d] bare soil evaporation and $r_g$ [mm/d] drainage to the saturated zone. Interception evaporation relies on parameter $\alpha_{ei}$ [-], representing the fraction of precipitation $P$ that is intercepted. The implicit assumption is that this evaporates before the next
precipitation event. Transpiration uses forest fraction $M$ [-], potential evapotranspiration $E_p$ [mm/d] and the estimated field capacity $S_{usfc}$ through parameter $fc$ [-]. Bare soil evaporation relies also on the maximum soil moisture storage $S_{max}$ [mm].

\[
\frac{dS_{sat}}{dt} = r_g - E_{sat,\text{veg}} - E_{sat,bs} - Q_{se} - Q_{ss}
\]  

(66)

\[
E_{sat,\text{veg}} = \frac{S_{sat}}{S_{max}} \times M \times E_p
\]  

(67)

\[
E_{sat,bs} = \frac{S_{sat}}{S_{max}} \times (1 - M) \times E_p
\]  

(68)

\[
Q_{se} = \begin{cases} 
  r_g, & \text{if } S_{us} \geq S_{max} \\
  0, & \text{otherwise}
\end{cases}
\]  

(69)

\[
Q_{ss} = \alpha_{ss} \times S_{sat}
\]  

(70)

Where $S_{sat}$ [mm] is the current storage in the saturated zone, $E_{sat,\text{veg}}$ [mm/d] transpiration through vegetation, $E_{sat,bs}$ [mm/d] bare soil evaporation, $Q_{se}$ [mm/d] saturation excess overland flow and $Q_{ss}$ [mm/d] subsurface flow. Subsurface flow uses time parameter $\alpha_{ss}$ [d$^{-1}$] Total flow:

\[
Q = Q_{se} + Q_{ss}
\]  

(71)
S2.9 Susannah Brook model v1-5 (model ID: 09)

The Susannah Brook model v1-5 (fig. S10) is part of a top-down modelling exercise designed to use auxiliary data (Son and Sivapalan, 2007). It has 2 stores and 6 parameters ($S_b$, $S_{fc}$, $M$, $a$, $b$ and $r$). The model aims to represent:

- Evaporation from soil and transpiration from vegetation;
- Saturation excess and non-linear subsurface flow;
- Groundwater recharge and baseflow.

S2.9.1 File names

Model:  m_09_susannah1_6p_2s
Parameter ranges:  m_09_susannah1_6p_2s_parameter_ranges

S2.9.2 Model equations

\[
\frac{dS_{uz}}{dt} = P - E_{bs} - E_{veg} - Q_{se} - Q_{ss} \tag{72}
\]

\[
E_{bs} = \frac{S}{S_b} (1 - M) E_p \tag{73}
\]

\[
E_{veg} = \begin{cases} 
M \ast E_p, & \text{if } S > S_{fc} \\
\frac{S}{S_{fc}} M \ast E_p, & \text{otherwise}
\end{cases} \tag{74}
\]

\[
Q_{se} = \begin{cases} 
P, & \text{if } S \geq S_b \\
0, & \text{otherwise}
\end{cases} \tag{75}
\]

\[
Q_{ss} = \begin{cases} 
\left(\frac{(S-S_{fc})}{a}\right)^{\frac{1}{b}}, & \text{if } S > S_{fc} \\
0, & \text{otherwise}
\end{cases} \tag{76}
\]

Where $S_{uz}$ is current storage in the upper zone [mm]. $P$ [mm/d] is the precipitation input. $E_{bs}$ is bare soil evaporation [mm/d] based on soil depth $S_b$ [mm] and forest fraction $M$. $E_{veg}$ is transpiration from vegetation, using the wilting point $S_{fc}$ [mm] and forest fraction $M$. $Q_{se}$ is saturation excess flow [mm/d]. $Q_{ss}$ is non-linear subsurface flow, using the wilting point $S_{fc}$ as a threshold for flow generation and two flow parameters $a$ [d] and $b$ [-]. $Q_r$ is groundwater recharge [mm/d].
\[
\frac{DS_{gw}}{dt} = Q_r - Q_b \tag{77}
\]

\[
Q_r = r \ast Q_{ss} \tag{78}
\]

\[
Q_b = \left(\frac{1}{a} S_{gw}\right)^\frac{1}{b} \tag{79}
\]

Where \( S_{gw} \) is the groundwater storage \([\text{mm}]\), and \( Q_b \) the baseflow flux \([\text{mm/d}]\).

Total flow \([\text{mm}]\):

\[
Q = Q_{se} + (Q_{ss} - Q_r) + Q_b \tag{80}
\]
S2.10 Susannah Brook model v2 (model ID: 10)

The Susannah Brook model v2 model (fig. S11) is part of a top-down modelling exercise designed to use auxiliary data (Son and Sivapalan, 2007). It has 2 stores and 6 parameters ($S_b, \phi, f_c, r, c, d$). For consistency with other model formulations, $S_b$ is used as a parameter, instead of being broken down into its constitutive parts $D$ and $\phi$. The model aims to represent:

- Separation of saturated zone and a variable-size unsaturated zone;
- Evaporation from unsaturated and saturated zones;
- Saturation excess and non-linear subsurface flow;
- Deep groundwater recharge.

S2.10.1 File names

Model: m_10_susannah2_6p_2s
Parameter ranges: m_10_susannah2_6p_2s_parameter_ranges

S2.10.2 Model equations

\begin{align*}
\frac{dS_{us}}{dt} &= P - E_{us} - r_g - S_e \quad (81) \\
E_{us} &= \frac{S_{us}}{S_b} \times E_p \quad (82) \\
S_b &= D \times \phi \quad (83) \\
r_g &= \begin{cases} P, & \text{if } S_{us} > S_{usfc} \\ 0, & \text{otherwise} \end{cases} \quad (84) \\
S_e &= \begin{cases} S_{us} - S_{usfc}, & \text{if } S_{us} > S_{usfc} \\ 0, & \text{otherwise} \end{cases} \quad (85) \\
S_{usfc} &= (S_b - S_{sat}) \times \frac{f_c}{\phi} \quad (86)
\end{align*}

Where $S_{us}$ is the current storage in the unsaturated store [mm], $P$ the current precipitation [mm], $S_b$ [mm] the maximum storage of the soil profile, based on the soil depth $D$ [mm] and the porosity $\phi$ [-]. $r_g$ is drainage from the unsaturated store to the saturated store [mm], based on the variable field capacity $S_{usfc}$ [mm]. $S_{usfc}$ is based on the current storage on the saturated zone $S_{sat}$ [mm], the maximum soil moisture storage $S_b$ [mm], the field capacity $f_c$ [-] and the porosity $\phi$ [-]. $S_e$ [mm]
is the storage excess, resulting from a decrease of $S_{usfc}$ that leads to more water being stored in the unsaturated zone than should be possible.

\[
\frac{dS_{sat}}{dt} = r_g - E_{sat} - Q_{SE} - Q_{SS} - Q_R \tag{87}
\]

\[
E_{sat} = \frac{S_{sat}}{S_b} * E_p \tag{88}
\]

\[
Q_{SE} = \begin{cases} 
  r_g + S_e, & \text{if } S_{sat} > S_b \\
  0, & \text{otherwise}
\end{cases} \tag{89}
\]

\[
Q_{SS} = (1 - r) * c * (S_{sat})^d \tag{90}
\]

\[
Q_R = r * c * (S_{sat})^d \tag{91}
\]

Where $S_{sat}$ is the current storage in the saturated zone [mm], $E_{sat}$ is the evaporation from the saturated zone [mm], $Q_{SE}$ saturation excess runoff [mm] that occurs when the saturated zone reaches maximum capacity $S_b$ [mm], $Q_{SS}$ is subsurface flow [mm] and $Q_R$ is recharge of deep groundwater [mm]. Both $Q_{SS}$ and $Q_R$ are based on the dimensionless fraction $r$ and subsurface flow constants $c [d^{-1}]$ and $d [-]$. Total runoff is the sum of $Q_{SE}$ and $Q_{SS}$:

\[
Q = Q_{SE} + Q_{SS} \tag{92}
\]
S2.11 Collie River Basin 3 (model ID: 11)

The Collie River Basin 3 model (fig. S12) is part of a top-down modelling exercise and is originally applied at the daily scale (Jothityangkoon et al., 2001). It has 2 stores and 6 parameters ($S_{\text{max}}, S_{fc}, a, M, b, \lambda$). The model aims to represent:

- Separate bare soil and vegetation evaporation;
- Saturation excess surface runoff;
- Non-linear subsurface runoff;
- Non-linear groundwater runoff.

S2.11.1 File names

Model: m_11_collie3_6p_2s
Parameter ranges: m_11_collie3_6p_2s_parameter_ranges

S2.11.2 Model equations

\[
\begin{align*}
\frac{dS}{dt} &= P - E_b - E_v - Q_{se} - Q_{ss} \quad (93) \\
E_b &= \frac{S}{S_{\text{max}}} (1 - M) * E_p \quad (94) \\
E_v &= \begin{cases} 
M * E_p, & \text{if } S > S_{fc} \\
\frac{s}{S_{fc}} * M * E_p, & \text{otherwise} 
\end{cases} \quad (95) \\
Q_{se} &= \begin{cases} 
P, & \text{if } S > S_{\text{max}} \\
0, & \text{otherwise} 
\end{cases} \quad (96) \\
Q_{ss} &= \begin{cases} 
(a * (S - S_{fc}))^b, & \text{if } S > S_{fc} \\
0, & \text{otherwise} 
\end{cases} \quad (97)
\end{align*}
\]

Figure S12: Structure of the Collie River Basin 3 model

Where $S$ [mm] is the current storage in the soil moisture and $P$ the precipitation input [mm/d]. Actual evaporation is split between bare soil evaporation $E_b$ [mm/d] and transpiration through vegetation $E_v$ [mm/d], controlled through the forest fraction $M$. The evaporation estimates are based on the current storage $S$, the potential evapotranspiration $E_p$ [mm/d] and the maximum soil moisture storage $S_{\text{max}}$ [mm], and field capacity $S_{fc}$ [mm] respectively. $Q_{se}$ [mm/d] is saturation excess overland flow. $Q_{ss}$ [mm/d] is non-linear subsurface flow regulated by runoff coefficients $a$ [d$^{-1}$] and $b$ [-].
\[
\frac{dG}{dt} = Q^*_{ss} - Q_{sg} \tag{98}
\]
\[
Q^*_{ss} = \lambda \ast Q_{ss} \tag{99}
\]
\[
Q_{sg} = (a \ast G)^b \tag{100}
\]

Where \( G \, [\text{mm}] \) is groundwater storage. \( Q^*_{ss} \, [\text{mm/d}] \) is the fraction of \( Q_{ss} \) directed to groundwater. \( Q_{sg} \, [\text{mm/d}] \) is non-linear groundwater flow that relies on the same parameters as subsurface flow uses. Total runoff:

\[
Q = Q_{se} + (1 - \lambda) \ast Q_{ss} + Q_{sg} \tag{101}
\]
**S2.12 Alpine model v2 (model ID: 12)**

The Alpine model v2 (fig. S13) is part of a top-down modelling exercise and represents a daily water balance model (Eder et al., 2003). It has 2 stores and 6 parameters ($T_t$, $ddf$, $S_{max}$, $C_{fc}$, $t_{c,in}$, $t_{c,bf}$). The model aims to represent:

- Snow accumulation and melt;
- Saturation excess overland flow;
- Linear subsurface runoff.

**S2.12.1 File names**
- Model: m_12_alpine2_6p_2s
- Parameter ranges: m_12_alpine2_6p_2s_parameter_ranges

**S2.12.2 Model equations**

\[
\frac{dS_n}{dt} = P_s - Q_N \\
\]

\[
P_s = \begin{cases} 
    P, & \text{if } T \leq T_t \\
    0, & \text{otherwise}
\end{cases}
\]

\[
Q_N = \begin{cases} 
    ddf \times (T - T_t), & \text{if } T \geq T_t \\
    0, & \text{otherwise}
\end{cases}
\]

Where $S_n$ is the current snow storage [mm], $P_s$ the precipitation that falls as snow [mm/d], $Q_N$ snow melt [mm/d] based on a degree-day factor (ddf, [mm/°C/d]) and threshold temperature for snowfall and snowmelt ($T_t$, [°C]).
\[
\frac{dS}{dt} = Pr + Q_N - E_a - Q_{se} - Q_{in} - Q_{bf}
\]  
(105)

\[
Pr = \begin{cases} 
P, & \text{if } T > TT \\
0, & \text{otherwise}
\end{cases}
\]  
(106)

\[
E_a = \begin{cases} 
E_p, & \text{if } S > 0 \\
0, & \text{otherwise}
\end{cases}
\]  
(107)

\[
Q_{se} = \begin{cases} 
Pr + Q_N, & \text{if } S \geq S_{max} \\
0, & \text{otherwise}
\end{cases}
\]  
(108)

\[
Q_{in} = \begin{cases} 
t_{c,in} \cdot (S - S_{fc}), & \text{if } S > S_{fc} \\
0, & \text{otherwise}
\end{cases}
\]  
(109)

\[
Q_{bf} = t_{c,bf} \cdot S
\]  
(110)

Where \( S \) [mm] is the current soil moisture storage, which is assumed to evaporate at the potential rate \( E_p \) [mm/d] when possible. When \( S \) exceeds the maximum storage \( S_{max} \) [mm], water leaves the model as saturation excess runoff \( Q_{se} \). If \( S \) exceeds field capacity \( S_{fc} \) [mm], interflow \( Q_{in} \) [mm/d] is generated controlled by time parameter \( t_{c,in} \) [d\(^{-1}\)]. \( Q_{bf} \) represents baseflow controlled by time scale parameter \( t_{c,bf} \) [d\(^{-1}\)]. Total runoff \( Q_t \) [mm/d] is:

\[
Q_t = Q_{se} + Q_{in} + Q_{bf}
\]  
(111)
S2.13 Hillslope model (model ID: 13)

The Hillslope model (fig. S14) is a conceptualization of the perceived dominant processes in a typical Western European hillslope (Savenije, 2010). It belongs to a 3-part topography driven modelling exercise, together with a wetland and plateau conceptualization. Each model is provided in isolation here, because they are well-suited for isolating specific model structure choices. It has 2 store and 7 parameters ($D_w$, $S_{h,max}$, $\beta_h$, $a$, $T_h$, $C$ and $K_h$). The model aims to represent:

- Stylized interception by vegetation;
- Evaporation;
- Separation between rapid subsurface flow and groundwater recharge;
- Capillary rise and linear relation runoff from groundwater.

S2.13.1 File names

Model: m_13_hillslope_7p_2s
Parameter ranges: m_13_hillslope_7p_2s_parameter_ranges

S2.13.2 Model equations

\[
\frac{dS_w}{dt} = P_e + C - (E_t + E_s) - Q_{se} \tag{112}
\]

\[
P_e = \max(P - D_h, 0) \tag{113}
\]

\[
C = c. \tag{114}
\]

\[
E_t + E_s = \begin{cases} E_p, & \text{if } S_w > 0 \\ 0, & \text{otherwise} \end{cases} \tag{115}
\]

\[
Q_{se} = \left(1 - \left(1 - \frac{S_h}{S_{h,max}}\right)^{\beta_h}\right) * P_e \tag{116}
\]

Figure S14: Structure of the Hillslope model

Where $S_w$ is the current soil water storage [mm]. Incoming precipitation $P$ [mm/d] is reduced by interception $D_h$ [mm/d], which is assumed to evaporate before the next precipitation event. $C$ is capillary rise from groundwater [mm/d], given as a constant rate. Evaporation from soil moisture $E_t + E_s$ [mm/d] occurs at the potential rate $E_p$ whenever possible. Storage excess surface runoff $Q_{se}$ [mm/d] depends on the fraction of the catchment that is currently saturated, expressed through parameters $S_{h,max}$ [mm] and $\beta_h$ [−].
\[
\frac{dS_{h,gw}}{dt} = Q_{se,g} - C - Q_{h,gw} \\
Q_{se,g} = (1 - a) * Q_{se} \\
Q_{h,gw} = K_h * S_{h,gw}
\]

(118) (119) (120)

Where \( S_{h,gw} \) is current groundwater storage [mm]. \( Q_{se,g} \) is the groundwater fraction of storage excess flow \( Q_{se} \) [mm/d], with \( Q_{se,s} \) as its complementary part. \( a \) is the parameter controlling this division [-]. Groundwater flow \( Q_{h,gw} \) [mm/d] depends linearly on current storage \( S_{h,gw} \) through parameter \( K_h \) [d\(^{-1}\)]. Total flow \( Q_t \) is the sum of \( Q_{h,gw} \) and \( Q_{h,srf} \), the latter of which is \( Q_{se,s} \) lagged over \( T_h \) days.
S2.14 TOPMODEL (model ID: 14)

The TOPMODEL (fig. S15) is originally a semi-distributed model that relies on topographic information (Beven and Kirkby, 1979). The model has undergone many revisions and significant differences can be seen between various publications. The version presented here is mostly based on Beven et al. (1995), with several necessary simplifications. Following Clark et al. (2008), the model is simplified to a lumped model (removing the distributed routing component) and all parameters are calibrated. This means the distribution of topographic index values that characterizes TOPMODEL are estimated using a shifted 2-parameter gamma distribution instead of being based on DEM data (Sivapalan et al., 1987; Clark et al., 2008). For simplicity of the evaporation calculations, the root zone store and unsaturated zone store are combined into a single threshold store with identical functionality to the original 2-store concept. The model has 2 stores and 7 parameters \( (S_{UZ,max}, S_t, K_d, q_0, f, \chi, \phi) \). The model aims to represent:

- Variable saturated area with direct runoff from the saturated part;
- Infiltration and saturation excess flow;
- Leakage to, and non-linear baseflow from, a deficit store.

S2.14.1 File names

Model: m_14_topmodel_7p_2s
Parameter ranges: m_14_topmodel_7p_2s_parameter_ranges

S2.14.2 Model equations

\[
\begin{align*}
\frac{dS_{UZ}}{dt} &= P_{eff} - Q_{ex} - E_a - Q_v \\
P_{eff} &= P - Q_{of} = P - A_C \cdot P \\
Q_{ex} &= \begin{cases} P_{eff}, & \text{if } S_{UZ} = S_{uz,max} \\ 0, & \text{otherwise} \end{cases} \\
E_a &= \begin{cases} E_p, & \text{if } S_{UZ} > S_t \cdot S_{UZ,max} \\ \frac{S_{UZ}}{S_t \cdot S_{UZ,max}} \cdot E_p, & \text{otherwise} \end{cases} \\
Q_v &= \begin{cases} k_d \frac{S_{UZ} - S_t \cdot S_{UZ,max}}{S_{UZ,max} (1 - S_t)}, & \text{if } S_{UZ} > S_t \cdot S_{UZ,max} \\ 0, & \text{otherwise} \end{cases}
\end{align*}
\]

Where \( S_{UZ} \) [mm] is the current storage in the combined unsaturated zone and root zone, with \( S_t \) [-] (fraction of \( S_{UZ,max} \)) indicating the boundary between the two...
and being the threshold above which drainage to the saturated zone can occur. $P_{eff}$ [mm/d] is the fraction of precipitation that does not fall on the saturated area $A_c \text{ [] }$, $E_o$ [mm/d] is evaporation that occurs at the potential rate for the unsaturated zone and scaled linearly with storage in the root zone, $Q_{ex}$ [mm/d] is overflow when the bucket reaches maximum capacity $S_{UZ,max}$ [mm], and $Q_v$ [mm/d] is drainage to the saturated zone, depending on time parameter $k_d$ [d$^{-1}$] and the relative storage in the unsaturated zone compared to the current deficit in the saturated zone.

$$\frac{dS_{SZ}}{dt} = -Q_v + Q_b$$ \hspace{1cm} (126)

$$Q_b = q_0 e^{-f*S_{SZ}}$$ \hspace{1cm} (127)

Where $S_{SZ}$ [mm] is the current storage deficit in the saturated zone store, which is increased by baseflow $Q_b$ [mm/d] and decreased by drainage $Q_v$. $Q_b$ relies on saturated flow rate $q_0$ [mm/d], parameter $f$ [mm$^{-1}$] and current deficit $S_{SZ}$. Total flow:

$$Q = Q_{of} + Q_{ex} + Q_b$$ \hspace{1cm} (128)

$$Q_{of} = A_c \times P$$ \hspace{1cm} (129)

The saturated area $A_c$ is calculated as follows. First, the within-catchment distribution of topographic index values is estimated with a shifted 2-parameter gamma distribution (Sivapalan et al., 1987; Clark et al., 2008):

$$f(\zeta) = \begin{cases} \frac{1}{\Gamma(\phi)} \left( \frac{\zeta-\mu}{\chi} \right)^{\phi-1} e^{\left( -\frac{\zeta-\mu}{\chi} \right)}, & \text{if } \zeta > \mu \\ 0, & \text{otherwise} \end{cases}$$ \hspace{1cm} (130)

Where $\Gamma$ is the gamma function and $\chi$, $\phi$ and $\mu$ are parameters of the gamma distribution. Following Clark et al. (2008), $\mu$ is fixed at $\mu = 3$ and $\chi$ and $\phi$ are calibration parameters. $\zeta$ represents the topographic index $\ln(a/tan\beta)$ with mean value $\lambda = \chi \phi + \mu$. Saturated area $A_c$ is computed as the fraction of the catchment that is above a deficit-dependent critical value $\zeta_{crit}$:

$$A_c = \int_{\zeta_{crit}}^{\infty} f(\zeta) d\zeta$$ \hspace{1cm} (131)

$$\zeta_{crit} = f * S_{SZ} + \lambda$$ \hspace{1cm} (132)
S2.15 Plateau model (model ID: 15)

The Plateau model (fig. S16) is a conceptualization of the perceived dominant processes in a typical Western European plateau (Savenije, 2010). It belongs to a 3-part topography driven modelling exercise, together with a wetland and hillslope conceptualization. Each model is provided in isolation here, because they are well-suited for isolating specific model structure choices. It has 2 stores and 8 parameters ($F_{\text{max}}, D_p, S_{u,max}, l_p, p, T_p, C$ and $K_p$). The model aims to represent:

- Stylized interception by vegetation;
- Evaporation controlled by a wilting point and moisture constrained transpiration;
- Separation between infiltration and infiltration excess flow;
- Capillary rise and linear relation runoff from groundwater.

S2.15.1 File names

Model: m_15_plateau_8p_2s
Parameter ranges: m_15_plateau_8p_2s_parameter_ranges

S2.15.2 Model equations

\[ \frac{dS_u}{dt} = P_i + C - E_t - R \]  
(133)

\[ P_i = \min(P_e, F_{\text{max}}) \]  
(134)

\[ = \min(\max(P - D_p, 0), F_{\text{max}}) \]  
(135)

\[ C = c. \]  
(136)

\[ E_t = E_p \times \max\left(\frac{P - s_u - s_{wp}}{s_{u,max} - s_{wp}}, 0\right) \]  
(137)

\[ R = \begin{cases} P_e + C, & \text{if } s_u = s_{u,max} \\ 0, & \text{otherwise} \end{cases} \]  
(138)

Where $S_u$ is the current soil water storage [mm]. Incoming precipitation $P$ [mm/d] is reduced by interception $D_p$ [mm/d], which is assumed to evaporate before the next precipitation event. $P_e$ is further divided into infiltration $P_i$ [mm/d] based on the maximum infiltration rate $F_{\text{max}}$ [mm/d] and infiltration excess $P_{ie} = P_e - P_i$ [mm/d]. $C$ is capillary rise from groundwater [mm/d], given as a constant rate.

Figure S16: Structure of the Plateau model
Evaporation from soil moisture $E_t$ [mm/d] occurs at the potential rate $E_p$ when $S_u$ is above the wilting point $S_{wp}$ [mm] (here defined as $S_{wp} = l_p * S_{u,max}$) and is further constrained by coefficient $p$ [-], which is between 0 and 1. Storage excess $R$ [mm/d] flows into the groundwater.

$$\frac{dS_{p, gw}}{dt} = R - C - Q_{p, gw}$$ \hspace{1cm} (139)

$$Q_{p, gw} = K_p * S_{p, gw}$$ \hspace{1cm} (140)

Where $S_{p, gw}$ is current groundwater storage [mm]. Groundwater flow $Q_{p, gw}$ [mm/d] depends linearly on current storage $S_{p, gw}$ through parameter $K_p$ [d$^{-1}$]. Total flow $Q_t$ is the sum of $Q_{p, gw}$ and $Q_{p, ieo}$, the latter of which is $P_{ie}$ lagged over $T_p$ days.
S2.16 New Zealand model v2 (model ID: 16)

The New Zealand model v2 (fig. S17) is part of a top-down modelling exercise that focusses on several catchments in New Zealand (Atkinson et al., 2002). It has 2 stores and 8 parameters ($I_{\text{max}}, S_{\text{max}}, S_{\text{fc}}, M, a, b$ and $t_{c,bf}, d$). The model aims to represent:

- Interception by vegetation;
- Separate vegetation and bare soil evaporation;
- Saturation excess overland flow;
- Subsurface runoff when soil moisture exceeds field capacity;
- Baseflow;
- Flow routing.

S2.16.1 File names

Model: m_16_newzealand2_8p_2s
Parameter ranges: m_16_newzealand2_8p_2s_parameter_ranges

S2.16.2 Model equations

\[
\frac{dS_i}{dt} = P - E_{\text{int}} - Q_{tf} \quad (141)
\]
\[
E_{\text{int}} = E_p \quad (142)
\]
\[
Q_{tf} = \begin{cases} P, & \text{if } S_i \geq I_{\text{max}} \\ 0, & \text{otherwise} \end{cases} \quad (143)
\]

Figure S17: Structure of the New Zealand model v1

Where $S_i$ [mm] is the current interception storage which gets replenished through daily precipitation $P$ [mm/d]. Intercepted water is assumed to evaporate ($E_{\text{int}}$ [mm/d]) at the potential rate $E_p$ [mm/d] when possible. $Q_{tf}$ [mm/d] represents throughfall towards soil moisture.
\[
\frac{dS_m}{dt} = Q_{tf} - E_{veg} - E_{bs} - Q_{se} - Q_{ss} - Q_{bf}
\] (144)

\[
E_{veg} = \begin{cases} 
M * E_p, & \text{if } S > S_{fc} \\
\frac{S}{S_{fc}} * M * E_p, & \text{otherwise}
\end{cases}
\] (145)

\[
E_{bs} = \frac{S}{S_{max}} (1 - M) * E_p
\] (146)

\[
Q_{se} = \begin{cases} 
P, & \text{if } S \geq S_{max} \\
0, & \text{otherwise}
\end{cases}
\] (147)

\[
Q_{ss} = \begin{cases} 
(a * (S - S_{fc}))^b, & \text{if } S \geq S_{fc} \\
0, & \text{otherwise}
\end{cases}
\] (148)

\[
Q_{bf} = t_{c,bf} * S
\] (149)

Where \(S_m \text{[mm]}\) is the current soil moisture storage which gets replenished through daily precipitation \(P \text{[mm/d]}\). Evaporation through vegetation \(E_{veg} \text{[mm/d]}\) depends on the forest fraction \(M \text{[-]}\) and field capacity \(S_{fc} \text{[mm]}\). \(E_{bs} \text{[mm/d]}\) represents bare soil evaporation. When \(S\) exceeds the maximum storage \(S_{max} \text{[mm]}\), water leaves the model as saturation excess runoff \(Q_{se}\). If \(S\) exceeds field capacity \(S_{fc} \text{[mm]}\), subsurface runoff \(Q_{ss} \text{[mm/d]}\) is generated controlled by time parameter \(a \text{[d}^{-1}\)] and nonlinearity parameter \(b \text{[-]}\). \(Q_{bf}\) represents baseflow controlled by time scale parameter \(t_{c,bf} \text{[d}^{-1}\)].

Total runoff \(Q_t \text{[mm/d]}\) is:

\[
Q_t = Q_{se} + Q_{ss} + Q_{bf}
\] (150)

Total flow is delayed by a triangular routing scheme controlled by time parameter \(d \text{[d]}\).
S2.17 Penman model (model ID: 17)

The Penman model (fig. S18) is based on the drying curve concept described in Penman (1950) (Wagener et al., 2002). It has 3 stores and 4 parameters ($S_{\text{max}}$, $\phi$, $\alpha$, $k_1$). The model aims to represent:

- Moisture accumulation and evaporation from the root zone;
- Bypass of excess moisture to the stream;
- Deficit-based groundwater accounting;
- Linear flow routing.

S2.17.1 File names

Model: m_17_penman_4p_3s
Parameter ranges: m_17_penman_4p_3s_parameter_ranges

S2.17.2 Model equations

\[
\frac{dS_{rz}}{dt} = P - E_a - Q_{ex} \quad (151)
\]

\[
E_a = \begin{cases} 
E_p, & \text{if } S_{rz} > 0 \\
0, & \text{otherwise}
\end{cases} \quad (152)
\]

\[
P_{ex} = \begin{cases} 
P, & \text{if } S_{rz} = S_{\text{max}} \\
0, & \text{otherwise}
\end{cases} \quad (153)
\]

Where $S_{rz}$ [mm] is the current storage in the root zone, refilled by precipitation $P$ [mm/d] and drained by evaporation $E_a$ [mm/d] and moisture excess $q_{ex}$ [mm/d]. $E_a$ occurs at the potential rate $E_p$ [mm/d] whenever possible. $q_{ex}$ occurs only when the store is at maximum capacity $S_{\text{max}}$ [mm].

\[
\frac{dS_{\text{def}}}{dt} = E_t + u_2 - q_{12} \quad (154)
\]

\[
E_t = \begin{cases} 
\gamma \times E_p, & \text{if } S_{rz} = 0 \\
0, & \text{otherwise}
\end{cases} \quad (155)
\]

\[
u_2 = \begin{cases} 
q_{12}, & \text{if } S_{\text{def}} = 0 \\
0, & \text{otherwise}
\end{cases} \quad (156)
\]

\[
q_{12} = (1 - \phi) \times q_{ex} \quad (157)
\]

Where $S_{\text{def}}$ [mm] is the current moisture deficit, which is increased by evaporation $E_t$ [mm/d] and reduced by inflow $q_{12}$ [mm/d]. $E_t$ occurs only when the upper store...
$S_{rez}$ is empty and at a fraction $\gamma$ [-] of $E_p$. Inflow $q_{12}$ is the fraction $(1-\phi)$ [-] of $q_{ex}$ that does not bypass the lower soil layer. Saturation excess $u_2$ [mm/d] occurs only when there is zero deficit.

\[
\frac{dC_{res}}{dt} = u_1 + u_2 - Q \tag{158}
\]

\[
Q = k_1 \ast C_{res} \tag{159}
\]

Where $C_{res}$ [mm] is the current storage in the routing reservoir, increased by $u_1$ and $u_2$, and drained by runoff $Q$ [mm/d]. $Q$ has a linear relationship with storage through time scale parameter $k_1$ [$d^{-1}$].
S2.18 SIMHYD (model ID: 18)

The SIMHYD model (fig. S19) is a simplified version of MODHYDROLOG, originally developed for use in Australia (Chiew et al., 2002). It has 3 stores (I, SMS, GW) and 7 parameters (INSC, COEFF, SQ, SMSC, SUB, CRAK, K). The model aims to represent:

- Interception by vegetation;
- Infiltration and infiltration excess flow;
- Preferential groundwater recharge, interflow and saturation excess flow;
- Groundwater recharge resulting from filling up of soil moisture storage capacity;
- Slow flow from groundwater.

S2.18.1 File names

Model: m_18_simhyd_7p_3s
Parameter ranges: m_18_simhyd_7p_3s_parameter_ranges

S2.18.2 Model equations

\[
\frac{dI}{dt} = P - E_i - EXC
\]

\[
E_i = \begin{cases} 
E_p, & \text{if } I > 0 \\
0, & \text{otherwise} 
\end{cases}
\]

\[
EXC = \begin{cases} 
P, & \text{if } I = INSC \\
0, & \text{otherwise} 
\end{cases}
\]

Where \( I \) is the current interception storage [mm], \( P \) precipitation [mm/d], \( E_i \) the evaporation from the interception store [mm/d] and \( EXC \) the excess rainfall [mm/d]. Evaporation is assumed to occur at the potential rate when possible. When \( I \) exceeds the maximum interception capacity \( INSC \) [mm], water is routed to the rest of the model as excess precipitation \( EXC \).
\[
\begin{align*}
\frac{dSMS}{dt} &= SMF - ET - GWF & (163) \\
SMF &= INF - INT - REC & (164) \\
INF &= \min\left(\text{COEFF} \times \exp\left(-SQ \times \frac{SMS}{SMSC}\right), \text{EXC}\right) & (165) \\
INT &= \text{SUB} \times \frac{SMS}{SMSC} \times \text{INF} & (166) \\
REC &= \text{CRAK} \times \frac{SMS}{SMSC} \times (INF - INT) & (167) \\
ET &= \min\left(10 \times \frac{SMS}{SMSC}, \text{PET}\right) & (168) \\
GWF &= \begin{cases} 
SMF, & \text{if } SMS = SMSC \\
0, & \text{otherwise} 
\end{cases} & (169)
\end{align*}
\]

Where SMS is the current storage in the soil moisture store [mm]. INF is total infiltration [mm/d] from excess precipitation, based on maximum infiltration loss parameter COEFF [-], the infiltration loss exponent SQ [-] and the ratio between current soil moisture storage SMS and the maximum soil moisture capacity SMSC [mm]. INT represents interflow and saturation excess flow [mm/d], using a constant of proportionality SUB [-]. REC is preferential recharge of groundwater [mm/d] based on another constant of proportionality CRAK [-]. SMF is flow into soil moisture storage [mm/d]. ET evaporation from the soil moisture that occurs at the potential rate when possible [mm/d], and GWF the flow to the groundwater store [mm/d]:

\[
\begin{align*}
\frac{dGW}{dt} &= REC + GWF - BAS & (170) \\
BAS &= K \times GW & (171)
\end{align*}
\]

Where GW is the current storage [mm] in the groundwater reservoir. Outflow BAS [mm/d] from the reservoir has a linear relation with storage through the linear recession parameter K [d^-1]. Total outflow \( Q_t \) [mm/d] is the sum of three parts:

\[
\begin{align*}
Q_t &= SRUN + INT + BAS & (172) \\
SRUN &= EXC - INF & (173)
\end{align*}
\]
S2.19  Australia model (model ID: 19)

The Australia model (fig. S20) is part of a top-down modelling exercise designed to use auxiliary data (Farmer et al., 2003). Some adjustments were made to the evaporation equations: these were originally separated between vegetation and bare soil evaporation, scaled between the unsaturated and saturated zone. This has been simplified to separation between unsaturated and saturated evaporation only. The model has 3 stores and 8 parameters \( S_b, \phi, f_c, \alpha_{SS}, \beta_{SS}, K_{deep}, \alpha_{BF}, \beta_{BF} \). For consistency with other model formulations, \( S_b \) is is used as a parameter, instead of being broken down into its constitutive parts \( D \) and \( \phi \). The model aims to represent:

- Separation of saturated zone and a variable-size unsaturated zone;
- Evaporation from unsaturated and saturated zones;
- Saturation excess and non-linear subsurface flow;
- Deep groundwater recharge and baseflow.

S2.19.1  File names

Model:  m_19_australia_8p_3s
Parameter ranges: m_19_australia_8p_3s_parameter_ranges

S2.19.2  Model equations

\[
\frac{dS_{us}}{dt} = P - E_{us} - r_g - s_e \tag{174}
\]

\[
E_{us} = \frac{S_{us}}{S_b} \cdot E_p \tag{175}
\]

\[
S_b = D \cdot \phi \tag{176}
\]

\[
r_g = \begin{cases} P, & \text{if } S_{us} > S_{usfc} \\ 0, & \text{otherwise} \end{cases} \tag{177}
\]

\[
s_e = \begin{cases} S_{us} - S_{usfc}, & \text{if } S_{us} > S_{usfc} \\ 0, & \text{otherwise} \end{cases} \tag{178}
\]

\[
S_{usfc} = (S_b - S_{sat}) \cdot \frac{f_c}{\phi} \tag{179}
\]

Where \( S_{us} \) is the current storage in the unsaturated store [mm], \( P \) the current precipitation [mm/d], \( S_b \) [mm] the maximum storage of the soil profile, based on the soil depth \( D \) [mm] and the porosity \( \phi \) [-]. \( r_g \) [mm/d] is drainage from the unsaturated store to the saturated store, based on the variable field capacity \( S_{usfc} \) [mm]. \( S_{usfc} \) is...
based on the current storage on the saturated zone $S_{sat}$ [mm], the maximum soil moisture storage $S_b$ [mm], the field capacity $f_c$ [-] and the porosity $\phi$ [-]. $s_e$ [mm/d] is the storage excess, resulting from a decrease of $S_{usfc}$ that leads to more water being stored in the unsaturated zone than should be possible.

$$\frac{dS_{sat}}{dt} = r_g - E_{sat} - Q_{SE} - Q_{SS} - Q_R \quad (180)$$

$$E_{sat} = \frac{S_{sat}}{S_b} \cdot E_p \quad (181)$$

$$Q_{SE} = \begin{cases} r_g + S_e, & \text{if } S_{sat} > S_b \\ 0, & \text{otherwise} \end{cases} \quad (182)$$

$$Q_{SS} = \alpha_{SS} \cdot (S_{sat})^{\beta_{SS}} \quad (183)$$

$$Q_R = K_{deep} \cdot S_{sat} \quad (184)$$

Where $S_{sat}$ is the current storage in the saturated zone [mm], $E_{sat}$ is the evaporation from the saturated zone [mm], $Q_{SE}$ saturation excess runoff [mm/d] that occurs when the saturated zone reaches maximum capacity $S_b$ [mm], $Q_{SS}$ is subsurface flow [mm/d] and $Q_R$ is recharge of deep groundwater [mm/d]. Both $Q_{SS}$ and $Q_R$ are based on the dimensionless fraction $r$ and subsurface flow constants $c$ [d$^{-1}$] and $d$ [-].

$$\frac{dG_w}{dt} = Q_R - Q_{BF} \quad (185)$$

$$Q_{BF} = \alpha_{BF} \cdot (G_w)^{\beta_{BF}} \quad (186)$$

Where $G_w$ is the current groundwater storage [mm] and $Q_{BF}$ baseflow, dependent on parameters $\alpha_{BF}$ [d$^{-1}$] and $\beta_{BF}$ [-]. Total runoff is the sum of $Q_{SE}$, $Q_{SS}$ and $Q_{BF}$:

$$Q = Q_{SE} + Q_{SS} + Q_{BF} \quad (188)$$
S2.20 Generalized Surface inFiltration Baseflow model (model ID: 20)

The GSFB model (fig. S21) is originally developed for use in Australian ephemeral catchments (Nathan and McMahon, 1990; Ye et al., 1997). It has 3 stores and 8 parameters ($C$, $NDC$, $S_{max}$, $E_{max}$, $F_{rate}$, $B$, $DPF$, $SDR_{max}$). The model aims to represent:

- Saturation excess surface runoff;
- Threshold-based infiltration;
- Threshold-based baseflow;
- Deep percolation and water rise to meet evaporation demand.

S2.20.1 File names

Model: m_20_gsfb_8p_3s
Parameter ranges: m_20_gsfb_8p_3s_parameter_ranges

S2.20.2 Model equations

\[
\frac{dS}{dt} = P + Q_{dr} - E_a - Q_s - F \tag{189}
\]

\[
Q_{dr} = \begin{cases} 
    C \cdot DS \cdot \left(1 - \frac{S}{NDC \cdot S_{max}}\right), & \text{if } S \leq NDC \cdot S_{max} \\
    0, & \text{otherwise}
\end{cases} \tag{190}
\]

\[
E_a = \begin{cases} 
    E_p, & \text{if } S > NDC \cdot S_{max} \\
    \min \left( E_p, E_{max} \frac{S}{NDC \cdot S_{max}} \right), & \text{otherwise}
\end{cases} \tag{191}
\]

\[
Q_s = \begin{cases} 
    P, & \text{if } S = S_{max} \\
    0, & \text{otherwise}
\end{cases} \tag{192}
\]

\[
F = \begin{cases} 
    F_{rate}, & \text{if } S > NDC \cdot S_{max} \\
    0, & \text{otherwise}
\end{cases} \tag{193}
\]

Where $S$ [mm] is the current storage in the upper zone, refilled by precipitation $P$ [mm/d] and recharge from deep groundwater $Q_{dr}$ [mm/d]. The store is drained by evaporation $E_a$ [mm/d], surface runoff $Q_s$ [mm/d] and infiltration $F$ [mm/d]. $E_a$ occurs at the potential rate $E_p$ [mm/d] if the store is above a threshold capacity given as the fraction $NDC$ [-] of maximum storage $S_{max}$ [mm]. Evaporation occurs at a
reduced rate scaled by maximum evaporation rate $E_{max}$ [mm/d] if the store is below this threshold. $Q_s$ occurs only if the store is at maximum capacity $S_{max}$. $F$ occurs at a constant rate $F_{rate}$ if the store is above threshold $NDC * S_{max}$. Recharge from deep percolation only occurs if the store is below threshold capacity $NDC * S_{max}$ and uses time parameter $C$ [d$^{-1}$] and current deep storage $DS$ [mm].

\[
\frac{dSS}{dt} = F - Q_b - D_p
\]

\[
Q_b = \begin{cases} 
B \times DPF \times (SS - SDR_{max}), & \text{if } SS > SDR_{max} \\
0, & \text{otherwise}
\end{cases}
\]

\[
D_p = (1 - B) \times DPF \times SS
\]

Where $SS$ [mm] is the current storage in the subsurface store, refilled by infiltration $F$ and drained by baseflow $Q_b$ [mm/d] and deep percolation $D_p$ [mm/d]. Outflow from this store is given as a function of storage $DS$ and time coefficient $DPF$ [d$^{-1}$]. A fraction $1 - B \times -$ of this outflow is deep percolation $D_p$. The remaining fraction $B \times -$ is baseflow $Q_b$, provided the store is above threshold $SDR_{max}$ [mm].

\[
\frac{dDS}{dt} = D_p - Q_{dr}
\]

Where $DS$ [mm] is the current storage in the deep store, refilled by a deep percolation $D_p$ and drained by recharge to the upper store $Q_{dr}$. Total flow:

\[
Q_t = Q_s + Q_b
\]
S2.21 Flex-B (model ID: 21)

The Flex-B model (fig. S22) is the basis of a model development study (Fenicia et al., 2008). It has 3 stores and 9 parameters \( UR_{\text{max}}, \beta, D, \text{Perc}_{\text{max}}, L_p, N_{\text{lag,f}}, N_{\text{lag,s}}, K_f, K_s \). The model aims to represent:

- Infiltration and saturation excess flow based on a distribution of different soil depths;
- A split between fast saturation excess flow and preferential recharge to a slow store;
- Percolation from the unsaturated zone to a slow runoff store.

S2.21.1 File names

Model: m_21_flexb_9p_3s
Parameter ranges: m_21_flexb_9p_3s_parameter_ranges

S2.21.2 Model equations

\[
\frac{dUR}{dt} = R_u - E_{ur} - R_p \quad (200)
\]

\[
R_U = (1 - C_r) \ast P_{eff} \quad (201)
\]

\[
C_r = \left[ 1 + \exp \left( \frac{-UR}{UR_{\text{max}}} + \frac{1}{2} \right) \beta \right]^{-1} \quad (202)
\]

\[
E_{ur} = E_p \ast \min \left( 1, \frac{UR}{UR_{\text{max}}} \right) \quad (203)
\]

\[
P_s = \text{Perc}_{\text{max}} \ast \frac{UR}{UR_{\text{max}}} \quad (204)
\]

Where \( UR \) is the current storage in the unsaturated zone \([\text{mm}]\). \( R_u \) \([\text{mm/d}]\) is the inflow into UR based on its current storage compared to maximum storage \( UR_{\text{max}} \) \([\text{mm}]\) and a shape distribution parameter \( \beta \) \([-]\). \( E_{ur} \) the evaporation \([\text{mm/d}]\) from UR which follows a linear relation between current and maximum storage until a threshold \( L_p \) \([-]\) is exceeded. \( P_s \) is the percolation from UR to the slow reservoir SR \([\text{mm/d}]\), based on a maximum percolation rate \( \text{Perc}_{\text{max}} \) \([\text{mm}]\), relative to the fraction of current storage and maximum storage. \( P_{eff} \) is routed towards the unsaturated zone based on \( C_r \), with the remainder being divided into preferential recharge \( R_s \) \([\text{mm/d}]\) and fast runoff \( R_f \) \([\text{mm/d}]\):
\begin{align*}
R_s &= (P_{s,ff} - R_u) \times D \\
R_f &= (P_{s,ff} - R_u) \times (1 - D)
\end{align*}

Where \( R_s \) and \( R_f \) are the flows [\text{mm/d}] to the slow and fast runoff reservoir respectively, based on runoff partitioning coefficient \( D [\cdot] \). Both are lagged by linearly increasing triangular transformation functions with parameters \( N_{\text{lag},s} [\text{d}] \) and \( N_{\text{lag},f} [\text{d}] \) respectively. Percolation \( R_p \) is added to \( R_s \) before the transformation to \( R_{s,l} \) occurs.

\begin{align*}
\frac{dFR}{dt} &= R_{f,l} - Q_f \\
Q_f &= K_f \ast FR
\end{align*}

Where \( FR \) is the current storage [\text{mm}] in the fast flow reservoir. Outflow \( Q_f [\text{mm/d}] \) from the reservoir has a linear relation with storage through time scale parameter \( K_f [\text{d}^{-1}] \).

\begin{align*}
\frac{dSR}{dt} &= R_{s,l} - Q_s \\
Q_s &= K_s \ast SR
\end{align*}

Where \( SR \) is the current storage [\text{mm}] in the slow flow reservoir. Outflow \( Q_s [\text{mm/d}] \) from the reservoir has a linear relation with storage through time scale parameter \( K_s [\text{d}^{-1}] \). Total outflow \( Q [\text{mm/d}] \):

\begin{equation}
Q = Q_f + Q_s
\end{equation}
S2.22 Variable Infiltration Capacity (VIC) model (model ID: 22)

The VIC model (fig. S23) is originally developed for use with General Circulation Models and uses latent and sensible heat fluxes to determine the rainfall-runoff relationship (Liang et al., 1994). For consistency with other models in this framework, we use a conceptualized version based in part of the VIC implementation in Clark et al. (2008). In addition, the original Leaf-Area-Index-based interception capacity is replaced with a sinusoidal curve-based approximation of interception capacity. The model has 3 stores and 10 parameters ($\bar{I}$, $I_s$, $S_{sm,max}$, $b$, $c_1$, $S_{gw,max}$, $k_2$, $c_2$). The model aims to represent:

- Time-varying interception by vegetation;
- Variable infiltration and saturation excess flow;
- Interflow and baseflow from a deeper groundwater layer.

S2.22.1 File names
Model: m_22_vic_10p_3s
Parameter ranges: m_22_vic_10p_3s_parameter_ranges

S2.22.2 Model equations

$\frac{dS_i}{dt} = P - E_i - P_{eff} - I_{ex}$  \hspace{1cm} (212)

$E_i = \frac{S_i}{I_{max}} \cdot E_p$  \hspace{1cm} (213)

$I_{max} = \bar{I} (1 + I_s \times \sin (2\pi (t + I_s)))$  \hspace{1cm} (214)

$P_{eff} = \begin{cases} P, & \text{if } S_i = I_{max} \\ 0, & \text{otherwise} \end{cases}$  \hspace{1cm} (215)

$I_{ex} = \text{max}(S_i - I_{max})$  \hspace{1cm} (216)

Where $S_i$ [mm] is the current interception storage, refilled by precipitation $P$ [mm/d] and drained by evaporation $E_i$ [mm/d] and interception excess flows $P_{eff}$ [mm/d] and $I_{ex}$ [mm/d]. $E_i$ decreases linearly with storage, based on maximum storage $I_{max}$ [mm]. $I_{max}$ is determined using the mean interception $\bar{I}$ [mm], fractional seasonal interception change $I_s$ [-] and time shift $I_s$ [-]. It is implicitly assumed that 1 sinusoidal period corresponds with a growing season of 1 year. $P_{eff}$ is effective rainfall when the store is at maximum capacity. $I_{ex}$ is an auxiliary flux used when a change in storage size result in current storage $S_i$ exceeding $I_{max}$.
\[
dS_{sm} \frac{dt}{dt} = inf - Et_1 - Q_{ex1} - pc
\]  \hspace{1cm} (217)

\[
inf = (P_{eff} + I_{ex}) - Q_{ie}
\]  \hspace{1cm} (218)

\[
Q_{ie} = (P_{eff} + I_{ex}) \ast \left(1 - \left(1 - \frac{S_{sm}}{S_{sm,max}}\right)^b\right)
\]  \hspace{1cm} (219)

\[
E_{t1} = \frac{S_{sm}}{S_{sm,max}} \ast (E_p - E_i)
\]  \hspace{1cm} (220)

\[
Q_{ex1} = \begin{cases} inf, & \text{if } S_{sm} = S_{sm,max} \\ 0, & \text{otherwise} \end{cases}
\]  \hspace{1cm} (221)

\[
pc = k_1 \ast \left(\frac{S_{sm}}{S_{sm,max}}\right)^{c_1}
\]  \hspace{1cm} (222)

Where \(S_{sm} [\text{mm}]\) is the current soil moisture storage, refilled by infiltration \(inf [\text{mm/d}]\), and drained by evapotranspiration \(Et_1 [\text{mm/d}]\), storage excess \(Q_{ex1} [\text{mm/d}]\) and percolation \(pc [\text{mm/d}]\). \(inf\) relies on the value of infiltration excess \(Q_{ie}\), which is calculated using the maximum soil moisture storage \(S_{sm,max} [\text{mm}]\) and shape parameter \(b [-]\). \(Et_1\) scales linearly with current storage. \(Q_{ex1}\) equals \(inf\) when the store is at maximum capacity. \(pc\) has a potentially non-linear relationship with current storage through time parameter \(k_1 [\text{d}^{-1}]\) and shape parameter \(c_1\).

\[
dS_{gw} \frac{dt}{dt} = pc - Et_2 - Q_{ex2} - Q_b
\]  \hspace{1cm} (223)

\[
Et_2 = \frac{S_{gw}}{S_{gw,max}} \ast (E_p - E_i - Et_1)
\]  \hspace{1cm} (224)

\[
Q_{ex2} = \begin{cases} pc, & \text{if } S_{gw} = S_{gw,max} \\ 0, & \text{otherwise} \end{cases}
\]  \hspace{1cm} (225)

\[
Q_b = k_2 \ast \left(\frac{S_{gw}}{S_{gw,max}}\right)^{c_2}
\]  \hspace{1cm} (226)

Where \(S_{gw} [\text{mm}]\) is the current groundwater storage, refilled through percolation \(pc [\text{mm/d}]\) and drained by evapotranspiration \(Et_2 [\text{mm/d}]\), excess flow \(Q_{ex2} [\text{mm/d}]\) and baseflow \(Q_b [\text{mm/d}]\). \(Et_2\) is scaled linearly with current storage based on maximum storage \(S_{gw,max} [\text{mm}]\). \(Q_{ex2}\) equals \(pc\) when the store is at maximum capacity. \(Q_b\) has a potentially non-linear relationship with current storage through time parameter \(k_2\) and shape parameter \(c_2\). Total outflow:

\[
Q_t = Q_{ie} + Q_{ex1} + Q_{ex2} + Q_b
\]  \hspace{1cm} (227)
S2.23 Large-scale catchment water and salt balance model element (model ID: 23)

The large-scale catchment water and salt balance model (LASCAM) (fig. S24) is part of a study that investigates soil water and salt concentration before and after forest clearing (Sivapalan et al., 1996). It is a semi-distributed model made up of individual elements, such as described below. The model presented here simulates the water balance only (salt is ignored). It has 3 stores and 24 parameters (αf, βf, Bmax, Fmax, αc, βc, Amin, Amax, αss, βss, c, αg, βg, γf, δf, t4, αb, βb, γa, δa, αa, βa, γb, δb). The model aims to represent:

- Stylized interception;
- Saturation and infiltration excess surface runoff;
- An inner layout representing near-stream saturated storage, deep saturated storage and medium-depth unsaturated storage;
- Subsurface saturation and infiltration excess flow to the near-stream store;
- Percolation to and capillary rise from groundwater.

S2.23.1 File names

Model: m_23_lascam_24p_3s
Parameter ranges: m_23_lascam_24p_3s_parameter_ranges

S2.23.2 Model equations
Where $F$ [mm] is the current storage in the unsaturated infiltration store, which
controls the amount of subsurface runoff generated on the boundary of a more
permeable top layer (store $A$) with a less permeable bottom layer (store $F$). $F$ is refilled
by actual infiltration $f_a$ [mm/d], and drained by recharge $r_f$ [mm/d] and evaporation
$E_b$ [mm/d]. $f_a$ depends on the actual infiltration rate $P_c$ [mm/d], the fraction
saturated catchment area $\phi_{ss}$ [-], the fraction variable area contributing to overland
flow $\phi_c$ [-] and a catchment-scale infiltration capacity $f_{ss}^*$ [mm/d]. $f_{ss}^*$ depends on a
scaling parameter $\alpha_f$ [mm/d], the relative storage in groundwater $B/B_{max}$, the relative
infiltration volume in the catchment $F/F_{max}$ and non-linearity parameter $\beta_f$ [-].
$B_{max}$ [mm] and $F_{max}$ [mm] are storage scaling parameters [-]. $\phi_c$ uses the minimum
contributing area $A_{min}$ [mm], maximum contributing area $A_{max}$ [mm] and shape
parameters $\alpha_c$ [-] and $\beta_c$ [-] to control the shape of this distribution. $\phi_{ss}$ takes a similar
shape as $\phi_c$, using parameters $\alpha_{ss}$ [-] and $\beta_{ss}$ [-]. $P_c$ is the lesser of throughfall rate
$P_g$ [mm/d] minus saturation excess $q_{se}$ [mm/d], and the catchment infiltration capacity
$f_{ss}^*$ [mm/d]. $f_{ss}^*$ is assumed to have a constant rate $c$ [mm/d]. $q_{se}$ is determined as that part of throughfall $P_g$ that falls on the variable contributing catchment area
given by $\phi_c$. $P_g$ is determined as a fixed interception rate $\alpha_g$ [mm/d] and a fractional
interception $\beta_g$ [-]. Evaporation $E_f$ uses the potential rate $E_p$ [mm/d] scaled by the
relative storage in $F$ and two shape parameters $\gamma_f$ [-] and $\delta_f$ [-]. Recharge $r_f$ [mm/d]
has a linear relation with storage through time parameter $t_d$ [d$^{-1}$].
\[
\frac{dA}{dt} = q_{sse} + q_{sie} + q_b - E_a - q_a - r_a \quad (239)
\]

\[
q_{sse} = \frac{\phi_{ss} - \phi_c}{1 - \phi_c} P_c \quad (240)
\]

\[
q_{sie} = \max \left( P_c \left( 1 - \frac{\phi_{ss}}{1 - \phi_c} \right), 0 \right) \quad (241)
\]

\[
q_b = \beta_b \left( \exp \left( \alpha_b \frac{B}{B_{\max}} \right) - 1 \right) \quad (242)
\]

\[
E_a = \phi_c * E_p + \gamma_a * E_p \left( \frac{A}{A_{\max}} \right)^{\delta_a} \quad (243)
\]

\[
q_a = \begin{cases} 
\alpha_a \left( \frac{A - A_{\min}}{A_{\max} - A_{\min}} \right)^{\beta_a}, & \text{if } A > A_{\min} \\
0, & \text{otherwise} 
\end{cases} \quad (244)
\]

\[
r_a = \phi_{ss} * f_{ss}^* \quad (245)
\]

Where \( A [\text{mm}] \) is the current storage in the more permeable upper zone (above less permeable lower zone \( F \)), refilled by sub-surface saturation excess \( q_{sse} [\text{mm/d}] \), sub-surface infiltration excess \( q_{sie} [\text{mm/d}] \) and discharge from groundwater \( q_b [\text{mm/d}] \). The store is drained by evaporation \( E_a \), sub-surface stormflow \( q_a [\text{mm/d}] \) and recharge \( r_a [\text{mm/d}] \). Flow from store \( B \), \( q_b \), decreases exponentially as the store dries out, controlled by parameters \( \beta_b \) and \( \alpha_b \). Evaporation \( E_a \) occurs at the potential rate \( E_p \) from the variable saturated area \( \phi_c \) and additionally at a rate scaled by the relative storage in \( A \) and two shape parameters \( \gamma_a [-] \) and \( \delta_a [-]/ \). Recharge \( r_a \) is a function of the saturated subsurface area \( \phi_{ss} \) and the sub-surface infiltration rate \( f_{ss}^* \).

\[
\frac{dB}{dt} = r_f + r_a - E_b - q_b \quad (246)
\]

\[
E_b = \gamma_b * E_p \left( \frac{B}{B_{\max}} \right)^{\delta_b} \quad (247)
\]

Where \( B [\text{mm}] \) is the current storage in the deep layers, refilled by recharge from stores \( A (r_a) \) and \( F (r_f) \), and drained by evaporation \( E_b \) and groundwater discharge \( q_b \). \( E_b \) uses the potential rate \( E_p \) scaled by the relative storage in \( B \) and two shape parameters \( \gamma_b [-] \) and \( \delta_b [-]/ \). Total flow:

\[
Q_t = q_{sc} + q_{ic} + q_a \quad (248)
\]

\[
q_{ic} = P_g - q_{sc} - P_c \quad (249)
\]

Where \( q_{ic} [\text{mm/d}] \) is infiltration excess on the surface.
S2.24 MOPEX-1 (model ID: 24)

The MOPEX-1 model (fig. S25) is part of a model improvement study that investigates the relationship between dominant processes and model structures for 197 catchments in the MOPEX database (Ye et al., 2012). It has 4 stores and 5 parameters ($S_{b1}$, $t_w$, $t_u$, $S_e$, $t_c$). The model aims to represent:

- Saturation excess flow;
- Infiltration to deeper soil layers;
- A split between fast and slow runoff.

S2.24.1 File names
Model: m_24_mopex1_5p_4s
Parameter ranges: m_24_mopex1_5p_4s_parameter_ranges

S2.24.2 Model equations

\[
\frac{dS_1}{dt} = P - ET_1 - Q_{1f} - Q_w
\]  
\[(250)\]

\[
ET_1 = \frac{S_1}{S_{b1}} \times Ep
\]  
\[(251)\]

\[
Q_{1f} = \begin{cases} 
P, & \text{if } S_1 \geq S_{b1} \\
0, & \text{otherwise}
\end{cases}
\]  
\[(252)\]

\[
Q_w = t_w \times S_1
\]  
\[(253)\]

Where $S_1$ [mm] is the current storage in soil moisture and $P$ precipitation [mm/d]. Evaporation $ET_1$ [mm/d] depends linearly on current soil moisture, maximum soil moisture $S_{b1}$ [mm] and potential evapotranspiration $Ep$ [mm/d]. Saturation excess flow $Q_{1f}$ [mm/d] occurs when the soil moisture bucket exceeds its maximum capacity. Infiltration to deeper groundwater $Q_w$ [mm/d] depends on current soil moisture and time parameter $t_w$ [d$^{-1}$].

\[
\frac{dS_2}{dt} = Q_w - ET_2 - Q_{2u}
\]  
\[(254)\]

\[
ET_2 = \frac{S_2}{S_e} \times Ep
\]  
\[(255)\]

\[
Q_{2u} = t_u \times S_2
\]  
\[(256)\]

Where $S_2$ [mm] is the current groundwater storage, refilled by infiltration from $S_1$. Evaporation $ET_2$ [mm/d] depends linearly on current groundwater and root zone
storage capacity $S_c$ [mm]. Leakage to the slow runoff store $Q_{2u}$ [mm/d] depends on current groundwater level and time parameter $t_u$ [d$^{-1}$].

$$\frac{dS_{c1}}{dt} = Q_{1f} - Q_f$$  \hspace{1cm} (257)
$$Q_f = t_c \times S_{c1}$$  \hspace{1cm} (258)

Where $S_{c1}$ [mm] is current storage in the fast flow routing reservoir, refilled by $Q_{1f}$. Routed flow $Q_f$ depends on the mean residence time parameter $t_c$ [d$^{-1}$].

$$\frac{dS_{c2}}{dt} = Q_{2u} - Q_u$$  \hspace{1cm} (259)
$$Q_u = t_c \times S_{c2}$$  \hspace{1cm} (260)

Where $S_{c2}$ [mm] is current storage in the slow flow routing reservoir, refilled by $Q_{2u}$. Routed flow $Q_u$ depends on the mean residence time parameter $t_c$ [d$^{-1}$]. Total simulated flow $Q_t$ [mm/d]:

$$Q_t = Q_f + Q_u$$  \hspace{1cm} (261)
S2.25  Thames Catchment Model (model ID: 25)

The Thames Catchment Model (TCM) model (fig. S26) is originally intended to be used in zones with similar surface characteristics, rather than catchments as a whole (Moore and Bell, 2001). It has 4 stores and 6 parameters (\(\phi, rc, \gamma, k_1, c_a, k_2\)). The model aims to represent:

- Effective rainfall before infiltration;
- Preferential recharge;
- Catchment drying through prolonged soil moisture depletion;
- Groundwater abstraction;
- Non-linear groundwater flow.

S2.25.1  File names

Model:  m_25_tcm_6p_4s
Parameter ranges:  m_25_tcm_6p_4s_parameter_ranges

S2.25.2  Model equations

\[
\frac{dS_{rz}}{dt} = P_{in} - E_a - q_{ex1} \tag{262}
\]

\[
P_{in} = (1 - \phi) \times P_n \tag{263}
\]

\[
P_n = \max(P - E_p, 0) \tag{264}
\]

\[
E_a = \begin{cases} 
E_p, & \text{if } S_{rz} > 0 \\
0, & \text{otherwise}
\end{cases} \tag{265}
\]

\[
q_{ex1} = \begin{cases} 
P_{in}, & \text{if } S_{rz} > rc \\
0, & \text{otherwise}
\end{cases} \tag{266}
\]

Where \(S_{rz} \,[\text{mm}]\) is the current storage in the root zone, refilled by infiltrated precipitation \(P_{in} \,[\text{mm} / \text{d}]\), and drained by evaporation \(E_a \,[\text{mm} / \text{d}]\) and storage excess flow \(q_{ex1} \,[\text{mm} / \text{d}]\). \(P_{in}\) is the fraction \((1 - \phi) \,[\text{]}\) of net precipitation \(P_n \,[\text{mm} / \text{d}]\) that is not preferential recharge. \(P_n\) is the difference between precipitation \(P \,[\text{mm} / \text{d}]\) and potential evapotranspiration \(E_p \,[\text{mm} / \text{d}]\) per time step. \(E_a\) occurs at the net potential rate whenever possible. \(q_{ex1}\) occurs only when the store is at maximum capacity \(rc \,[\text{mm}]\). 

Figure S26: Structure of the TCM model

53
\[
\frac{dS_{\text{def}}}{dt} = E_t + q_{ex2} - q_{ex1}
\] (267)

\[
E_t = \begin{cases} 
\gamma \ast E_p, & \text{if } S_{rz} = 0 \\
0, & \text{otherwise}
\end{cases}
\] (268)

\[
q_{ex2} = \begin{cases} 
q_{ex1}, & \text{if } S_{\text{def}} = 0 \\
0, & \text{otherwise}
\end{cases}
\] (269)

Where \(S_{\text{def}} \text{[mm]}\) is the current storage in the soil moisture deficit store. The deficit is increased by evaporation \(E_t \text{[mm/d]}\) and percolation \(q_{ex2} \text{[mm/d]}\). The deficit is decreased by overflow from the upper store \(q_{ex1}\). \(E_t\) only occurs when the upper zone is empty and at a fraction \(\gamma \text{[-]}\) of \(E_p\). \(q_{ex2}\) only occurs when the deficit is zero.

\[
\frac{dS_{u\text{z}}}{dt} = P_{by} + q_{ex2} - q_{u\text{z}}
\] (270)

\[
P_{by} = \phi \ast P_n
\] (271)

\[
q_{u\text{z}} = k_1 \ast S_{u\text{z}}
\] (272)

Where \(S_{u\text{z}}\) is the current storage in the unsaturated zone, refilled by preferential recharge \(P_{by} \text{[mm/d]}\) and percolation \(q_{ex2} \text{[mm/d]}\), and drained by groundwater flow \(q_{u\text{z}} \text{[mm/d]}\). \(P_{by}\) is a fraction \(\phi \text{[-]}\) of \(P_n\). \(q_{u\text{z}}\) has a linear relation with storage through time parameter \(k_1 \text{[d\text{^{-1}}]}\).

\[
\frac{dS_{s\text{z}}}{dt} = q_{u\text{z}} - a - Q
\] (273)

\[
a = c_a
\] (274)

\[
Q = \begin{cases} 
k_2 \ast S_{s\text{z}}^2, & \text{if } S_{s\text{z}} > 0 \\
0, & \text{otherwise}
\end{cases}
\] (275)

Where \(S_{s\text{z}} \text{[mm]}\) is the current storage in the saturated zone, refilled by groundwater flow \(q_{u\text{z}} \text{[mm/d]}\) and drained by abstractions \(a \text{[mm/d]}\) and outflow \(Q \text{[mm/d]}\). \(a\) occurs at a constant rate \(c_a \text{[mm/d]}\). Abstractions can draw down the aquifer below the runoff generating threshold. \(Q\) has a quadratic relation with storage through parameter \(k_2 \text{[mm}^{-1}\text{d}^{-1}]\).
**S2.26 Flex-I (model ID: 26)**

The Flex-I model (fig. S27) is the part of a model development exercise (Fenicia et al., 2008). It has 4 stores and 10 parameters ($I_{max}$, $UR_{max}$, $\beta$, $D$, $Per_{cmax}$, $L_p$, $N_{lag,f}$, $N_{lag,s}$, $K_f$, $K_s$). The model aims to represent:

- Interception by vegetation;
- Infiltration and saturation excess flow based on a distribution of different soil depths;
- A split between fast saturation excess flow and preferential recharge to a slow store;
- Percolation from the unsaturated zone to a slow runoff store.

**S2.26.1 File names**

Model: m_26_flexi_10p_4s

Parameter ranges: m_26_flexi_10p_4s_parameter_ranges

**S2.26.2 Model equations**

Figure S27: Structure of the Flex-I model

\[
\frac{dI}{dt} = P - E_i - P_{eff} \tag{276}
\]

\[
E_i = \begin{cases} 
Ep, & \text{if } I > 0 \\
0 & \text{otherwise}
\end{cases} \tag{277}
\]

\[
P_{eff} = \begin{cases} 
P, & \text{if } I \geq I_{max} \\
0 & \text{otherwise}
\end{cases} \tag{278}
\]

Where $I$ is the current interception storage [mm], $P$ [mm/d] incoming precipitation, $E_i$ [mm/d] evaporation from the interception store and $P_{eff}$ [mm/d] interception excess routed to soil moisture. Evaporation occurs at the potential rate $Ep$ [mm/d] whenever possible. Interception excess occurs when the interception store exceeds its maximum capacity $I_{max}$ [mm].
\[
\frac{dUR}{dt} = Ru - E_{ur} - R_p
\] (279)

\[
RU = (1 - Cr) \times P_{eff}
\] (280)

\[
Cr = \left[ 1 + \exp\left(\frac{-UR/UR_{max} + 1/2}{\beta}\right) \right]^{-1}
\] (281)

\[
E_{ur} = E_p \times \min\left(1, \frac{UR}{UR_{max}} \frac{1}{L_p}\right)
\] (282)

\[
P_s = Perc_{max} \times \frac{-UR}{UR_{max}}
\] (283)

Where UR is the current storage in the unsaturated zone [mm]. Ru [mm/d] is the inflow into UR based on its current storage compared to maximum storage UR_{max} [mm] and a shape distribution parameter \(\beta\) [-]. E_{ur} the evaporation [mm/d] from UR which follows a linear relation between current and maximum storage until a threshold \(L_p\) [-] is exceeded. \(P_s\) is the percolation from UR to the slow reservoir SR [mm/d], based on a maximum percolation rate \(Perc_{max}\) [mm], relative to the fraction of current storage and maximum storage. \(P_{eff}\) is routed towards the unsaturated zone based on \(Cr\), with the remainder being divided into preferential recharge \(R_s\) [mm/d] and fast runoff \(R_f\) [mm/d]:

\[
R_s = (P_{eff} - Ru) \times D
\] (284)

\[
R_f = (P_{eff} - Ru) \times (1 - D)
\] (285)

Where \(R_s\) and \(R_f\) are the flows [mm/d] to the slow and fast runoff reservoir respectively, based on runoff partitioning coefficient D [-]. Both are lagged by linearly increasing triangular transformation functions with parameters \(N_{lag,s}\) [d] and \(N_{lag,f}\) [d] respectively, that give the number of days over which \(R_s\) and \(R_f\) need to be transformed. Percolation \(R_p\) is added to \(R_s\) before the transformation to \(R_{s,l}\) occurs.

\[
\frac{dFR}{dt} = R_{f,l} - Q_f
\] (286)

\[
Q_f = K_f \times FR
\] (287)

Where FR is the current storage [mm] in the fast flow reservoir. Outflow \(Q_f\) [mm/d] from the reservoir has a linear relation with storage through time scale parameter \(K_f\) [d^{-1}].

\[
\frac{dSR}{dt} = R_{s,l} - Q_s
\] (288)

\[
Q_s = K_s \times SR
\] (289)
Where SR is the current storage [mm] in the slow flow reservoir. Outflow $Q_s$ [mm/d] from the reservoir has a linear relation with storage through time scale parameter $K_s$ [d$^{-1}$]. Total outflow $Q$ [mm/d]:

$$Q = Q_f + Q_s$$  \hspace{2cm} (290)
S2.27 Tank model (model ID: 27)

The Tank Model (fig. S28) is originally developed for use constantly saturated soils in Japan (Sugawara, 1995). It has 4 stores and 12 parameters ($A_0, A_1, A_2, t_1, t_2, B_0, B_1, t_3, C_0, C_1, t_4, D_1$). The model aims to represent:

- Runoff on increasing time scales with depth.

S2.27.1 File names

Model: m_27_tank_12p_4s
Parameter ranges: m_27_tank_12p_4s_parameter_ranges

S2.27.2 Model equations

\[
\frac{dS_1}{dt} = P - E_1 - F_{12} - Y_2 - Y_1 \tag{291}
\]

\[
E_1 = \begin{cases} 
E_p, & \text{if } S_1 > 0 \\
0, & \text{otherwise} 
\end{cases} \tag{292}
\]

\[
F_{12} = A_0 \cdot S_1 \tag{293}
\]

\[
Y_2 = \begin{cases} 
A_2 \cdot (S_1 - t_2), & \text{if } S_1 > t_2 \\
0, & \text{otherwise} 
\end{cases} \tag{294}
\]

\[
Y_1 = \begin{cases} 
A_1 \cdot (S_1 - t_1), & \text{if } S_1 > t_1 \\
0, & \text{otherwise} 
\end{cases} \tag{295}
\]

Where $S_1$ [mm] is the current storage in the upper zone, refilled by precipitation $P$ [mm/d] and drained by evaporation $E_1$ [mm/d], drainage $F_{12}$ [mm/d] and surface runoff $Y_1$ [mm/d] and $Y_2$ [mm/d]. $E_1$ occurs at the potential rate $E_p$ [mm/d] if water is available. Drainage to the intermediate layer has a linear relationship with storage through time scale parameter $A_0$ [d$^{-1}$]. Surface runoff $Y_2$ and $Y_1$ occur when $S_1$ is above thresholds $t_2$ [mm] and $t_1$ [mm] respectively. Both are linear relationships through time parameters $A_2$ [d$^{-1}$] and $A_1$ [d$^{-1}$] respectively.

Figure S28: Structure of the Tank Model
\[
\frac{dS_2}{dt} = F_{12} - E_2 - F_{23} - Y_3 \quad (296)
\]

\[
E_2 = \begin{cases} 
E_p, & \text{if } S_1 = 0 & S_2 > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (297)
\]

\[
F_{23} = B_0 \times S_2 \quad (298)
\]

\[
Y_3 = \begin{cases} 
B_1 \times (S_2 - t_3), & \text{if } S_2 > t_3 \\
0, & \text{otherwise} 
\end{cases} \quad (299)
\]

Where \( S_2 \) [mm] is the current storage in the intermediate zone, refilled by drainage \( F_{12} \) from the upper zone and drained by evaporation \( E_2 \) [mm/d], drainage \( F_{23} \) [mm/d] and intermediate discharge \( Y_3 \) [mm/d]. \( E_2 \) occurs at the potential rate \( E_p \) if water is available and the upper zone is empty. Drainage to the third layer \( F_{23} \) has a linear relationship with storage through time scale parameter \( B_0 \) [d\(^{-1}\)]. Intermediate runoff \( Y_3 \) occurs when \( S_2 \) is above threshold \( t_3 \) [mm] and has a linear relationship with storage through time scale parameter \( B_1 \) [d\(^{-1}\)].

\[
\frac{dS_3}{dt} = F_{23} - E_3 - F_{34} - Y_4 \quad (300)
\]

\[
E_3 = \begin{cases} 
E_p, & \text{if } S_1 = 0 & S_2 = 0 & S_3 > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (301)
\]

\[
F_{34} = C_0 \times S_3 \quad (302)
\]

\[
Y_4 = \begin{cases} 
C_1 \times (S_3 - t_4), & \text{if } S_3 > t_4 \\
0, & \text{otherwise} 
\end{cases} \quad (303)
\]

Where \( S_3 \) [mm] is the current storage in the sub-base zone, refilled by drainage \( F_{23} \) from the intermediate zone and drained by evaporation \( E_3 \) [mm/d], drainage \( F_{34} \) [mm/d] and sub-base discharge \( Y_4 \) [mm/d]. \( E_3 \) occurs at the potential rate \( E_p \) if water is available and the upper zones are empty. Drainage to the fourth layer \( F_{34} \) has a linear relationship with storage through time scale parameter \( C_0 \) [d\(^{-1}\)]. Sub-base runoff \( Y_4 \) occurs when \( S_3 \) is above threshold \( t_4 \) [mm] and has a linear relationship with storage through time scale parameter \( C_1 \) [d\(^{-1}\)].

\[
\frac{dS_4}{dt} = F_{34} - E_4 - Y_5 \quad (304)
\]

\[
E_4 = \begin{cases} 
E_p, & \text{if } S_1 = 0 & S_2 = 0 & S_3 = 0 & S_4 > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (305)
\]

\[
Y_5 = D_1 \times S_4 \quad (306)
\]

Where \( S_4 \) [mm] is the current storage in the base layer, refilled by drainage \( F_{34} \) from the sub-base zone and drained by evaporation \( E_4 \) [mm/d] and baseflow \( Y_5 \) [mm/d].
$E_4$ occurs at the potential rate $E_p$ if water is available and the upper zones are empty. Baseflow $Y_5$ has a linear relationship with storage through time scale parameter $D_1$ $[d^{-1}]$. Total runoff:

$$Q_t = Y_1 + Y_2 + Y_3 + Y_4 + Y_5$$

(307)
S2.28 Xinanjiang model (model ID: 28)

The Xinanjiang model (fig. S29) is originally intended for use in humid or semi-humid regions in China (Zhao, 1992). The model uses a variable contributing area to simulate runoff. The version presented here uses a double parabolic curve to simulate tension water capacities within the catchment (Jayawardena and Zhou, 2000), instead of the original single parabolic curve. The model has 4 stores and 12 parameters ($A_{im}$, $a$, $b$, $W_{max}$, $LM$, $c$, $S_{max}$, $Ex$, $k_I$, $k_G$, $c_I$, $c_G$). The model aims to represent:

- Runoff from impervious areas;
- Variable distribution of tension water storage capacities in the catchment;
- Variable contributing area of free water storages;
- Direct surface runoff from the contributing free area;
- Delayed interflow and baseflow from the contributing free area.

S2.28.1 File names

Model: m_28_xinanjiang_12p_4s
Parameter ranges: m_28_xinanjiang_12p_4s_parameter_ranges

S2.28.2 Model equations

\[
\frac{dW}{dt} = P_i - E - R \\
P_i = (1 - A_{im}) \times P \\
R = \begin{cases} 
P_i \times (0.5 - a)^{1-b} \left( \frac{W}{W_{max}} \right)^b, & \text{if } W_{max} \leq 0.5 - a \\
P_i \times (1 - (0.5 + a)^{1-b} \left( 1 - \frac{W}{W_{max}} \right)^b, & \text{otherwise}
\end{cases} \\
E = \begin{cases} 
E_p, & \text{if } W > LM \\
\frac{W}{LM}E_p, & \text{if } c \times LM \geq W \leq LM \\
c \times E_p, & \text{otherwise}
\end{cases}
\]

Figure S29: Structure of the Xinanjiang model

Where $W$ [mm] is the current tension water storage, refilled by an infiltration $P_i$ [mm/d] and drained by evaporation $E$ [mm/d] and runoff $R$ [mm/d]. $P_i$ is the fraction of precipitation $P$ [mm/d] that does not fall on impervious area $A_{im}$ [−]. Runoff...
generation \( R \) uses a double parabolic curve to determine the fraction of catchment area that is at full tension storage and thus can contribute to runoff generation. This curve relies on shape parameters \( a [-] \) and \( b [-] \), and maximum tension water storage \( W_{\text{max}} [\text{mm}] \). Evaporation rate \( E \) declines as tension water storage decreases. Evaporation occurs at the potential rate \( E_p [\text{mm/d}] \) if storage \( W \) is above threshold \( L_M [\text{mm}] \), and reduces linearly below that up to a second threshold \( c * L_M [-][\text{mm}] \). Below this threshold evaporation occurs at a constant rate \( c * E_p \).

\[
dS\frac{dt}{dt} = R - R_S - R_I - R_G \\
R_S = R * \left( 1 - \left( 1 - \frac{S}{S_{\text{max}}} \right)^{Ex} \right) \\
R_I = k_I * S * \left( 1 - \left( 1 - \frac{S}{S_{\text{max}}} \right)^{Ex} \right) \\
R_G = k_G * S * \left( 1 - \left( 1 - \frac{S}{S_{\text{max}}} \right)^{Ex} \right)
\]

Where \( S [\text{mm}] \) is the current storage of free water, refilled by runoff \( R \) from filled tension water areas, and drained by surface runoff \( R_S [\text{mm/d}] \), interflow \( R_I [\text{mm/d}] \) and baseflow \( R_G [\text{mm/d}] \). All runoff components rely on a parabolic equation to simulate variable contributing areas of the catchment, dependent on maximum free water storage \( S_{\text{max}} [\text{mm}] \) and shape parameter \( Ex [-] \). \( R_I \) also uses a time coefficient \( k_I [d^{-1}] \). \( R_G \) uses a time coefficient \( k_G [d^{-1}] \).

\[
dS_{I}\frac{dt}{dt} = R_I - Q_I \\
Q_I = c_I * S_{I}
\]

Where \( S_I [\text{mm}] \) is the current storage in the interflow routing reservoir, filled by interflow from free water \( R_I \) and drained by delayed interflow \( Q_I [\text{mm/d}] \). \( Q_I \) uses a time coefficient \( c_I [d^{-1}] \).

\[
dS_{G}\frac{dt}{dt} = R_G - Q_G \\
Q_G = c_G * S_{G}
\]

Where \( S_G [\text{mm}] \) is the current storage in the baseflow routing reservoir, filled by baseflow from free water \( R_G \) and drained by delayed baseflow \( Q_G [\text{mm/d}] \). \( Q_G \) uses a time coefficient \( c_G [d^{-1}] \). Total flow depends on four separate runoff components:
\[ Q_I = Q_S + Q_I + Q_G \quad (320) \]
\[ Q_S = R_S + R_B \quad (321) \]
\[ R_B = A_{im} \times P \quad (322) \]

Where \( R_B \) [mm/d] is direct runoff generated by precipitation \( P \) [mm/d] on the fraction impervious area \( A_{im} \).
S2.29 HyMOD (model ID: 29)

The HyMOD model (fig. S30) combines a PDM-like soil moisture routine (e.g. Moore (2007)) with a Nash cascade of three linear reservoirs that simulates fast flow and a single linear reservoir intended to simulate slow flow (Wagener et al., 2001; Boyle, 2001). Although the model was originally intended as a flexible structure where the user defines which processes to include, this study includes only a single version that is commonly used. It has 5 parameters ($S_{max}$, $b$, $a$, $k_f$ and $k_s$) and 5 stores. The model aims to represent:

- Different soil depths throughout the catchment;
- Separation of flow into fast and slow flow.

S2.29.1 File names

Model: m_29_hymod_5p_5s
Parameter ranges: m_29_hymod_5p_5s_parameter_ranges

S2.29.2 Model equations

![Figure S30: Structure of the HyMOD model]

\[
\frac{dS_m}{dt} = P - E_a - P_e \tag{323}
\]

\[
E_a = \frac{S_m}{S_{max}} \times E_p \tag{324}
\]

\[
P_e = \left(1 - \left(1 - \frac{S}{S_{max}}\right)^b\right) \times P \tag{325}
\]

Where \( S_m \) is the current storage in \( S_m \) [mm], \( S_{max} \) [mm] is the maximum storage in \( S_m \), \( E_a \) and \( E_p \) the actual and potential evapotranspiration respectively [mm/d] and \( b \) is the soil depth distribution parameter [-]. \( P \) [mm/d] is the precipitation input.

\[
\frac{dF_1}{dt} = P_f - Q_{f,1} \tag{326}
\]

\[
P_f = a \times P_e \tag{327}
\]

\[
Q_{f,1} = k_f \times S_{f,1} \tag{328}
\]

Where \( F_1 \) is the current storage in store \( F_1 \) [mm], \( a \) the fraction of \( P_e \) that flows into the fast stores and \( k_f \) the runoff coefficient of the fast stores. Stores \( F_2 \) and \( F_3 \) take...
the outflow of the previous store as input ($Q_{f,1}$ and $Q_{f,2}$ respectively) and generate outflow analogous to the equations above.

\[
\frac{dS}{dt} = P_s - Q_s 
\]
(329)

\[
P_s = (1 - a) \times P_e 
\]
(330)

\[
Q_s = k_s \times S 
\]
(331)

Where $S$ is the current storage in store $S$ [mm], $1 - a$ [-] the fraction of $P_e$ that flows into the slow store and $k_s$ the runoff coefficient of the slow store. Total outflow:

\[
Q_t = Q_s + Q_f,3 
\]
(332)
S2.30 MOPEX-2 (model ID: 30)
The MOPEX-2 model (fig. S31) is part of a model improvement study that investigates the relationship between dominant processes and model structures for 197 catchments in the MOPEX database (Ye et al., 2012). It has 5 stores and 7 parameters ($T_{\text{crit}}$, $ddf$, $S_{b1}$, $t_w$, $tu$, $S_e$, $tc$). The model aims to represent:

- Snow accumulation and melt;
- Saturation excess flow;
- Infiltration to deeper soil layers;
- A split between fast and slow runoff.

S2.30.1 File names
Model: m_30_mopex2_7p_5s
Parameter ranges: m_30_mopex2_7p_5s_parameter_ranges

S2.30.2 Model equations

\[
\begin{align*}
\frac{dS_n}{dt} &= P_s - Q_n \\
P_s &= \begin{cases} P, & \text{if } T \leq T_{\text{crit}} \\ 0, & \text{otherwise} \end{cases} \\
Q_n &= \begin{cases} ddf \cdot (T - T_{\text{crit}}), & \text{if } T > T_{\text{crit}} \\ 0, & \text{otherwise} \end{cases}
\end{align*}
\]

Where $S_n$ [mm] is the current snow pack. Precipitation occurs as snowfall $P_s$ [mm/d] when current temperature $T$ [$^\circ\text{C}$] is below threshold $T_{\text{crit}}$ [$^\circ\text{C}$]. Snowmelt $Q_N$ [mm/d] occurs when the temperature rises above the threshold temperature and relies in the degree-day factor $ddf$ [mm/$^\circ\text{C}$/d].

Figure S31: Structure of the MOPEX-2 model
\[
d\frac{S_1}{dt} = P_r - ET_1 - Q_{1f} - Q_w \tag{336}
\]

\[
P_r = \begin{cases} 
  P, & \text{if } T > T_{\text{crit}} \\
  0, & \text{otherwise} 
\end{cases} \tag{337}
\]

\[
ET_1 = \frac{S_1}{S_{b1}} * E_p \tag{338}
\]

\[
Q_{1f} = \begin{cases} 
  P, & \text{if } S_1 \geq S_{b1} \\
  0, & \text{otherwise} \tag{339}
\end{cases}
\]

\[
Q_w = t_w * S_1 \tag{340}
\]

Where \( S_1 [\text{mm}] \) is the current storage in soil moisture and \( P_r \) precipitation as rain [\text{mm/d}]. Evaporation \( ET_1 [\text{mm/d}] \) depends linearly on current soil moisture, maximum soil moisture \( S_{b1} [\text{mm}] \) and potential evapotranspiration \( E_p [\text{mm/d}] \). Saturation excess flow \( Q_{1f} [\text{mm/d}] \) occurs when the soil moisture bucket exceeds its maximum capacity. Infiltration to deeper groundwater \( Q_w [\text{mm/d}] \) depends on current soil moisture and time parameter \( t_w [\text{d}^{-1}] \).

\[
d\frac{S_2}{dt} = Q_w - ET_2 - Q_{2u} \tag{341}
\]

\[
ET_2 = \frac{S_2}{S_e} * E_p \tag{342}
\]

\[
Q_{2u} = t_u * S_2 \tag{343}
\]

Where \( S_2 [\text{mm}] \) is the current groundwater storage, refilled by infiltration from \( S_1 \). Evaporation \( ET_2 [\text{mm/d}] \) depends linearly on current groundwater and root zone storage capacity \( S_e [\text{mm}] \). Leakage to the slow runoff store \( Q_{2u} [\text{mm/d}] \) depends on current groundwater level and time parameter \( t_u [\text{d}^{-1}] \).

\[
d\frac{S_{c1}}{dt} = Q_{1f} - Q_f \tag{344}
\]

\[
Q_f = t_c * S_{c1} \tag{345}
\]

Where \( S_{c1} [\text{mm}] \) is current storage in the fast flow routing reservoir, refilled by \( Q_{1f} \). Routed flow \( Q_f \) depends on the mean residence time parameter \( t_c [\text{d}^{-1}] \).

\[
d\frac{S_{c2}}{dt} = Q_{2u} - Q_u \tag{346}
\]

\[
Q_u = t_c * S_{c2} \tag{347}
\]

Where \( S_{c2} [\text{mm}] \) is current storage in the slow flow routing reservoir, refilled by \( Q_{2u} \). Routed flow \( Q_u \) depends on the mean residence time parameter \( t_c [\text{d}^{-1}] \). Total simulated flow \( Q_t [\text{mm/d}] \):
Knoben et al., 2018

\[ Q_t = Q_f + Q_u \]  

(348)
S2.31 MOPEX-3 (model ID: 31)

The MOPEX-3 model (fig. S32) is part of a model improvement study that investigates the relationship between dominant processes and model structures for 197 catchments in the MOPEX database (Ye et al., 2012). It has 5 stores and 8 parameters \( (T_{\text{crit}}, ddf, S_{b1}, t_w, S_{b2}, t_u, S_e, t_c) \). The model aims to represent:

- Snow accumulation and melt;
- Saturation excess flow;
- Infiltration to deeper soil layers;
- Subsurface-influenced fast flow;
- A split between fast and slow runoff.

S2.31.1 File names

Model: m_31_mopex3_8p_5s
Parameter ranges: m_31_mopex3_8p_5s_parameter_ranges

S2.31.2 Model equations

\[
\frac{dS_n}{dt} = P_s - Q_n \tag{349}
\]

\[
P_s = \begin{cases} 
P, & \text{if } T \leq T_{\text{crit}} \\
0, & \text{otherwise}
\end{cases} \tag{350}
\]

\[
Q_n = \begin{cases} 
ddf \times (T - T_{\text{crit}}), & \text{if } T > T_{\text{crit}} \\
0, & \text{otherwise}
\end{cases} \tag{351}
\]

Where \( S_n \) [mm] is the current snow pack. Precipitation occurs as snowfall \( P_s \) [mm/d] when current temperature \( T \) [°C] is below threshold \( T_{\text{crit}} \) [°C]. Snowmelt \( Q_N \) [mm/d] occurs when the temperature rises above the threshold temperature and relies on the degree-day factor \( dd \) [mm/°C/d].

Figure S32: Structure of the MOPEX-3 model
\[
\frac{dS_1}{dt} = P_r - ET_1 - Q_{1f} - Q_w \tag{352}
\]

\[
P_r = \begin{cases} P, & \text{if } T > T_{\text{crit}} \\ 0, & \text{otherwise} \end{cases} \tag{353}
\]

\[
ET_1 = \frac{S_1}{S_{b1}} * Ep \tag{354}
\]

\[
Q_{1f} = \begin{cases} P, & \text{if } S_1 \geq S_{b1} \\ 0, & \text{otherwise} \end{cases} \tag{355}
\]

\[
Q_w = t_w * S_1 \tag{356}
\]

Where \( S_1 \) [mm] is the current storage in soil moisture and \( P_r \) precipitation as rain [mm/d]. Evaporation \( ET_1 \) [mm/d] depends linearly on current soil moisture, maximum soil moisture \( S_{b1} \) [mm] and potential evapotranspiration \( Ep \) [mm/d]. Saturation excess flow \( Q_{1f} \) [mm/d] occurs when the soil moisture bucket exceeds its maximum capacity. Infiltration to deeper groundwater \( Q_w \) [mm/d] depends on current soil moisture and time parameter \( t_w \) [d^{-1}].

\[
\frac{dS_2}{dt} = Q_w - ET_2 - Q_{2u} - Q_{2f} \tag{357}
\]

\[
ET_2 = \frac{S_2}{S_{e}} * Ep \tag{358}
\]

\[
Q_{2u} = t_u * S_2 \tag{359}
\]

\[
Q_{2f} = \begin{cases} Q_w, & \text{if } S_2 \geq S_{b2} \\ 0, & \text{otherwise} \end{cases} \tag{360}
\]

Where \( S_2 \) [mm] is the current groundwater storage, refilled by infiltration from \( S_1 \). Evaporation \( ET_2 \) [mm/d] depends linearly on current groundwater and root zone storage capacity \( S_e \) [mm]. Leakage to the slow runoff store \( Q_{2u} \) [mm/d] depends on current groundwater level and time parameter \( t_u \) [d^{-1}]. When the store reaches maximum capacity \( S_{b2} \) [mm], excess flow \( Q_{2f} \) [mm/d] is routed towards the fast response routing store.

\[
\frac{dS_{c1}}{dt} = Q_{1f} + Q_{2f} - Q_f \tag{361}
\]

\[
Q_f = t_c * S_{c1} \tag{362}
\]

Where \( S_{c1} \) [mm] is current storage in the fast flow routing reservoir, refilled by \( Q_{1f} \) and \( Q_{2f} \). Routed flow \( Q_f \) depends on the mean residence time parameter \( t_c \) [d^{-1}].
\[
\frac{dS_{c2}}{dt} = Q_{2u} - Q_u \tag{363}
\]
\[
Q_u = t_c \ast S_{c2} \tag{364}
\]

Where \( S_{c2} \) [mm] is current storage in the slow flow routing reservoir, refilled by \( Q_{2u} \). Routed flow \( Q_u \) depends on the mean residence time parameter \( t_c \) [d\(^{-1}\)]. Total simulated flow \( Q_t \) [mm/d]:

\[
Q_t = Q_f + Q_u \tag{365}
\]
S2.32 MOPEX-4 (model ID: 32)

The MOPEX-4 model (fig. S33) is part of a model improvement study that investigates the relationship between dominant processes and model structures for 197 catchments in the MOPEX database (Ye et al., 2012). It has 5 stores and 10 parameters ($T_{crit}$, $ddf$, $S_{b1}$, $t_w$, $I_s$, $I$, $S_{b2}$, $t_u$, $S_e$, $t_c$). The original model relies on observations of Leaf Area Index and a calibrated interception fraction. Liang et al. (1994) show typical Leaf Area Index time series, and a sinusoidal function is a reasonable approximation of this. Therefore, the model is slightly modified to use a calibrated sinusoidal function, so that the data input requirements for MOPEX-4 are consistent with other models. The model aims to represent:

- Snow accumulation and melt;
- Time-varying interception;
- Saturation excess flow;
- Infiltration to deeper soil layers;
- A split between fast and slow runoff.

S2.32.1 File names

Model: m_32_mopex4_10p_5s
Parameter ranges: m_32_mopex4_10p_5s_parameter_ranges

S2.32.2 Model equations

\[
\frac{dS_n}{dt} = P_s - Q_n \tag{366}
\]

\[
P_s = \begin{cases} P, & \text{if } T \leq T_{crit} \\ 0, & \text{otherwise} \end{cases} \tag{367}
\]

\[
Q_n = \begin{cases} ddf \ast (T - T_{crit}), & \text{if } T > T_{crit} \\ 0, & \text{otherwise} \end{cases} \tag{368}
\]

Where $S_n$ [mm] is the current snow pack. Precipitation occurs as snowfall $P_s$ [mm/d] when current temperature $T$ [$^\circ$C] is below threshold $T_{crit}$ [$^\circ$C]. Snowmelt $Q_N$ [mm/d] occurs when the temperature rises above the threshold temperature and relies in the degree-day factor $ddf$ [mm/$^\circ$C/d].

Figure S33: Structure of the MOPEX-4 model
\[ \frac{dS_1}{dt} = P_r - ET_1 - I - Q_{1f} - Q_w \]  
\[ P_r = \begin{cases} P, & \text{if } T > T_{\text{crit}} \\ 0, & \text{otherwise} \end{cases} \]  
\[ ET_1 = \frac{S_1}{S_{b1}} \cdot Ep \]  
\[ I = \max \left(0, I_\alpha + (1 - I_\alpha) \cos \left(2\pi \frac{t - I_s}{t_{\text{max}}} \right) \right) \cdot P_r \]  
\[ Q_{1f} = \begin{cases} P, & S_1 \geq S_{b1} \\ 0, & \text{otherwise} \end{cases} \]  
\[ Q_w = t_w \cdot S_1 \]  

Where \( S_1 [\text{mm}] \) is the current storage in soil moisture and \( P_r \) precipitation as rain \([\text{mm/d}]\). Evaporation \( ET_1 [\text{mm/d}] \) depends linearly on current soil moisture, maximum soil moisture \( S_{b1} [\text{mm}] \) and potential evapotranspiration \( Ep [\text{mm/d}] \). Evaporation \( ET_1 [\text{mm/d}] \) depends on the mean intercepted fraction \( I_\alpha [-] \), the maximum Leaf Area Index timing \( I_s [\text{d}] \) and the length of the seasonal cycle \( t_{\text{max}} [\text{d}] \) (usually set at 365 days). Saturation excess flow \( Q_{1f} [\text{mm/d}] \) occurs when the soil moisture bucket exceeds its maximum capacity. Infiltration to deeper groundwater \( Q_w [\text{mm/d}] \) depends on current soil moisture and time parameter \( t_w [\text{d}^{-1}] \).

\[ \frac{dS_2}{dt} = Q_w - ET_2 - Q_{2u} - Q_{2f} \]  
\[ ET_2 = \frac{S_2}{S_{e}} \cdot Ep \]  
\[ Q_{2u} = t_u \cdot S_2 \]  
\[ Q_{2f} = \begin{cases} Q_w, & S_2 \geq S_{b2} \\ 0, & \text{otherwise} \end{cases} \]  

Where \( S_2 [\text{mm}] \) is the current groundwater storage, refilled by infiltration from \( S_1 \). Evaporation \( ET_2 [\text{mm/d}] \) depends linearly on current groundwater and root zone storage capacity \( S_e [\text{mm}] \). Leakage to the slow runoff store \( Q_{2u} [\text{mm/d}] \) depends on current groundwater level and time parameter \( t_u [\text{d}^{-1}] \). When the store reaches maximum capacity \( S_{b2} [\text{mm}] \), excess flow \( Q_{2f} [\text{mm/d}] \) is routed towards the fast response routing store.

\[ \frac{dS_{c1}}{dt} = Q_{1f} + Q_{2f} - Q_f \]  
\[ Q_f = t_c \cdot S_{c1} \]
Where $S_{c1} [\text{mm}]$ is current storage in the fast flow routing reservoir, refilled by $Q_{1f}$ and $Q_{2f}$. Routed flow $Q_f$ depends on the mean residence time parameter $t_c [d^{-1}]$.

$$\frac{dS_{c2}}{dt} = Q_{2u} - Q_u \quad (381)$$
$$Q_u = t_c \ast S_{c2} \quad (382)$$

Where $S_{c2} [\text{mm}]$ is current storage in the slow flow routing reservoir, refilled by $Q_{2u}$. Routed flow $Q_u$ depends on the mean residence time parameter $t_c [d^{-1}]$. Total simulated flow $Q_t [\text{mm/d}]$:

$$Q_t = Q_f + Q_u \quad (383)$$
S2.33 SACRAMENTO model (model ID: 33)

The SACRAMENTO model (fig. S34) is part of an ongoing model development project by the National Weather Service, which started several decades ago (Burnash, 1995; National Weather Service, 2005). The documentation mentions a specific order of flux computations. For consistency with other models, here all fluxes are computed simultaneously. It has 5 stores and 13 parameters ($PCTIM$, $UZTWM$, $UZFWM$, $k_{uz}$, $PBASE$, $ZPERC$, $REXP$, $LZTW$, $LZFWPM$, $LZFWSM$, $PFREE$, $k_{zp}$, $k_{zs}$). The model also uses several coefficients derived from the calibration parameters (Koren et al., 2000): $PBASE$ and $ZPERC$. The model aims to represent:

- Impervious and direct runoff;
- Within soil division of water storage between tension and free water;
- Surface runoff, interflow and percolation to deeper soil layers;
- Multiple baseflow processes.

S2.33.1 File names

Model: m_33_sacramento_11p_5s
Parameter ranges: m_33_sacramento_11p_5s_parameter_ranges

S2.33.2 Model equations

\[
\frac{dUZTW}{dt} = P_{eff} + R_u - E_{uztw} - Tw_{ex,u}
\]  
(384)

\[P_{eff} = (1 - PCTIM) \times P\]  
(385)

\[R_u = \begin{cases} 
UZTWM \times UZFW - UZFWM \times UZTW, \\
UZTWM + UZFWM \times Tw_{ex,u}, \\
0, \quad \text{otherwise}
\end{cases}
\]  
(386)

\[E_{uztw} = \frac{UZTW}{UZTWM + UZFWM} \times E_p\]  
(387)

\[Tw_{ex,u} = \begin{cases} 
P_{eff}, \quad \text{if } UZTW = UZTWM \\
0, \quad \text{otherwise}
\end{cases}\]  
(388)

Where $UZTW$ [mm] is upper zone tension water, refilled by effective precipitation.

Figure S34: Structure of the SACRAMENTO model
$P_{eff} \ [mm/d]$ and redistribution of free water $R_u \ [mm/d]$, and is drained by evaporation $E_{uztw} \ [mm/d]$ and tension water excess $T_{wex,u} \ [mm/d]$. $P_{eff}$ is the fraction \((1 - PCTIM) \ [-]\) of precipitation $P$ that does not fall on impervious fraction $PCTIM$ [-]. $R_u$ is only active when the relative deficit in tension water is greater than that in free water, and rebalances the available water in the upper zone. This uses the current storages, $UZTW$ and $UZFW$, and maximum storages, $UZTWM$ [mm] and $UZFWM$ [mm], of tension and free water stores respectively. Evaporation is determined with a linear relation between available, maximum upper zone tension storage and potential evapotranspiration $E_p \ [mm/d]$. $T_{wex,u}$ occurs only when the store is at maximum capacity.

$$\frac{dUZFW}{dt} = T_{wex,u} - E_{uzfw} - Q_{sur} - Q_{int} - Pc - R_u$$ \hspace{1cm} (389)

$$E_{uzfw} = \begin{cases} E_p - E_{uztw}, & \text{if } UZFW > 0 \& E_p > E_{uztw} \\ 0, & \text{otherwise} \end{cases}$$ \hspace{1cm} (390)

$$Q_{sur} = \begin{cases} T_{wex,u}, & \text{if } UZFW = UZFWM \\ 0, & \text{otherwise} \end{cases}$$ \hspace{1cm} (391)

$$Q_{int} = k_{uz} \cdot UZFW$$ \hspace{1cm} (392)

$$Pc = Pc_{demand} \cdot \frac{UZFW}{UZFWM}$$ \hspace{1cm} (393)

$$Pc_{demand} = PBASE \cdot \left(1 + ZPERC \cdot \left(\frac{\sum LZ_{deficiency}}{\sum LZ_{capacity}}\right)^{1 + REXP}\right)$$ \hspace{1cm} (394)

$$LZ_{deficiency} = [LZTWM - LZTW] + [LZFWPM - LZWFP] + [LZFWSM - LZFWS]$$ \hspace{1cm} (395)

$$LZ_{capacity} = LZTWM + LZFWPM + LZFWSM$$ \hspace{1cm} (396)

Where $UZFW$ [mm] is upper zone free water, refilled by excess water $T_{wex,u}$ that can not be stored as tension water, and drained by evaporation $E_{uzfw}$ [mm/d], surface runoff $Q_{sur}$ [mm/d], interflow $Q_{int}$ [mm/d], and percolation to deeper groundwater $Pc$ [mm/d]. Evaporative demand unmet by the upper tension water store is taken from upper free water storage at the potential rate. $Q_{sur}$ occurs only when the store is at maximum capacity $UZFWM$ [mm]. $Q_{int}$ uses time coefficient $k_{uz}$ [$d^{-1}$] to simulate interflow. Percolation $Pc$ is calculated as a balance between the fraction water
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availability in upper zone free storage, and demand from the lower zone \( P_{\text{demand}} \). The demand can be between a base percolation rate \( P_{\text{BASE}} \) [mm/d] and an upper limit of \( Z_{\text{PERC}} \) \([-]\) times \( P_{\text{BASE}} \). This demand is scaled by the ratio between total deficiency and maximum lower zone storage. \( LZTWM \) [mm], \( LZWFP \) [mm], \( LZFWS \) [mm] are the maximum capacity of the lower zone tension store, primary free water store and supplemental free water store respectively. The lower zone percolation demand is potentially non-linear through exponent \( \text{REGX} \) \([-\] . \( P_{\text{BASE}} \) is calculated as \( k_{\text{lp}} \cdot LZWFP + k_{\text{ls}} \cdot LZFWS \).

\[
\frac{dLZTW}{dt} = Pc_{tw} + R_{l,p} + R_{l,s} - E_{ltw} - Tw_{ex,l} \tag{397}
\]

\[
Pc_{tw} = (1 - PFREE) \cdot Pc \tag{398}
\]

\[
R_{l,p} = \begin{cases} 
LZWFP \cdot \frac{LZTW(LZWFP + LZFWS)}{(LZWFP + LZFWS)(LZTW + LZWFP + LZFWS)}, & \text{if } LZTW < \frac{LZWFP + LZFWS}{LZTW} \\
0, & \text{otherwise}
\end{cases} \tag{399}
\]

\[
R_{l,s} = \begin{cases} 
LZWFS \cdot \frac{LZTW(LZWFP + LZFWS)}{(LZWFP + LZFWS)(LZTW + LZWFP + LZFWS)}, & \text{if } LZTW < \frac{LZWFP + LZFWS}{LZTW} \\
0, & \text{otherwise}
\end{cases} \tag{400}
\]

\[
E_{ltw} = \begin{cases} 
(E_p - E_{uztw} - E_{uzfw}) \cdot \frac{LZTW}{LZTW + LZTWM}, & \text{if } LZTW > 0 \& E_p > (E_{uztw} + E_{uzfw}) \\
0, & \text{otherwise}
\end{cases} \tag{401}
\]

\[
Tw_{ex,l} = \begin{cases} 
Pc_{tw}, & \text{if } LZTW = LZTWM \\
0, & \text{otherwise}
\end{cases} \tag{402}
\]

Where \( LZTW \) [mm] is lower zone tension water, refilled by percolation \( Pc_{tw} \) [mm/d] and drained by evaporation \( E_{ztw} \) [mm/d] and tension water excess \( Tw_{ex,l} \) [mm/d]. Evaporative demand unmet by the upper zone can be satisfied from the lower zone tension water store, scaled by the current lower zone storage relative to total tension zone storage. Both \( R_{l,p} \) and \( R_{l,s} \) are only active when the relative deficit in tension water is greater than that in free water, and rebalances the available water in the lower zone. This uses the current storages, \( LZTW \), \( LZWFP \) and \( LZFWS \), and maximum storages, \( LZTWM \) [mm], \( LZWFP \) [mm] and \( LZFWS \) [mm], of the tension and free water stores respectively. \( Pc_{tw} \) is the fraction \( (1 - PFREE) \) \([-\] of percolation \( Pc \) that does not go into free storage. \( Tw_{ex,l} \) occurs only when the store is at maximum capacity \( LZTWM \) [mm].
\[
\frac{dLZFWP}{dt} = PC_{fw,p} + TW_{ex,lp} - Q_{bfp}
\]

(403)

\[
PC_{fw,p} = \left[ \frac{LZFWSM - LZFWP}{LZFWSM} \right] \left[ \frac{LZFWPM - LZFWP}{LZFWPM + LZFWSM - LZFWS} \right] \ast (PFREE \ast PC)
\]

(404)

\[
TW_{ex,lp} = \left[ \frac{LZFWSM - LZFWP}{LZFWSM} \right] \left[ \frac{LZFWPM - LZFWP}{LZFWPM + LZFWSM - LZFWS} \right] \ast TW_{ex,l}
\]

(405)

\[
Q_{bfp} = k_{lzp} \ast LZFWP
\]

(406)

Where \(LZFWP\) [mm] is current storage in the primary lower zone free water store, refilled by excess tension water \(TW_{ex,lp}\) [mm/d] and percolation \(PC_{fw,p}\) [mm/d] and drained by primary baseflow \(Q_{bfp}\) [mm/d]. Refilling of both lower zone free water stores (primary and supplemental) is divided between the two based on their relative, scaled moisture deficiency. Percolation from the upper zone \(PC_{fw,p}\) is scaled according to the relative current moisture deficit \(\frac{LZFWSM - LZFWP}{LZFWSM - LZFWS}\) compared to the total relative deficit in the lower free water stores \(\frac{LZFWPM - LZFWP + LZFWSM - LZFWS}{LZFWSM - LZFWS}\). \(TW_{ex,lp}\) is a similarly scaled part of \(TW_{ex,l}\). \(Q_{bfp}\) uses time parameter \(K_{lzp} [d^{-1}]\) to estimate primary baseflow.

\[
\frac{dLZFWS}{dt} = PC_{fw,s} + TW_{ex,ls} - Q_{bfs}
\]

(407)

\[
PC_{fw,s} = \left[ \frac{LZFWSM - LZFWS}{LZFWSM} \right] \left[ \frac{LZFWPM - LZFWP}{LZFWPM + LZFWSM - LZFWS} \right] \ast (PFREE \ast PC)
\]

(408)

\[
TW_{ex,ls} = \left[ \frac{LZFWSM - LZFWS}{LZFWSM} \right] \left[ \frac{LZFWPM - LZFWP}{LZFWPM + LZFWSM - LZFWS} \right] \ast TW_{ex,l}
\]

(409)

\[
Q_{bfs} = k_{lzs} \ast LZFWS
\]

(410)

Where \(LZFWS\) [mm] is current storage in the supplemental free water lower zone store, refilled by excess tension water \(TW_{ex,ls}\) [mm/d] and percolation \(PC_{fw,s}\) [mm/d], and drained by supplemental baseflow \(Q_{bfs}\) [mm/d]. \(PC_{fw,s}\) is determined based on relative deficits in the lower zone free stores, as is \(TW_{ex,ls}\). \(Q_{bfs}\) uses time parameter \(K_{lzs} [d^{-1}]\) to estimate supplementary baseflow. Total simulated outflow:

\[
Q_t = Q_{dir} + Q_{sur} + Q_{int} + Q_{bfp} + Q_{bfs}
\]

(411)
S2.34 FLEX-IS (model ID: 34)

The FLEX-IS model (fig. S35) is a combination of the FLEX-B model expanded with an interception (I) routine (Fenicia et al., 2008) and a snow (S) module (Nijzink et al., 2016). It has 5 stores and 12 parameters ($TT$, $ddf$, $I_{max}$, $UR_{max}$, $\beta$, $L_p$, $Perc_{max}$, $D$, $N_{lag,f}$, $N_{lag,s}$, $K_f$, $K_s$). The model aims to represent:

- Snow accumulation and melt;
- Interception by vegetation;
- Infiltration and saturation excess flow based on a distribution of different soil depths;
- A split between fast saturation excess flow and preferential recharge to a slow store;
- Percolation from the unsaturated zone to a slow runoff store.

S2.34.1 File names

Model: m_34_flexis_12p_5s
Parameter ranges: m_34_flexis_12p_5s_parameter_ranges

S2.34.2 Model equations

\[
\frac{dS}{dt} = P_s - M 
\]

(412)

\[
P_s = \begin{cases} 
P, & \text{if } T \leq TT \\
0, & \text{otherwise} 
\end{cases} 
\]

(413)

\[
M = \begin{cases} 
ddf \times (T - TT), & \text{if } T \geq TT \\
0, & \text{otherwise} 
\end{cases} 
\]

(414)

Where $S$ [mm] is the current snow storage, $P_s$ the precipitation that falls as snow [mm/d], $M$ the snowmelt [mm/d] based on a degree-day factor ($ddf$, [mm/°C/d]) and threshold temperature for snowfall and snowmelt (TT, [°C]).

Figure S35: Structure of the FLEX-IS model
\[
\frac{dI}{dt} = P_I + M - E_I - P_{eff} \quad (415)
\]
\[
P_I = \begin{cases} 
P, & \text{if } T > T_T \\
0, & \text{otherwise}
\end{cases} \quad (416)
\]
\[
E_I = \begin{cases} 
E_p, & \text{if } I > 0 \\
0, & \text{otherwise}
\end{cases} \quad (417)
\]
\[
P_{eff} = \begin{cases} 
P_I, & \text{if } I = I_{max} \\
0, & \text{otherwise}
\end{cases} \quad (418)
\]

Where \( P_I [\text{mm/d}] \) is the incoming precipitation, \( I \) is the current interception storage [mm], which is assumed to evaporate \( (E_I [\text{mm/d}]) \) at the potential rate \( E_p [\text{mm/d}] \) when possible. When \( I \) exceeds the maximum interception storage \( I_{max} [\text{mm}] \), water is routed to the rest of the model as \( P_{eff} [\text{mm/d}] \).

\[
\frac{dUR}{dt} = R_u - E_{ur} - R_p \quad (419)
\]
\[
R_u = (1 - \left[ 1 + \exp\left(\frac{-UR/UR_{max} + 1/2}{\beta}\right)\right]^{-1}) \cdot P_{eff} \quad (420)
\]
\[
E_{ur} = E_p \cdot \min\left(1, \frac{UR}{UR_{max} L_p}\right) \quad (421)
\]
\[
R_p = Perc_{max} \cdot \frac{-UR}{UR_{max}} \quad (422)
\]

Where \( UR \) is the current storage in the unsaturated zone [mm]. \( R_u [\text{mm/d}] \) is the inflow into \( UR \) based on its current storage compared to maximum storage \( UR_{max} [\text{mm}] \) and a shape distribution parameter \( \beta [-] \). \( E_{ur} \) the evaporation \( [\text{mm/d}] \) from \( UR \) which follows a linear relation between current and maximum storage until a threshold \( L_p [-] \) is exceeded. \( R_p [\text{mm/d}] \) is the percolation from \( UR \) to the slow reservoir \( SR [\text{mm}] \), based on a maximum percolation rate \( Perc_{max} [\text{mm/d}] \), relative to the fraction of current storage and maximum storage.

\[
R_s = (P_{eff} - R_u) \cdot D \quad (423)
\]
\[
R_f = (P_{eff} - R_u) \cdot (1 - D) \quad (424)
\]

Where \( R_s \) and \( R_f \) are the flows \( [\text{mm/d}] \) to the slow and fast runoff reservoir respectively, based on runoff partitioning coefficient \( D [-] \). Both are lagged by linearly increasing triangular transformation functions with parameters \( N_{lag,s} \) and \( N_{lag,f} \) respectively, that give the number of time steps over which \( R_s \) and \( R_f \) need to be transformed. \( R_p \) is added to \( R_s \) before the transformation occurs.
\[
\frac{dFR}{dt} = R_{f,l} - Q_f 
\] (425)
\[
Q_f = K_f \times FR 
\] (426)

Where FR is the current storage [mm] in the fast flow reservoir. Outflow \( Q_f \) [mm/d] from the reservoir has a linear relation with storage through time scale parameter \( K_f \) [d\(^{-1}\)].

\[
\frac{dSR}{dt} = R_{s,l} - Q_s 
\] (427)
\[
Q_s = K_s \times SR 
\] (428)

Where SR is the current storage [mm] in the slow flow reservoir. Outflow \( Q_s \) [mm/d] from the reservoir has a linear relation with storage through time scale parameter \( K_s \) [d\(^{-1}\)].

\[
Q = Q_f + Q_s 
\] (429)

Where \( Q \) [mm/d] is the total simulated flow as the sum of \( Q_s \) and \( Q_f \).
S2.35 MOPEX-5 (model ID: 35)

The MOPEX-5 model (fig. S36) is part of a model improvement study that investigates the relationship between dominant processes and model structures for 197 catchments in the MOPEX database (Ye et al., 2012). It has 5 stores and 12 parameters ($T_{\text{crit}}, ddf, S_b1, t_w, I_o, I_s, T_{\text{min}}, T_{\text{max}}, S_b2, t_u, S_e, t_c$). The original model relies on observations of Leaf Area Index and a calibrated interception fraction. Liang et al. (1994) show typical Leaf Area Index time series, and a sinusoidal function is a reasonable approximation of this. Therefore, the model is slightly modified to use a calibrated sinusoidal function, so that the data input requirements for MOPEX-5 are consistent with other models. The model aims to represent:

- Snow accumulation and melt;
- Time-varying interception and the impact of phenology on transpiration;
- Saturation excess flow;
- Infiltration to deeper soil layers;
- A split between fast and slow runoff.

S2.35.1 File names

Model: m_35_mopex5_12p_5s
Parameter ranges: m_35_mopex5_12p_5s_parameter_ranges

S2.35.2 Model equations

\[
\frac{dS_n}{dt} = P_s - Q_n \tag{430}
\]

\[
P_s = \begin{cases} 
P, & \text{if } T \leq T_{\text{crit}} \\
0, & \text{otherwise} \end{cases} \tag{431}
\]

\[
Q_n = \begin{cases} 
ddf \times (T - T_{\text{crit}}), & \text{if } T > T_{\text{crit}} \\
0, & \text{otherwise} \end{cases} \tag{432}
\]

Where $S_n$ [mm] is the current snow pack. Precipitation occurs as snowfall $P_s$ [mm/d] when current temperature $T$ [$^\circ$C] is below threshold $T_{\text{crit}}$ [$^\circ$C]. Snowmelt $Q_n$ [mm/d] occurs when the temperature rises above the threshold temperature and relies in the degree-day factor $dd$ [mm/$^\circ$C/d].

Figure S36: Structure of the MOPEX-5 model
\[
\frac{dS_1}{dt} = P_r - ET_1 - I - Q_{1f} - Q_w
\] (433)

\[
P_r = \begin{cases} P, & \text{if } T > T_{\text{crit}} \\ 0, & \text{otherwise} \end{cases}
\] (434)

\[
ET_{c1} = \frac{S_1}{S_{b1}} \ast Ep_c
\] (435)

\[
I = \max \left(0, I_\alpha + (1 - I_\alpha) \sin \left(2\pi \frac{t + I_s}{365/d}\right)\right)
\] (436)

\[
Q_{1f} = \begin{cases} P, & \text{if } S_1 \geq S_{b1} \\ 0, & \text{otherwise} \end{cases}
\] (437)

\[
Q_w = t_w \ast S_1
\] (438)

Where \(S_1 \text{ [mm]}\) is the current storage in soil moisture and \(P_r\) precipitation as rain [mm/d]. Evaporation \(ET_1\) [mm/d] depends linearly on current soil moisture, maximum soil moisture \(S_{b1} \text{ [mm]}\) and phenology-corrected potential evapotranspiration:\n
\[
EP_c = Ep \ast GSI
\] (439)

\[
GSI = \begin{cases} 0, & \text{if } T < T_{\text{min}} \\ \frac{T - T_{\text{min}}}{T_{\text{max}} - T_{\text{min}}}, & \text{if } T_{\text{min}} \geq T < T_{\text{max}} \\ 1, & \text{if } T \geq T_{\text{max}} \end{cases}
\] (440)

Where GSI is a growing season index based on parameters \(T_{\text{min}} \text{ [°C]}\) and \(T_{\text{max}} \text{ [°C]}\). Interception \(I\) [mm/d] depends on the mean intercepted fraction \(I_\alpha \) and the maximum Leaf Area Index timing \(I_s\) [d]. Saturation excess flow \(Q_{1f}\) [mm/d] occurs when the soil moisture bucket exceeds its maximum capacity. Infiltration to deeper groundwater \(Q_w\) [mm/d] depends on current soil moisture and time parameter \(t_w\) [d−1].

\[
\frac{dS_2}{dt} = Q_w - ET_2 - Q_{2u} - Q_{2f}
\] (441)

\[
ET_{c2} = \frac{S_2}{S_{c2}} \ast Ep_c
\] (442)

\[
Q_{2u} = t_u \ast S_2
\] (443)

\[
Q_{2f} = \begin{cases} Q_w, & \text{if } S_2 \geq S_{b2} \\ 0, & \text{otherwise} \end{cases}
\] (444)

Where \(S_2 \text{ [mm]}\) is the current groundwater storage, refilled by infiltration from \(S_1\). Evaporation \(ET_2\) [mm/d] depends linearly on current groundwater and root zone
storage capacity $S_e$ [mm]. Leakage to the slow runoff store $Q_{2u}$ [mm/d] depends on current groundwater level and time parameter $t_u$ [d$^{-1}$]. When the store reaches maximum capacity $S_{b2}$ [mm], excess flow $Q_{2f}$ [mm/d] is routed towards the fast response routing store.

$$\frac{dS_{c1}}{dt} = Q_{1f} + Q_{2f} - Q_f$$  \hspace{1cm} (445)

$$Q_f = t_c * S_{c1}$$  \hspace{1cm} (446)

Where $S_{c1}$ [mm] is current storage in the fast flow routing reservoir, refilled by $Q_{1f}$ and $Q_{2f}$. Routed flow $Q_f$ depends on the mean residence time parameter $t_c$ [d$^{-1}$].

$$\frac{dS_{c2}}{dt} = Q_{2u} - Q_u$$  \hspace{1cm} (447)

$$Q_u = t_c * S_{c2}$$  \hspace{1cm} (448)

Where $S_{c2}$ [mm] is current storage in the slow flow routing reservoir, refilled by $Q_{2u}$. Routed flow $Q_u$ depends on the mean residence time parameter $t_c$ [d$^{-1}$]. Total simulated flow $Q_t$ [mm/d]:

$$Q_t = Q_f + Q_u$$  \hspace{1cm} (449)
S2.36 MODHYDROLOG (model ID: 36)

The MODHYDROLOG model (fig. S37) is an elaborate groundwater recharge model, originally created for use in Australia (Chiew, 1990; Chiew and McMahon, 1994). It has 5 stores (I, D, SMS, GW and CH) and 15 parameters (INSC, COEFF, SQ, SMSC, SUB, CRAK, EM, DSC, ADS, MD, VCOND, DLEV, \( k_1 \), \( k_2 \) and \( k_3 \)). It originally includes a routing scheme that allows linking sub-basins together, which has been removed here. The model aims to represent:

- Interception by vegetation;
- Infiltration and infiltration excess flow;
- Depression storage and delayed infiltration;
- Preferential groundwater recharge, interflow and saturation excess flow;
- Groundwater recharge resulting from filling up of soil moisture storage capacity;
- Water exchange between shallow and deep aquifers;
- Water exchange between aquifer and river channel.

S2.36.1 File names

Model: m_36_modhydrolog_15p_5s
Parameter ranges: m_36_modhydrolog_15p_5s_parameter_ranges

S2.36.2 Model equations

\[
\frac{dI}{dt} = P - E_i - EXC \quad (450)
\]

\[
E_i = \begin{cases} 
E_p, & \text{if } I > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (451)
\]

\[
EXC = \begin{cases} 
P, & \text{if } I = INSC \\
0, & \text{otherwise} 
\end{cases} \quad (452)
\]

Where \( I \) [mm] is the current interception storage, \( P \) the rainfall [mm/d], \( E_i \) the evaporation from the interception store [mm/d] and \( EXC \) the excess rainfall [mm/d]. Evaporation is assumed to occur at the potential rate \( E_p \) [mm/d] when possible. When \( I \) exceeds the maximum interception capacity \( INSC \), water is routed to the rest of the model as excess precipitation \( EXC \). The soil moisture store...
SMS is instrumental in dividing runoff between infiltration and surface flow:

\[
\frac{d\text{SMS}}{dt} = \text{SMF} + \text{DINF} - \text{ET} - \text{GWF}
\]

(453)

\[
\text{SMF} = \text{INF} - \text{INT} - \text{REC}
\]

(454)

\[
\text{INF} = \min \left( \text{COEFF} \exp \left( -\text{SQ} \cdot \frac{\text{SMS}}{\text{SMSC}} \right), \text{EXC} \right)
\]

(455)

\[
\text{INT} = \text{SUB} \cdot \frac{\text{SMS}}{\text{SMSC}} \cdot \text{INF}
\]

(456)

\[
\text{REC} = \text{CRAK} \cdot \frac{\text{SMS}}{\text{SMSC}} \cdot (\text{INF} - \text{INT})
\]

(457)

\[
\text{ET} = \min \left( \text{EM} \cdot \frac{\text{SMS}}{\text{SMSC}}, \text{PET} \right)
\]

(458)

\[
\text{GWF} = \begin{cases} 
\text{SMF}, & \text{if } \text{SMS} = \text{SMSC} \\
0, & \text{otherwise}
\end{cases}
\]

(459)

Where SMS is the current storage in the soil moisture store [mm]. SMF [mm/d] and DINF [mm/d] are the infiltration and delayed infiltration respectively. INF is total infiltration [mm/d] from excess precipitation, based on maximum infiltration loss parameter COEFF [-], the infiltration loss exponent SQ and the ratio between current soil moisture storage SMS [mm] and the maximum soil moisture capacity SMSC [mm]. INT represents interflow and saturation excess flow [mm/d], using a constant of proportionality SUB [-]. REC is preferential recharge of groundwater [mm/d] based on another constant of proportionality CRAK [-]. SMF is flow into soil moisture storage [mm/d]. ET evaporation from the soil moisture that occurs at the potential rate when possible [mm/d], based on the maximum plant-controlled rate EM [mm/d]. GWF is the flow to the groundwater store [mm/d]:
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\[ \frac{dD}{dt} = TRAP - E_D - DINF \] (460)

\[ TRAP = ADS \ast \exp\left(-MD \frac{D}{DSC - D}\right) \ast RUN \] (461)

\[ RUN = EXC - INF \] (462)

\[ E_D = \begin{cases} ADS \ast E_p, & \text{if } D > 0 \\ 0, & \text{otherwise} \end{cases} \] (463)

\[ DINF = \begin{cases} ADS \ast RATE, & \text{if } D > 0 \\ 0, & \text{otherwise} \end{cases} \] (464)

\[ RATE = COEFF \ast \exp\left(-SQ \frac{SMS}{SMSC}\right) - INF - INT - REC \] (465)

Where TRAP [mm/d] is the part of overland flow captured in the depression store (equation taken from Porter and McMahon (1971)), \( E_D \) the evaporation from the depression store [mm/d], and DINF delayed infiltration to soil moisture [mm/d]. TRAP uses DSC as the maximum depression store capacity [mm], ADS as the fraction of land functioning as depression storage [-] and MD a depression storage parameter [-]. \( E_D \) relies on the potential evapotranspiration \( E_p \). The groundwater store has no defined upper and lower boundary and instead fluctuates around a datum DLEV:

\[ \frac{dGW}{dt} = REC + GWF - SEEP - FLOW \] (466)

\[ SEEP = VCOND \ast (GW - DLEV) \] (467)

\[ FLOW = \begin{cases} k_1 \ast |GW| + k_2 \ast (1 - \exp(-k_3 \ast |GW|)), & \text{if } GW \geq 0 \\ -(k_1 \ast |GW| + k_2 \ast (1 - \exp(-k_3 \ast |GW|))), & \text{if } GW < 0 \end{cases} \] (468)

Where SEEP [mm/d] is the exchange with a deeper aquifer (can be negative or positive) and FLOW [mm/d] the exchange with the channel (can be negative or positive). VCOND is a leakage coefficient, DLEV a datum around which the groundwater level can fluctuate, and \( k_1, k_2 \) and \( k_3 \) are runoff coefficients. The channel store aggregates incoming fluxes and produces the total runoff \( Q_t \) [mm/d]:

\[ \frac{dCH}{dt} = SRUN + INT + FLOW - Q \] (469)

\[ SRUN = RUN - TRAP \] (470)

\[ Q_t = \begin{cases} CH, & \text{if } CH > 0 \\ 0, & \text{otherwise} \end{cases} \] (471)
S2.37 HBV-96 (model ID: 37)

The HBV-96 model (fig. S38) was originally developed for use in Sweden, but has been widely applied beyond its original region (Lindström et al., 1997). It can account for different land types (forest, open ground, lakes) but that distinction has been removed here. It has 5 stores and 15 parameters ($TT$, $TTI$, $CFR$, $CFMAX$, $TTM$, $WHC$, $CFLUX$, $FC$, $LP$, $\beta$, $K_0$, $\alpha$, $c$, $K_1$, $MAXBAS$) parameters. The model aims to represent:

- Snow accumulation, melt and refreezing;
- Infiltration and capillary flow to, and evaporation from, soil moisture;
- A non-linear storage-runoff relationship from the upper runoff-generating zone;
- A linear storage-runoff relationship from the lower runoff-generating zone.

S2.37.1 File names

Model: m_37_hbv_15p_5s
Parameter ranges: m_37_hbv_15p_5s_parameter_ranges

S2.37.2 Model equations

\[
\frac{dSP}{dt} = sf + refr - melt \tag{472}
\]

\[
sf = \begin{cases} 
  P, & \text{if } T \leq TT - \frac{1}{2}TTI \\
  P \times \frac{TT + \frac{1}{2}TTI - T}{TTI}, & \text{otherwise} \\
  0, & \text{if } T \geq TT + \frac{1}{2}TTI 
\end{cases} \tag{473}
\]

\[
refr = \begin{cases} 
  CFR \times CFMAX \times (TTM - T), & \text{if } T < TTM \\
  0, & \text{otherwise} 
\end{cases} \tag{474}
\]

\[
melt = \begin{cases} 
  CFMAX \times (T - TTM), & \text{if } T \geq TTM \\
  0, & \text{otherwise} 
\end{cases} \tag{475}
\]

Where $SP$ is the current snow storage [mm]. $sf$ is precipitation that occurs as snowfall [mm/d] based on daily precipitation $P$ [mm/d], threshold temperature for snowfall $TT$ [$^\circ$C] and the snowfall threshold interval length $TTI$ [$^\circ$C]. $refr$ [mm/d] is the refreezing of liquid snow if the current temperature $T$ is below the melting threshold $TTM$ [$^\circ$C], using a coefficient of refreezing $CFR$ [-] and a...
degree-day factor $CF_{MAX}$ [mm/d/$\degree$C]. $melt$ represents snowmelt if the current temperature $T$ is below the melting threshold $TTM$, using the degree-day factor $CF_{MAX}$.

\[
\frac{dWC}{dt} = rf + melt - refr - in - S_{excess}
\]  

(476)

\[
rf = \begin{cases} 
0, & \text{if } T \leq TT - \frac{1}{2}TTI \\
\frac{P * (T - TT) + \frac{1}{2}TTI}{TTI}, & \text{otherwise} \\
P, & \text{if } T \geq TT + \frac{1}{2}TTI 
\end{cases}
\]  

(477)

\[
in = \begin{cases} 
rf + melt, & \text{if } WC \geq WHC * SP \\
0, & \text{otherwise}
\end{cases}
\]  

(478)

\[
S_{c} = \begin{cases} 
WC - WHC * SP, & \text{if } WC \geq WHC * SP \\
0, & \text{otherwise}
\end{cases}
\]  

(479)

Where $WC$ is the current liquid water content in the snow pack [mm], $rf$ is the precipitation occurring as rain [mm/d] based on temperature threshold parameters $TT$ and $TTI$, $refr$ is the refreezing flux, and $in$ is the infiltration to soil moisture [mm/d] that occurs when the water holding capacity of snow gets exceeded. $S_{excess}$ [mm/d] represents excess stored water that is freed when the total possible storage of liquid water in the snow pack is reduced.

\[
\frac{dSM}{dt} = (in + S_{excess}) + cf - E_{a} - r
\]  

(480)

\[
 cf = CFLUX * \left(1 - \frac{SM}{FC}\right)
\]  

(481)

\[
E_{a} = \begin{cases} 
E_{p}, & \text{if } SM \geq LP * FC \\
E_{p} * \frac{SM}{LP * FC}, & \text{otherwise}
\end{cases}
\]  

(482)

\[
r = (in + S_{excess}) * \left(\frac{SM}{FC}\right)^{\beta}
\]  

(483)

Where $SM$ is the current storage in soil moisture [mm], $in$ is the infiltration from the surface, $cf$ the capillary rise [mm/d] from the unsaturated zone, $E_{a}$ evaporation [mm/d] and $r$ the flow to the upper zone [mm/d]. Capillary rise depends on the maximum rate CFLUX [mm/d], scaled by the available storage in soil moisture, expressed as the ration between current storage $SM$ and maximum storage $FC$ [mm]. Evaporation $E_{a}$ occurs at the potential rate $E_{p}$ when current soil moisture is above the wilting point $LP$ [mm], and is scaled linearly below that. Runoff $r$ to the upper zone has a potentially non-linear relationship with infiltration in through parameter $\beta$ [-].
\[
\frac{dUZ}{dt} = r - cf - Q_0 - perc 
\]
(484)

\[
Q_0 = K_0 \ast UZ^{(1+\alpha)} 
\]
(485)

\[
perc = c. 
\]
(486)

Where UZ is the current storage [mm] in the upper zone. Outflow \( Q_0 \) [mm/d] from the reservoir has a non-linear relation with storage through time scale parameter \( K_0 \) \([d^{-1}]\) and \( \alpha \) [-]. Percolation \( perc \) [mm/d] to the lower zone is given as a constant rate \( c \) [mm/d].

\[
\frac{dLZ}{dt} = perc - Q_1 
\]
(487)

\[
Q_1 = K_1 \ast LZ 
\]
(488)

Where LZ is the current storage [mm] in the lower zone. Outflow \( Q_1 \) [mm/d] from the reservoir has a linear relation with storage through time scale parameter \( K_1 \) \([d^{-1}]\). Total outflow is generated by summing \( Q_0 \) and \( Q_1 \) and applying a triangular transform based on lag parameter MAXBAS [d].
S2.38 Tank Model - SMA (model ID: 38)
The Tank Model (fig. S39) is originally developed for use in constantly saturated soils in Japan. This alternative Tank model - SMA (soil moisture accounting) version was developed for regions that are not continuously saturated (Sugawara, 1995). This model is identical to the original tank model, but has an increased depth in the first store to represent primary soil moisture, and adds a new store to represent secondary soil moisture. It has 5 stores and 16 parameters ($s_{m1}$, $s_{m2}$, $k_1$, $k_2$, $A_0$, $A_1$, $A_2$, $t_1$, $t_2$, $B_0$, $B_1$, $t_3$, $C_0$, $C_1$, $t_4$, $D_1$). The model aims to represent:

- Runoff on increasing time scales with depth;
- Soil moisture storage;
- capillary rise to replenish soil moisture.

S2.38.1 File names
Model: m_38_tank2_16p_5s
Parameter ranges: m_38_tank2_16p_5s_parameter_ranges

S2.38.2 Model equations

$$\frac{dS_1}{dt} = P + T_1 - T_2 - E_1 - F_{12} - Y_2 - Y_1 \quad (489)$$

$$T_1 = k_1 \left(1 - \frac{S_1}{s_{m1}}\right), \text{if } S_1 < s_{m1} \quad (490)$$

$$T_2 = k_2 \left(\frac{\min(S_1, s_{m1})}{s_{m1}} - \frac{X_s}{s_{m2}}\right) \quad (491)$$

$$E_1 = \begin{cases} E_p, & \text{if } S_1 > 0 \\ 0, & \text{otherwise} \end{cases} \quad (492)$$

$$F_{12} = \begin{cases} A_0 \ast (S_1 - s_{m1}), & \text{if } S_1 > s_{m1} \\ 0, & \text{otherwise} \end{cases} \quad (493)$$

$$Y_2 = \begin{cases} A_2 \ast (S_1 - t_2), & \text{if } S_1 > t_2 \\ 0, & \text{otherwise} \end{cases} \quad (494)$$

$$Y_1 = \begin{cases} A_1 \ast (S_1 - t_1), & \text{if } S_1 > t_1 \\ 0, & \text{otherwise} \end{cases} \quad (495)$$

Where $S_1$ [mm] is the current storage in the upper zone, refilled by precipitation $P$.
Knoben et al., 2018

[mm/d] and drained by evaporation $E_1$ [mm/d], drainage $F_{12}$ [mm/d] and surface runoff $Y_1$ [mm/d] and $Y_2$ [mm/d]. If $S_1$ is below the soil moisture threshold $sm_1$ [mm], capillary rise $T_1$ [mm/d] from store $S_2$ can occur. Capillary rise has a base rate $k_1$ [mm/d] and decreases linearly as soil moisture $S_1$ nears $sm_1$. This store is connected to the secondary soil moisture store $X_s$ through transfer flux $T_2$ [mm/d]. This flux can work in either direction, based on a base rate $k_2$ [mm/d], the current storages $S_1$ [mm] and $X_s$ [mm] and the maximum soil moisture storages $sm_1$ [mm] and $sm_2$ [mm]. Evaporation $E_1$ occurs at the potential rate $E_p$ [mm/d] if water is available. Drainage to the intermediate layer has a linear relationship with storage through time scale parameter $A_0$ [$d^{-1}$]. Surface runoff $Y_2$ and $Y_1$ occur when $S_1$ is above thresholds $t_2$ [mm] and $t_1$ [mm] respectively. Both are linear relationships through time parameters $A_2$ [$d^{-1}$] and $A_1$ [$d^{-1}$] respectively.

$$\frac{dX_s}{dt} = T_2$$ (496)

Where $X_s$ [mm] is the current storage in the secondary soil moisture zone. This zone has a maximum capacity $sm_2$ [mm], used in the calculation of $T_2$. $T_2$ can be both positive and negative.

$$\frac{dS_2}{dt} = F_{12} - E_2 - T_1 - F_{23} - Y_3$$ (497)

$$E_2 = \begin{cases} E_p, & \text{if } S_1 = 0 \& S_2 > 0 \\ 0, & \text{otherwise} \end{cases}$$ (498)

$$F_{23} = B_0 * S_2$$ (499)

$$Y_3 = \begin{cases} B_1 * (S_2 - t_3), & \text{if } S_2 > t_3 \\ 0, & \text{otherwise} \end{cases}$$ (500)

Where $S_2$ [mm] is the current storage in the intermediate zone, refilled by drainage $F_{12}$ from the upper zone and drained by evaporation $E_2$ [mm/d], drainage $F_{23}$ [mm/d] and intermediate discharge $Y_3$ [mm/d]. $E_2$ occurs at the potential rate $E_p$ if water is available and the upper zone is empty. Drainage to the third layer $F_{23}$ has a linear relationship with storage through time scale parameter $B_0$ [$d^{-1}$]. Intermediate runoff $Y_3$ occurs when $S_2$ is above threshold $t_3$ [mm] and has a linear relationship with storage through time scale parameter $B_1$ [$d^{-1}$].
\[
\frac{dS_3}{dt} = F_{23} - E_3 - F_{34} - Y_4 
\]
(501)

\[
E_3 = \begin{cases} 
E_p, & \text{if } S_1 = 0 \& S_2 = 0 \& S_3 > 0 \\
0, & \text{otherwise} 
\end{cases} 
\]
(502)

\[
F_{34} = C_0 \ast S_3 
\]
(503)

\[
Y_4 = \begin{cases} 
C_1 \ast (S_3 - t_4), & \text{if } S_3 > t_4 \\
0, & \text{otherwise} 
\end{cases} 
\]
(504)

Where \( S_3 \) [mm] is the current storage in the sub-base zone, refilled by drainage \( F_{23} \) from the intermediate zone and drained by evaporation \( E_3 \) [mm/d], drainage \( F_{34} \) [mm/d] and sub-base discharge \( Y_4 \) [mm/d]. \( E_3 \) occurs at the potential rate \( E_p \) if water is available and the upper zones are empty. Drainage to the fourth layer \( F_{34} \) has a linear relationship with storage through time scale parameter \( C_0 \) [d\(^{-1}\)]. Sub-base runoff \( Y_4 \) occurs when \( S_3 \) is above threshold \( t_4 \) [mm] and has a linear relationship with storage through time scale parameter \( C_1 \) [d\(^{-1}\)].

\[
\frac{dS_4}{dt} = F_{34} - E_4 - Y_5 
\]
(505)

\[
E_4 = \begin{cases} 
E_p, & \text{if } S_1 = 0 \& S_2 = 0 \& S_3 = 0 \& S_4 > 0 \\
0, & \text{otherwise} 
\end{cases} 
\]
(506)

\[
Y_5 = D_1 \ast S_4 
\]
(507)

Where \( S_4 \) [mm] is the current storage in the base layer, refilled by drainage \( F_{34} \) from the sub-base zone and drained by evaporation \( E_4 \) [mm/d] and baseflow \( Y_5 \) [mm/d]. \( E_4 \) occurs at the potential rate \( E_p \) if water is available and the upper zones are empty. Baseflow \( Y_5 \) has a linear relationship with storage through time scale parameter \( D_1 \) [d\(^{-1}\)]. Total runoff:

\[
Q_t = Y_1 + Y_2 + Y_3 + Y_4 + Y_5 
\]
(508)
S2.39 Midlands Catchment Runoff Model (model ID: 39)

The Midlands Catchment Runoff model (fig. S40) is intended to be used in a flood-forecasting setting (Moore and Bell, 2001). To reduce the number of free parameters, the original evaporation routines and routing are somewhat simplified here. The model has 5 stores and 16 parameters ($S_{\text{max}}$, $c_{\text{max}}$, $c_0$, $c_1$, $c_e$, $D_{\text{surp}}$, $k_d$, $\gamma_d$, $q_{p,\text{max}}$, $k_g$, $\tau$, $S_{bf}$, $k_{cr}$, $\gamma_{cr}$, $k_{or}$, $\gamma_{or}$). The model aims to represent:

- Interception by vegetation;
- Direct runoff from a variable contributing area;
- A deficit-based approach to soil moisture accounting and interflow and percolation;
- Baseflow from groundwater;
- Uniform flood wave distribution in time;
- In-channel and out-of-channel flood routing.

S2.39.1 File names

Model: m_39_mcrm_16p_5s
Parameter ranges: m_39_mcrm_16p_5s_parameter_ranges

S2.39.2 Model equations

\[
\frac{dS}{dt} = P - E_c - q_t \tag{509}
\]

\[
E_c = \begin{cases} 
E_p, & \text{if } S > 0 \\
0, & \text{otherwise}
\end{cases} \tag{510}
\]

\[
q_t = \begin{cases} 
P, & \text{if } S = S_{\text{max}} \\
0, & \text{otherwise}
\end{cases} \tag{511}
\]

Where $S$ [mm] is the current interception storage, refilled by precipitation $P$ [mm/d] and drained by evaporation $E_c$ [mm/d] and throughfall $q_t$ [mm/d]. $E_c$ occurs at the potential rate whenever possible. $q_t$ occurs only when the store is at maximum capacity $S_{\text{max}}$ [mm].
\[
\frac{dD}{dt} = q_n - E_r - q_d - q_p \tag{512}
\]

\[
q_n = q_t - q_r \tag{513}
\]

\[
q_r = \min \left( c_{\text{max}}, c_0 + c_0 e^{c_1 D} \right) * q_t \tag{514}
\]

\[
E_r = \frac{1}{1 + e^{-c_2 D}} \left( E_p - E_c \right) \tag{515}
\]

\[
q_d = \begin{cases} 
  k_d (D_{\text{surp}} - D)^{\gamma_d}, & \text{if } D > D_{\text{surp}} \\
  0, & \text{otherwise}
\end{cases} \tag{516}
\]

\[
q_p = \begin{cases} 
  q_{p,\text{max}}, & \text{if } D \geq D_{\text{surp}} \\
  \frac{D}{D_{\text{surp}}} q_{p,\text{max}}, & \text{if } 0 < D < D_{\text{surp}} \\
  0, & \text{otherwise}
\end{cases} \tag{517}
\]

Where \( D [\text{mm}] \) is the current storage in soil moisture, refilled by net infiltration \( q_n [\text{mm/d}] \) and drained by evaporation \( E_r [\text{mm/d}] \), direct runoff \( q_d [\text{mm/d}] \) and percolation \( q_p [\text{mm/d}] \). Negative D-values are possible and indicate a moisture deficit. Net inflow \( q_n \) is calculated as the difference between throughfall \( q_t \) and rapid runoff \( q_r [\text{mm/d}] \). \( q_r \) varies depending on the current degree of saturation in the catchment, with a maximum fraction of the catchment area contributing to rapid runoff called \( c_{\text{max}} [-] \), a minimum contributing area of \( c_0 [-] \) and an exponential increase with increasing soil moisture storage, controlled through shape parameter \( c_1 [-] \), in between. \( E_r \) fulfils any remaining evaporation demand but decreases with increasing moisture deficit (negative D values). This relation is controlled through shape parameter \( c_2 \). \( q_d \) has a non-linear relation with storage above a threshold \( D_{\text{surp}} [\text{mm}] \) through time scale parameter \( k_d [d^{-1}] \) and non-linearity parameter \( \gamma_d [-] \). Percolation \( q_p \) has a maximum rate of \( q_{p,\text{max}} \) if \( D \) is above threshold \( D_{\text{surp}} \) and decreases linearly between \( D = D_{\text{surp}} \) and \( D = 0 \).

\[
\frac{dS_g}{dt} = q_p - q_b \tag{518}
\]

\[
q_b = k_b * S_g^{1.5} \tag{519}
\]

Where \( S_g [\text{mm}] \) is the current groundwater storage, refilled by percolation \( q_p \) and drained by baseflow \( q_b [\text{mm/d}] \). \( q_b \) uses time parameter \( k_b [d^{-1}] \) and a fixed non-linearity coefficient of 1.5. Next, \( q_r, q_d \) and \( q_b \) are summed together and distributed uniformly over timespan \( \tau [d] \), giving delayed flow \( u_{ib} [\text{mm/d}] \).
\[
\frac{dS_{ic}}{dt} = u_{ib} - u_{ob} - q_{ic} \tag{520}
\]

\[
u_{ob} = \begin{cases} 
  u_{ib}, & \text{if } S_{ic} = S_{bf} \\
  0, & \text{otherwise} 
\end{cases} \tag{521}
\]

\[
q_{ic} = \begin{cases} 
  k_{cr} \ast S_{ic}^{\gamma_{cr}}, & \text{if } q_{ic} < \frac{3}{4} S_{ic} \\
  \frac{3}{4} S_{ic}, & \text{otherwise} \tag{522}
\end{cases}
\]

Where \( S_{ic} \) [mm] is the current in-channel storage, refilled by \( u_{ic} \) and drained by in-channel flow \( q_{ic} \) [mm/d] and out-of-bank flow \( u_{ob} \) [mm/d]. \( u_{ob} \) only occurs when the store is at maximum capacity \( S_{bf} \) [mm]. \( q_{ic} \) uses time parameter \( k_{cr} \) [d\(^{-1}\)] and non-linearity parameter \( \gamma_{cr} \) [-].

\[
\frac{dS_{oc}}{dt} = u_{ob} - q_{oc} \tag{523}
\]

\[
q_{oc} = \begin{cases} 
  k_{or} \ast S_{oc}^{\gamma_{or}}, & \text{if } q_{oc} < \frac{3}{4} S_{oc} \\
  \frac{3}{4} S_{oc}, & \text{otherwise} \tag{524}
\end{cases}
\]

Where \( S_{oc} \) [mm] is the current out-of-channel storage, refilled by \( u_{ob} \) and drained by out-of-channel flow \( q_{oc} \) [mm/d]. \( q_{oc} \) uses time parameter \( k_{or} \) [d\(^{-1}\)] and non-linearity parameter \( \gamma_{or} \) [-]. Total flow:

\[
Q_t = q_{oc} + q_{ic} \tag{525}
\]
S2.40 SMAR (model ID: 40)

The SMAR model (fig. S41) is the result of a series of modifications to the original 'layers-model' (O’Connell et al., 1970) and summarized by Tan and O’Connor (1996). The model uses an arbitrary number of soil moistures stores connected in series, with each store having a depth of 25mm. The number of stores is an optimization parameter. The current storage in the upper 5 stores features in various equations. For consistency within this framework, the process is reversed: the model uses a fixed number of 5 soil moisture stores, but the depth of each store is variable and given as $S_{n,max} = S_{max}/5$. It has 6 stores and 8 parameters ($H, Y, S_{max}, C, K_G, N, K$). The model aims to represent:

- Saturation excess overland flow;
- Infiltration excess overland flow;
- Gradual infiltration into soil moisture and declining evaporation potential when water is sourced from further underground;
- Groundwater flow;
- Routing of non-groundwater flow.

S2.40.1 File names

Model: m_40_smar_8p_6s
Parameter ranges: m_40_smar_8p_6s_parameter_ranges

S2.40.2 Model equations

\[
\frac{dS_1}{dt} = I - E_1 - q_1
\]  
(526)

\[
I = \begin{cases} 
Y, & \text{if } P^* - R_1 \geq Y \\
Y - P, & \text{otherwise} 
\end{cases}
\]  
(527)

\[
P^* = \begin{cases} 
P - E_p, & \text{if } P > E_p \\
0, & \text{otherwise} 
\end{cases}
\]  
(528)

\[
R_1 = P^* \cdot H \cdot \sum \frac{S_n}{S_{max}}
\]  
(529)

\[
R_! = (P^* - R_1) - I
\]  
(530)

\[
E_1 = C^{(1-1)} \cdot E_p^*
\]  
(531)

\[
E_p^* = \begin{cases} 
E_p - P, & \text{if } E_p > P \\
0, & \text{otherwise} 
\end{cases}
\]  
(532)

\[
q_1 = \begin{cases} 
(P^* - R_1 - R_2, & \text{if } S_1 \geq S_{max} \\
0, & \text{otherwise} 
\end{cases}
\]  
(533)

Figure S41: Structure of the SMAR model
Where $S_1$ [mm] is the current storage in the upper soil layer, $I$ [mm/d] infiltration into the soil, $P^*$ the effective precipitation [mm/d], $R_1$ [mm/d] is direct runoff, $R_2$ [mm/d] is infiltration excess runoff, $E_1$ [mm/d] evaporation and $q_1$ [mm/d] flow towards deeper soil layers. $I$ uses a constant infiltration rate $Y$ [mm/d]. Direct runoff $R_1$ relies on distribution parameter $H$ [-] and is scaled by the current soil moisture storage in all layers compared to the maximum soil moisture storage $S_{\text{max}}$ [mm] of all layers. Evaporation from this soil layer occurs at the effect potential rate $E^*_p$. Runoff to deeper layers $q_1$ only occurs when the current storage exceeds the store’s maximum capacity.

$$S_2 = q_1 - E_2 - q_2$$ (534)

$$E_2 = \begin{cases} C(2^{1-1}) \cdot E_p, & \text{if } S_1 = 0 \\ 0, & \text{otherwise} \end{cases}$$ (535)

$$q_2 = \begin{cases} q_1, & \text{if } S_2 \geq \frac{S_{\text{max}}}{5} \\ 0, & \text{otherwise} \end{cases}$$ (536)

Where $S_2$ [mm] is the current storage in the second soil layer, $E_2$ [mm/d] the evaporation scaled by parameter $C$ [-], and $q_2$ [mm/d] overflow into the next layer. Evaporation is assumed to occur only when the storage in the upper layers has been exhausted.

$$S_3 = q_2 - E_3 - q_3$$ (537)

$$E_3 = \begin{cases} C(3^{1-1}) \cdot E_p, & \text{if } S_2 = 0 \\ 0, & \text{otherwise} \end{cases}$$ (538)

$$q_3 = \begin{cases} q_2, & \text{if } S_3 \geq \frac{S_{\text{max}}}{5} \\ 0, & \text{otherwise} \end{cases}$$ (539)

Where $S_3$ [mm] is the current storage in the second soil layer, $E_3$ [mm/d] the evaporation scaled by parameter $C^2$ [-], and $q_3$ [mm/d] overflow into the next layer. Evaporation is assumed to occur only when the storage in the upper layers has been exhausted.
\[ S_4 = q_3 - E_4 - q_4 \]  
\[ E_4 = \begin{cases} 
C(4-1) \times E_p, & \text{if } S_4 = 0 \\
0, & \text{otherwise} 
\end{cases} \]  
\[ q_4 = \begin{cases} 
q_3, & \text{if } S_4 \geq \frac{S_{\text{max}}}{4} \\
0, & \text{otherwise} \end{cases} \]

Where \( S_4 \) [mm] is the current storage in the second soil layer, \( E_4 \) [mm/d] the evaporation scaled by parameter \( C^3 \) [\(-\)], and \( q_4 \) [mm/d] overflow into the next layer. Evaporation is assumed to occur only when the storage in the upper layers has been exhausted.

\[ S_5 = q_4 - E_5 - R_3 \]  
\[ E_5 = \begin{cases} 
C(5-1) \times E_p, & \text{if } S_4 = 0 \\
0, & \text{otherwise} 
\end{cases} \]  
\[ R_3 = \begin{cases} 
q_4, & \text{if } S_5 \geq \frac{S_{\text{max}}}{5} \\
0, & \text{otherwise} \end{cases} \]

Where \( S_5 \) [mm] is the current storage in the second soil layer, \( E_5 \) [mm/d] the evaporation scaled by parameter \( C^4 \) [\(-\)], and \( R_3 \) [mm/d] overflow towards groundwater. Evaporation is assumed to occur only when the storage in the upper layers has been exhausted.

\[ \frac{dG_w}{dt} = R_g - Q_g \]  
\[ R_g = G_3 \times R_3 \]  
\[ Q_g = K_3 \times G_3 \]

Where \( G_w \) [mm] is the current groundwater storage, refilled by fraction \( G \) [\(-\)] of \( R_3 \) [mm/d] and drained as a linear reservoir with time parameter \( K_3 \) [d\(^{-1}\)]. This groundwater flow \( Q_g \) [mm/d] contributes directly to simulated streamflow \( Q \). The fraction \( R_3^* = (1 - G) \times R_3 \) that does not reach the groundwater reservoir is combined with \( R_1 \) and \( R_2 \) and routed with a gamma function with parameters \( N \) and \( K \). The routing function approximates a Nash-cascade consisting of \( N \) reservoirs with storage coefficient \( K \):

\[ h = \frac{1}{K\Gamma(N)} \left( \frac{t}{K} \right)^{N-1} e^{-t/K} \]
Integration over the time step length $d$ provides the fraction of flow routed per time step $Q_r \ [mm/d]$. Total flow:

$$Q_t = Q_r + Q_g \quad (550)$$
S2.41 NAM model (model ID: 41)

The NAM model (fig. S42) is originally developed for use in Denmark (Nielsen and Hansen, 1973). Here a small modification is made by replacing runoff routing equations of the form $\frac{1}{k} e^{-t/k}$ with the linear reservoirs these equations represent. The model has 6 stores and 10 parameters ($C_s, C_{of}, L^*, C_{L1}, U^*, C_{of}, C_{L2}, K_0, K_1, K_b$). The model aims to represent:

- Snow accumulation and melt;
- Interflow when total soil moisture exceeds a threshold;
- Separation of saturation excess flow into overland flow and infiltration;
- Baseflow from groundwater;

S2.41.1 File names

Model: m_41_nam_10p_6s
Parameter ranges: m_41_nam_10p_6s_parameter_ranges

S2.41.2 Model equations

\[ \frac{dS}{dt} = P_s - M \] (551)
\[ P_s = \begin{cases} P, & \text{if } T \leq 0 \\ 0, & \text{otherwise} \end{cases} \] (552)
\[ M = \begin{cases} C_s * T, & \text{if } T > 0 \\ 0, & \text{otherwise} \end{cases} \] (553)

Where $S$ is the current snow storage [mm], $P_s$ [mm/d] the precipitation that falls as snow and $M$ the snowmelt [mm/d] based on a degree-day factor ($c_s$, [mm/°C/d]). The freezing point of 0° [C] is used as a threshold for snowfall and melt.

Figure S42: Structure of the NAM model
\[
\frac{dU}{dt} = P_r + M - E_U - If - P_n \quad (554)
\]

\[
P_r = \begin{cases} 
P, & \text{if } T > 0 \\
0, & \text{otherwise}
\end{cases} \quad (555)
\]

\[
E_U = \begin{cases} 
E_p, & \text{if } U > 0 \\
0, & \text{otherwise}
\end{cases} \quad (556)
\]

\[
If = \begin{cases} 
C_{if} \frac{L/L^* - C_{L2}}{1 - C_{L2}} U, & \text{if } L/L^* > C_{L1} \\
0, & \text{otherwise}
\end{cases} \quad (557)
\]

\[
P_n = \begin{cases} 
(P_r + M), & \text{if } U = U^* \\
0, & \text{otherwise}
\end{cases} \quad (558)
\]

Where \( U \) [mm] is the current storage in the upper zone, refilled by precipitation as rain \( P_r \) [mm/d] and snowmelt \( M \), and drained by evaporation \( E_U \) [mm/d], interflow \( If \) [mm/d] and net precipitation \( P_n \) [mm/d]. \( P_r \) occurs only when the current temperature exceeds the threshold of \( 0^\circ C \). \( E_U \) occurs at the potential rate \( E_p \) whenever possible. \( If \) occurs only if the fractional storage in the lower zone \( L/L^* \) (\( L \) is current lower zone storage, \( L^* \) is lower zone maximum storage) exceeds a threshold \( C_{L1} \). \( P_n \) occurs only when the upper zone exceeds its maximum storage capacity \( U^* \) [mm].

\[
\frac{dL}{dt} = Dl - E_t \quad (559)
\]

\[
Dl = (P_n - Of) \left(1 - \frac{L}{L^*}\right) \quad (560)
\]

\[
Of = \begin{cases} 
C_{of} \frac{L/L^* - C_{L2}}{1 - C_{L2}} P_n, & \text{if } L/L^* > C_{L2} \\
0, & \text{otherwise}
\end{cases} \quad (561)
\]

\[
E_t = \begin{cases} 
\frac{L}{L^*} E_p, & \text{if } U = 0 \\
0, & \text{otherwise}
\end{cases} \quad (562)
\]

Where \( L \) [mm] is the current storage in the lower zone, refilled by a fraction of infiltration \( Dl \) [mm/d] and drained by evaporation \( E_t \) [mm/d]. \( Dl \) is calculated as a fraction of infiltration \( P_n - Of \), dependent on the current deficit in the lower zone. Note that with the current formulation \( Dl \) might be larger than the lower zone deficit \( L^* - L \) and a constraint of the form \( Dl \leq L^* - L \) is needed. Overland flow \( Of \) [mm/d] is a fraction of \( P_n \) determined using the relative storage in the lower zone \( L/L^* \) and two coefficients \( C_{of} \) [-] and \( C_{L2} \) [-]. \( E_t \) occurs only when the upper zone is empty, and at a reduced rate that uses the relative storage in the lower zone.
\[
\frac{dO}{dt} = Of - Q_o \\
Q_o = K_0 \times O
\] (563)

Where \( O [\text{mm}] \) is the current storage in the overland flow routing store. \( Q_o \) is the routed overland flow, using time coefficient \( K_0 [d^{-1}] \).

\[
\frac{dI}{dt} = If - Q_i \\
Q_i = K_1 \times I
\] (565)

Where \( I [\text{mm}] \) is the current storage in the interflow routing store. \( Q_i \) is the routed interflow, using time coefficient \( K_1 [d^{-1}] \).

\[
\frac{dG}{dt} = Gw - Q_b \\
Gw = (P_n - Of) \left( \frac{L}{L^*} \right) \\
Q_b = K_b \times O
\] (567)

Where \( G [\text{mm}] \) is the current storage in the overland flow routing store, refilled by groundwater flow \( Gw [\text{mm/d}] \). \( Q_b \) is the routed baseflow, using time coefficient \( K_b [d^{-1}] \). Total flow:

\[
Q = Q_o + Q_i + Q_b
\] (570)
S2.42 HYCYMODEL (model ID: 42)

The HYCYMODEL (fig. S43) is originally developed for use in heavily forested catchments in Japan (Fukushima, 1988). The original model specifies evaporation from the $S_b$ store as $E_T = e_p(i) \cdot Q_b/Q_{bc}$, if $S_u < 0 \& S_b < S_{bc}$, with $Q_{bc} = f(S_{bc})$. However, no further details are given and $S_{bc}$ is not listed as a parameter. We assume that $S_{bc} [\text{mm}]$ is a threshold parameter and that evaporation potential declines linearly to zero when the store drops under this threshold. The model has 6 stores and 12 parameters ($C, I_{1,\text{max}}, \alpha, I_{2,\text{max}}, k_{in}, D_{50}, D_{16}, S_{bc}, k_b, p_b, k_h, k_c$). The model aims to represent:

- Split between channel and ground precipitation;
- Interception by canopy and stems/trunks;
- Overland flow from a variable contributing area;
- Non-linear channel flow, hillslope flow and baseflow;
- Channel evaporation.

S2.42.1 File names

Model: m_42_hycymodel_12p_6s
Parameter ranges: m_42_hycymodel_12p_6s_parameter_ranges

S2.42.2 Model equations

\[
\frac{dI_c}{dt} = R_g - E_{ic} - q_{ie} \quad (571)
\]

\[
R_g = (1 - C)P \quad (572)
\]

\[
E_{ic} = \begin{cases} 
(1 - C) \cdot E_p, & \text{if } I_c > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (573)
\]

\[
q_{ie} = \begin{cases} 
R_g, & \text{if } I_c = I_{1,\text{max}} \\
0, & \text{otherwise} 
\end{cases} \quad (574)
\]

Where $I_c [\text{mm}]$ is the current canopy storage, refilled by rainfall on ground $R_g [\text{mm/d}]$ and drained by evaporation $E_{ic} [\text{mm/d}]$ and canopy interception excess $Q_{ie} [\text{mm/d}]$. $R_g$ is the fraction (1-C) [\text{mm}] of rainfall $P [\text{mm/d}]$ that falls on ground (and not in the channel). This fraction appears several times in the model to scale evaporation values according to surface area.

Figure S43: Structure of the HYCYMODEL
$E_{te}$ occurs at the potential rate $E_p$ [mm/d] when possible. $q_{te}$ only occurs when the canopy store is at maximum capacity $I_{1,max}$ [mm].

\[
\frac{dI_s}{dt} = q_{is} - E_{is} - R_s \quad (575)
\]

\[
q_{is} = \alpha \cdot q_{ie} \quad (576)
\]

\[
E_{is} = \begin{cases} 
(1 - C) \cdot E_p, & \text{if } I_s > 0 \\
0, & \text{otherwise}
\end{cases} \quad (577)
\]

\[
R_s = \begin{cases} 
q_{is}, & \text{if } I_s = I_{2,max} \\
0, & \text{otherwise}
\end{cases} \quad (578)
\]

Where $I_s$ [mm] is the current stem and trunk storage, refilled by a fraction of canopy excess $q_{is}$ [mm/d] and drained by evaporation $E_{is}$ [mm/d] and stem flow $R_s$ [mm/d]. $q_{is}$ is the fraction $\alpha$ of canopy excess $q_{te}$. The remainder $(1 - \alpha)$ is throughfall $R_t$ [mm/d]. $E_{is}$ occurs at the potential rate $E_p$ when possible. $R_s$ occurs only when the store is at maximum capacity $I_{2,max}$.

\[
\frac{dS_u}{dt} = R_n - R_e - E_{su} - Q_{in} \quad (579)
\]

\[
R_n = R_t + R_s \quad (580)
\]

\[
R_e = m \cdot R_n \quad (581)
\]

\[
m = \int_{-\infty}^{\xi} \frac{1}{\sqrt{2\pi}} \exp \left( -\frac{\xi^2}{2} \right) \, d\xi \quad (582)
\]

\[
\xi = \frac{\log(S_u/D_{50})}{\log(D_{50}/D_{16})} \quad (583)
\]

\[
E_{su} = \begin{cases} 
(1 - C) \cdot E_p, & \text{if } E_{us} > 0 \\
0, & \text{otherwise}
\end{cases} \quad (584)
\]

\[
Q_{in} = k_{in} \cdot S_u \quad (585)
\]

Where $S_u$ [mm] is the current storage in the upper zone, refilled by net precipitation $R_n$ [mm/d] and drained by effective rainfall $R_e$ [mm/d], evaporation $E_{su}$ [mm/d] and infiltration $Q_{in}$ [mm/d]. $R_n$ is the sum of throughfall $R_t$ and stem flow $R_s$. $R_e$ is a fraction $m$ of $R_e$, determined from a variable contributing area concept. $m$ is calculated is an integral from a regular normal distribution, scaled by the current storage $S_u$ compared to two parameters $D_{50}$ [mm] and $D_{16}$ [mm]. These parameters represent the effective soil depths at which respectively 50% and 16% of the catchment area contribute to $R_e$. $E_{su}$ occurs at the potential rate $E_p$ when possible. $Q_{in}$ has a linear relation with storage through time parameter $k_{in}$ [d$^{-1}$].
\[
\frac{dS_b}{dt} = Q_{in} - E_{sb} - Q_b
\]  
(586)

\[
E_{sb} = \begin{cases} 
(1 - C) * E_p, & \text{if } S_u = 0 \& S_b \geq S_{bc} \\
(1 - C) * E_p \frac{S_b}{S_{bc}}, & \text{otherwise} 
\end{cases}
\]  
(587)

\[
Q_b = k_b * S_p^b
\]  
(588)

Where \(S_b [\text{mm}]\) is the current storage in the lower zone, refilled by infiltration \(Q_{in} [\text{mm/d}]\) and drained by evaporation \(E_{sb} [\text{mm/d}]\) and baseflow \(Q_b [\text{mm/d}]\). \(E_{sb}\) occurs at the potential rate when the store is above a threshold \(S_{bc}[\text{mm}]\), and declines linearly below that. \(Q_b\) has a potentially non-linear relation with storage through time parameter \(k_b [\text{d}^{-1}]\) and scale parameter \(p_b [-]\).

\[
\frac{dS_h}{dt} = R_e - Q_h
\]  
(589)

\[
Q_h = k_h * S_p^h
\]  
(590)

Where \(S_h [\text{mm}]\) is the current storage in the hillslope routing store, refilled by effective rainfall \(R_e\) and drained by hillslope runoff \(Q_h\). \(Q_h\) has a potentially non-linear relation with storage through time parameter \(k_h [\text{d}^{-1}]\) and scale parameter \(p_h [-]\). \(p_h\) is a fixed parameter in the original model with value 5/3.

\[
\frac{dS_c}{dt} = R_c - Q_c
\]  
(591)

\[
Q_c = k_c * S_p^c
\]  
(592)

Where \(S_c [\text{mm}]\) is the current storage in the channel routing store, refilled by rainfall on the channel \(R_c\) and drained by channel runoff \(Q_c\). \(Q_c\) has a potentially non-linear relation with storage through time parameter \(k_c [\text{d}^{-1}]\) and scale parameter \(p_c [-]\). \(p_c\) is a fixed parameter in the original model with value 5/3.

\[
Q_t = Q_c + Q_h + Q_b - E_c
\]  
(593)

\[
E_c = C * E_p
\]  
(594)

Where \(Q_t [\text{mm/d}]\) is the total flow as sum of the three individual flow fluxes minus channel evaporation \(E_c [\text{mm/d}]\).
S2.43 GSM-SOCONT model (model ID: 43)

The Glacier and SnowMelt - SOil CONTRibution model (GSM-SOCONT) model (fig. S44) is a model developed for alpine, partly glaciated catchments (Schaefli et al., 2005). For consistency with other models in this framework, several simplifications are used. The model does not use different elevation bands nor DEM data to estimate certain parameters, and does not calculate an annual glacier mass balance. The model has 6 stores and 12 parameters ($f_{ice}$, $T_0$, $a_{snow}$, $T_m$, $k_s$, $a_{ice}$, $k_i$, A, x, y, $k_{sl}$, $\beta$). The model aims to represent:

- Separate treatment of glacier and non-glacier catchment area;
- Snow accumulation and melt;
- Glacier melt;
- Soil moisture accounting in the non-glacier catchment area.

S2.43.1 File names

Model: m_43_gsmsocont_12p_6s
Parameter ranges: m_43_gsmsocont_12p_6s_parameter_ranges

S2.43.2 Model equations

\[
\frac{dH_{i,s}}{dt} = P_{ice,s} - M_{i,s} \quad (595)
\]

\[
P_{ice,s} = \begin{cases} 
    P_{ce}, & \text{if } T \leq T_0 \\
    0, & \text{otherwise}
\end{cases} \quad (596)
\]

\[
P_{ice} = f_{ice} \times P \quad (597)
\]

\[
M_{i,s} = \begin{cases} 
    a_{snow}(T - T_m), & \text{if } T > T_m \\
    0, & \text{otherwise}
\end{cases} \quad (598)
\]

Where $H_{i,s}$ [mm] is the current storage in the snow pack, refilled by precipitation-as-snow $P_{ce,s}$ [mm/d] and depleted by melt $M_{i,s}$ [mm/d]. $P_{ce,s}$ occurs only when the temperature $T$ [°C] is below a threshold temperature for snowfall $T_0$ [°C]. $P_{ice}$ is the fraction $f_{ice}$ [-] of precipitation $P$ [mm/d] that falls on the ice-covered part of the catchment. $M_{i,s}$ uses a degree-day-factor $a_{snow}$.

Figure S44: Structure of the GSM-SOCONT model
Knoben et al., 2018

\[ [\text{mm/}°\text{C/d}] \] to estimate snow melt if temperature is above a threshold for snow melt \( T_s \ [°\text{C}] \).

\[ \frac{dS_{i,s}}{dt} = M_{i,s} + P_{i,r,s} - Q_{is} \] (599)

\[ P_{i,r,s} = P_{i,\text{ce},r}, \quad \text{if } H_{i,s} > 0 \] (600)

\[ P_{i,\text{ce},r} = \begin{cases} P_{\text{ice}}, & \text{if } T > T_0 \\ 0, & \text{otherwise} \end{cases} \] (601)

\[ Q_{is} = k_s * S_{i,s} \] (602)

Where \( S_{i,s} \ [\text{mm}] \) is the current storage in the snow-water routing reservoir, refilled by snow melt \( M_{i,s} \ [\text{mm/d}] \) and rain-on-snow \( P_{i,\text{ce},s} \ [\text{mm/d}] \), and drained by runoff \( Q_{is} \). 

\[ \frac{dS_{i,i}}{dt} = M_{i,i} + P_{i,r,i} - Q_{ii} \] (603)

\[ P_{i,r,i} = P_{i,\text{ce},r}, \quad \text{if } H_{i,s} = 0 \] (604)

\[ M_{i,s} = \begin{cases} a_{\text{ice}}(T - T_m), & \text{if } T > T_m & \text{& } H_{i,s} = 0 \\ 0, & \text{otherwise} \end{cases} \] (605)

\[ Q_{ii} = k_i * S_{i,i} \] (606)

Where \( S_{i,i} \ [\text{mm}] \) is the current storage in the ice-water routing reservoir, refilled by glacier melt \( M_{i,i} \ [\text{mm/d}] \) and rain-on-ice \( P_{i,\text{ce},i} \ [\text{mm/d}] \), and drained by runoff \( Q_{ii} \). Both \( M_{i,i} \) and \( P_{i,\text{ce},i} \) are assumed to only occur once the snow pack \( H_{i,s} \) is depleted. \( M_{i,i} \) uses a degree-day-factor \( a_{\text{ice}} \ [\text{mm/}°\text{C/d}] \) to estimate glacier melt. Ice storage in the glacier is assumed to be infinite. \( P_{i,\text{ce},r,i} \) is equal to \( P_{i,\text{ce},r} \) if \( H_{i,s} = 0 \). \( Q_{ii} \) has a linear relation with storage through time parameter \( k_i \ [\text{d}^{-1}] \).

\[ \frac{dH_{ni,s}}{dt} = P_{ni,s} - M_{ni,s} \] (607)

\[ P_{ni,s} = \begin{cases} P_{\text{non-ice}}, & \text{if } T \leq T_0 \\ 0, & \text{otherwise} \end{cases} \] (608)

\[ P_{\text{non-ice}} = (1 - f_{\text{ice}}) * P \] (609)

\[ M_{ni,s} = \begin{cases} a_{\text{snow}}(T - T_m), & \text{if } T > T_m \\ 0, & \text{otherwise} \end{cases} \] (610)

Where \( H_{ni,s} \ [\text{mm}] \) is the current snow pack storage on the non-ice covered fraction \( 1 - f_{\text{ice}} \) of the catchment, which increases through snowfall \( F_{ni,s} \ [\text{mm/d}] \) and
decreases through snow melt $M_{ni,s}$ [$mm/d$]. Both fluxes are calculated in the same manner as those on the ice-covered part of the catchment (fluxes $P_{ice,s}$ and $M_{ice,s}$).

$$\frac{dS_{ni,s}}{dt} = P_{inf} - ET - Q_{sl}$$  \hspace{1cm} (611)

$$P_{inf} = P_{eq} - P_{eff}$$  \hspace{1cm} (612)

$$P_{eff} = P_{eq} \left( \frac{S_{ni,s}}{A} \right)^y$$  \hspace{1cm} (613)

$$P_{eq} = M_{ni,s} + P_{ni,r}$$  \hspace{1cm} (614)

$$ET = E_P \left( \frac{S_{ni,s}}{A} \right)^x$$  \hspace{1cm} (615)

$$Q_{sl} = k_{sl}S_{ni,s}$$  \hspace{1cm} (616)

Where $S_{ni,s}$ [$mm$] is the current storage in soil moisture, refilled by infiltrated precipitation $P_{inf}$ [$mm/d$] and drained by evapotranspiration $ET$ [$mm/d$] and slow flow $Q_{sl}$ [$mm/d$]. $P_{inf}$ depends on the effective precipitation $P_{eff}$. $P_{eq}$ is the total of snow melt $M_{ni,s}$ and precipitation-as-rain $P_{ni,r}$ [$mm/d$]. $P_{ni,r}$ is calculated in the same manner as $P_{i,r}$ (equation 7). $ET$ is a fraction potential evapotranspiration $E_P$ [$mm/d$], calculated using $A$ and non-linearity parameter $y$ [-]. $Q_{sl}$ has a linear relation with storage through time parameter $k_{sl}$ [d$^{-1}$].

$$\frac{dS_{ni,q}}{dt} = P_{eff} - Q_{qu}$$  \hspace{1cm} (617)

$$Q_{qu} = \beta S_{ni,q}^{5/3}$$  \hspace{1cm} (618)

Where $S_{ni,q}$ [$mm$] is the current storage in the direct runoff reservoir, refilled by effective precipitation $P_{eff}$ [$mm/d$] and by quick flow $Q_{qu}$ [$mm/d$]. $Q_{sl}$ has a non-linear relation with storage through time parameter $\beta$ [$mm^{5/3}/d$] and the factor 5/3. Total flow:

$$Q = Q_{qu} + Q_{sl} + Q_{is} + Q_{ii}$$  \hspace{1cm} (619)
S2.44 ECHO model (model ID: 44)

The ECHO model (fig. S45) is a single element from the Spatially Explicit Hydrologic Response (SEHR-ECHO) model (Schaefli et al., 2014). Because the model is used as a lumped model here, the "SEHR" prefix was dropped intentionally. For consistency with other models, soil moisture storage $S$ is given here in absolute terms [mm], rather than fractional terms that are used in the original reference. Rain- and snowfall equations are taken from Schaefli et al. (2005). The model has 6 stores and 16 parameters ($\rho, T_s, T_m, a_s, a_f, G_{max}, \theta, \phi, S_{max}, sw, sm, K_{sat}, c, L_{max}, k_f, k_s$). The model aims to represent:

- Interception by vegetation;
- Snowfall, snowmelt, ground-heat flux and storage and refreezing of liquid snow;
- Infiltration, infiltration excess and saturation excess;
- Fast and slow runoff.

S2.44.1 File names

Model: m_44_echo_16p_6s
Parameter ranges: m_44_echo_16p_6s_parameter_ranges

S2.44.2 Model equations

\[
\frac{dI}{dt} = P - E_i - P_n \quad (620)
\]
\[
E_i = \begin{cases} 
E_p, & \text{if } I > 0 \\
0, & \text{otherwise} 
\end{cases} \quad (621)
\]
\[
P_n = \begin{cases} 
P, & \text{if } I = \rho \\
0, & \text{otherwise} \end{cases} \quad (622)
\]

Where $I$ [mm] is the current interception storage, refilled by precipitation $P$ [mm/d] and drained by evaporation $E_i$ [mm/d] and net precipitation $P_n$ [mm/d]. $E_i$ occurs at the potential rate $E_p$ [mm/d] when possible. $P_n$ only occurs when the store is at maximum capacity $\rho$ [mm].

Figure S45: Structure of the ECHO model
\[
\frac{dH_s}{dt} = P_s + F_s - M_s - G_s \quad (623)
\]

\[
P_s = \begin{cases} 
P_{tn}, & \text{if } T \leq T_s \\ 0, & \text{otherwise} \end{cases} \quad (624)
\]

\[
M_s = \begin{cases} 
a_s(T - T_m), & \text{if } T > T_m, H_s > 0 \\ 0, & \text{otherwise} \end{cases} \quad (625)
\]

\[
F_s = \begin{cases} 
a_f a_s(T_m - T), & \text{if } T < T_m, H_w > 0 \\ 0, & \text{otherwise} \end{cases} \quad (626)
\]

\[
G_s = \begin{cases} 
G_{max}, & \text{if } H_s > 0 \\ 0, & \text{otherwise} \end{cases} \quad (627)
\]

Where \( H_s \) [mm] is the current storage in the snow pack, refilled by precipitation-as-snow \( P_s \) [mm/d] and refreezing of melted snow \( F_s \) [mm/d], and drained by snowmelt \( M_s \) [mm/d] and the ground-heat flux \( G_s \) [mm/d]. \( P_s \) is calculated as all effective rainfall after interception, provided the temperature is below a threshold \( T_s \) [°C]. \( M_s \) uses a degree-day factor \( a_s \) [mm/°C/d] and threshold temperature for snowmelt \( T_m \) [°C]. \( F_s \) occurs if the current temperature is below \( T_m \) and the degree-day rate reduced by factor \( a_f \) [-]. \( G_s \) occurs at a constant rate \( G_{max} \) [mm/d].

\[
\frac{dH_w}{dt} = P_r + M_s - F_s - M_w \quad (628)
\]

\[
P_r = \begin{cases} 
P_{tn}, & \text{if } T > T_s \\ 0, & \text{otherwise} \end{cases} \quad (629)
\]

\[
M_w = \begin{cases} 
P_r + M_s, & \text{if } H_w = \theta \cdot H_s \\ 0, & \text{otherwise} \end{cases} \quad (630)
\]

Where \( H_w \) [mm] is the current storage of liquid water in the snow pack, refilled by precipitation-as-rain \( P_r \) [mm/d] and snowmelt \( M_s \) [mm/d], and drained by refreezing \( F_s \) [mm/d] and outflow of melt water \( M_w \) [mm/d]. \( P_r \) is calculated as all effective rainfall after interception, provided the temperature is above a threshold \( T_s \) [°C]. \( M_w \) occurs only if the store is at maximum capacity, which is a fraction \( \theta \) [-] of the current snow pack height \( H_s \) [mm].
\[
\frac{dS}{dt} = F_i - R_D - E_t - L
\]  
(631)

\[
F_i = P_{eq} - R_H
\]  
(632)

\[
P_{eq} = M_w + G_s
\]  
(633)

\[
R_H = \begin{cases} 
\max(P_{eq} - \phi, 0), & \text{if } S < S_{max} \\
0, & \text{otherwise}
\end{cases}
\]  
(634)

\[
R_D = \begin{cases} 
P_{eq}, & \text{if } S = S_{max} \\
0, & \text{otherwise}
\end{cases}
\]  
(635)

\[
E_t = \min \left( \max \left( 0, E_{t,\text{pot}} \frac{S - sw}{sm - sw} \right), E_{t,\text{pot}} \right)
\]  
(636)

\[
E_{t,\text{pot}} = E_p - E_i
\]  
(637)

\[
L = K_{sat} S^c
\]  
(638)

Where \( S \) [mm] is the current storage in the soil moisture zone, refilled by infiltration \( F_i \) [mm/d] and drained by Dunne-type runoff \( R_D \) [mm/d], evapotranspiration \( E_t \) [mm/d] and leakage \( L \) [mm/d]. \( F_i \) is calculated as equivalent precipitation \( P_{eq} \) minus Horton-type runoff \( R_H \). \( P_{eq} \) is the sum of melt water \( M_w \) and the ground-heat flux \( G_s \). \( R_H \) occurs at fixed rate \( \phi \) [mm/d] and only if the soil moisture is not saturated. \( R_D \) is equal to equivalent precipitation \( P_{eq} \) but occurs only when the store is at maximum capacity \( S_{max} \) [mm]. \( E_t \) fulfils any leftover evaporation demand after interception. \( E_t \) occurs at the potential rate until the plant stress point \( sm \) [mm], decreases linearly until the wilting point \( sw \) [mm] and is zero for any lower storage values. \( L \) has a non-linear relationship with storage through time parameter \( K_{sat} \) [d\(^{-1}\)] and coefficient \( c \) [-].

\[
\frac{S_{fast}}{dt} = L_f - R_f
\]  
(639)

\[
L_f = L - L_s
\]  
(640)

\[
L_s = \min(L, L_{max})
\]  
(641)

\[
R_f = k_f \ast S_{fast}
\]  
(642)

Where \( S_{fast} \) [mm] is the current storage in the fast runoff reservoir, refilled by leakage-to-fast-flow \( L_f \) [mm/d] and drained by fast runoff \( R_f \) [mm/d]. \( L_f \) depends on leakage \( L \) from soil moisture and the leakage-to-slow-flow \( L_s \). \( L_s \) is calculated from a maximum leakage rate \( L_{max} \) [mm/d]. \( R_f \) has a linear relation with storage through time parameter \( k_f \) [mm/d].

\[
\frac{dS_{slow}}{dt} = L_s - R_s
\]  
(643)

\[
R_s = k_s \ast S_{slow}
\]  
(644)
Where $S_{slow}$ [mm] is the current storage in the slow runoff reservoir, refilled by leakage-to-slow-flow $L_s$ [mm/d] and drained by slow runoff $R_s$ [mm/d]. $R_s$ has a linear relation with storage through time parameter $k_s$ [mm/d]. Total flow:

$$Q = R_H + R_D + R_F + R_S$$  \hspace{1cm} (645)
S2.45 Precipitation-Runoff Modelling System (PRMS) (model ID: 45)

The PRMS model (fig. S46) is a modelling system that, in its most recent version, allows the user to specify a wide variety of catchment processes and flux equations (Markstrom et al., 2015). The version presented here is a simplified version of the original PRMS model (Leavesley et al., 1983). Simplifications involve the use of PET time series instead of within-model estimates based on temperature, and simpler interception and snow routines. The model has 7 stores and 18 parameters (TT, ddf, α, β, STOR, RETIP, SCN, SCX, REMX, SMAX, cgw, RESMAX, k1, k2, k3, k4, k5, k6). The model aims to represent:

- Snow accumulation and melt;
- Interception by vegetation;
- Depression storage and impervious surface areas;
- Direct runoff based on catchment saturation;
- Infiltration into soil moisture and connection with deeper groundwater;
- Potentially non-linear interflow, baseflow and groundwater sink.

S2.45.1 File names

Model: m_45_prms_18p_7s
Parameter ranges: m_45_prms_18p_7s_parameter_ranges

S2.45.2 Model equations

\[
\frac{dS}{dt} = P_s - M \tag{646}
\]

\[
P_s = \begin{cases} 
P, & \text{if } T \leq TT \\ 
0, & \text{otherwise} \end{cases} \tag{647}
\]

\[
M = \begin{cases} 
ddf \ast (T - TT), & \text{if } T \geq TT \\ 
0, & \text{otherwise} \end{cases} \tag{648}
\]

Where S is the current snow storage [mm], \(P_s\) the rain that falls as snow [mm], M the snowmelt [mm] based on a degree-day factor (ddf, [mm/°C/d]) and threshold temperature for snowfall and snowmelt (TT, [°C]).
\[
\frac{d{XIN}}{dt} = P_{in} - E_{in} - P_{tf} \tag{649}
\]
\[
P_{in} = \alpha \times P_{sm} \tag{650}
\]
\[
P_{sm} = \beta \times P_{r} \tag{651}
\]
\[
P_{r} = \begin{cases} 
  P, & \text{if } T > TT \\
  0, & \text{otherwise}
\end{cases} \tag{652}
\]
\[
E_{in} = \begin{cases} 
  \beta \times E_{p}, & \text{if } XIN > 0 \\
  0, & \text{otherwise}
\end{cases} \tag{653}
\]
\[
P_{tf} = \begin{cases} 
  P_{in}, & \text{if } XIN = STOR \\
  0, & \text{otherwise}
\end{cases} \tag{654}
\]

Where \(XIN\) [mm] is the current storage in the interception reservoir, recharged by intercepted rainfall \(P_{in}\) [mm/d] and drained by evaporation \(E_{i}\) [mm/d] and throughfall \(P_{tf}\) [mm/d]. \(P_{in}\) [mm/d] is the fraction \(\alpha\) of rainfall on non-impervious area \(P_{sm}\) [mm/d] that does not bypass the interception reservoir. \(P_{sm}\) [mm/d] is the fraction \(\beta\) of rainfall \(P_{r}\) [mm/d] that does not fall on impervious area. Rainfall is given as all precipitation \(P\) [mm/d] that occurs when temperature \(T\) [°C] is above a threshold \(TT\) [°C]. \(E_{i}\) [mm/d] occurs at the potential rate \(E_{p}\), corrected for the fraction of the catchment where interception can occur. Throughfall \(P_{tf}\) is all rainfall that reaches the interception reservoir when it is at maximum capacity \(STOR\) [mm].

\[
\frac{d{RSTOR}}{dt} = P_{im} + M_{im} - E_{im} - SAS \tag{655}
\]
\[
P_{im} = (1 - \beta) \times P_{r} \tag{656}
\]
\[
M_{im} = (1 - \beta) \times M \tag{657}
\]
\[
E_{im} = \begin{cases} 
  (1 - \beta) \times E_{p}, & \text{if } RSTOR > 0 \\
  0, & \text{otherwise}
\end{cases} \tag{658}
\]
\[
SAS = \begin{cases} 
  P_{im} + M_{im}, & \text{if } RSTOR = RETIP \\
  0, & \text{otherwise}
\end{cases} \tag{659}
\]

Where \(RSTOR\) [mm] is current depression storage, refilled by rainfall and snowmelt on impervious area, \(P_{im}\) [mm/d] and \(M_{im}\) [mm/d] respectively, and drained by evaporation \(E_{im}\) [mm/d] and surface runoff \(SAS\) [mm/d]. \(P_{im}\) is given as the fraction \(1 - \beta\) of rainfall \(P_{r}\). \(M_{im}\) is given as the fraction \(1 - \beta\) of snowmelt \(M\). \(E_{im}\) occurs at the potential rate \(E_{p}\), corrected for the fraction of the catchment where impervious areas can occur. \(SAS\) occurs when the depression store is at maximum capacity \(RETIP\) [mm].
\[
\frac{dRECHR}{dt} = INF - E_a - PC
\]

\[
INF = M_{sm} + P_{tf} + P_{by} - SRO
\]

\[
M_{sm} = \beta \cdot M
\]

\[
P_{by} = (1 - \alpha) \cdot P_{sm}
\]

\[
SRO = [SCN + (SCX - SCN) \cdot \frac{RECHR}{REMX}] \cdot (M_{sm} + P_{tf} + P_{by})
\]

\[
E_a = \frac{RECHR}{REMX} \cdot (E_p - E_i - E_{im})
\]

\[
PC = \begin{cases} INF, & \text{if } RECHR = REMX \\ 0, & \text{otherwise} \end{cases}
\]

Where \(RECHR \text{ [mm]}\) is the current storage in the upper soil moisture zone, recharged by infiltration \(INF \text{ [mm/d]}\) and drained by evaporation \(E_a \text{ [mm/d]}\) and percolation \(PC \text{ [mm/d]}\). \(INF\) is the difference between incoming snowmelt \(M_{sm} \text{ [mm/d]}\), throughfall \(P_{tf} \text{ [mm/d]}\) and interception bypass \(P_{by} \text{ [mm/d]}\), and surface runoff from saturated area \(SRO \text{ [mm/d]}\). \(S_{sm}\) is snowmelt from the fraction \(\beta \text{ [-]}\) of the catchment that is not impervious. \(P_{by}\) is the fraction \(1 - \alpha\) of rainfall over non-impervious area \(P_{sm}\) that bypasses the interception store. \(SRO\) has a linear relation between minimum contributing area \(SCN \text{ [-]}\) and maximum contributing area \(SCX \text{ [-]}\) based on current storage \(RECHR\) and maximum storage \(REMX \text{ [mm]}\). \(E_a\) uses a similar linear relationship and accounts for already fulfilled evaporation demand by interception and impervious areas. \(PC\) occurs when the store reaches maximum capacity.

\[
\frac{dSMAV}{dt} = PC - E_t - EXCS
\]

\[
E_t = \begin{cases} SMAV \cdot (E_p - E_{in} - E_{im} - E_a), & \text{if } RECHR < (E_p - E_{in} - E_{im}) \\ 0, & \text{otherwise} \end{cases}
\]

\[
EXCS = \begin{cases} PC, & \text{if } SMAV = SMAX - REMX \\ 0, & \text{otherwise} \end{cases}
\]

Where \(SMAV \text{ [mm]}\) is the current storage in the lower soil moisture zone, recharged by percolation from the upper zone \(PC \text{ [mm/d]}\) and drained by transpiration \(E_t \text{ [mm/d]}\) and soil moisture excess \(EXCS \text{ [mm/d]}\). \(E_t\) is corrected for already fulfilled evaporation demand and only occurs if the upper zone can not satisfy this demand. \(E_t\) uses a linear relationship between current storage and the maximum storage in the lower zone \(SMAX - REMX \text{ [mm]}\). \(EXCS\) only occurs when the store has reached maximum capacity \(SMAX - REMX\).
\[
\frac{dRES}{dt} = QRES - GAD - RAS \quad (670)
\]
\[
QRES = \min(EXCS - SEP, 0) \quad (671)
\]
\[
GAD = k_1 \left( \frac{RES}{RESMAX} \right)^{k_2} \quad (672)
\]
\[
RAS = k_3 \ast RES + k_4 \ast RES^2 \quad (673)
\]
\[
(674)
\]

Where \( RES \) [mm] is the current storage in the runoff reservoir, filled by the difference between soil moisture excess \( EXCS \) [mm/d] and constant groundwater recharge \( SEP \) [mm/d], and drained by groundwater drainage \( GAD \) [mm/d] and interflow component \( RAS \) [mm/d]. \( GAD \) is potentially non-linear using time coefficient \( k_1 \) [d\(^{-1}\)] and non-linearity coefficient \( k_2 \) [-], and is also scaled by the maximum reservoir capacity \( RESMAX \) [mm]. \( RAS \) is non-linear interflow based on coefficients \( k_3 \) [d\(^{-1}\)] and \( k_4 \) [mm\(^{-1}\)d\(^{-1}\)].

\[
\frac{dGW}{dt} = SEP + GAD - BAS - SNK \quad (675)
\]
\[
SEP = \min(c_{gw}, EXCS) \quad (676)
\]
\[
BAS = k_5 \ast GW \quad (677)
\]
\[
SNK = k_6 \ast GW \quad (678)
\]

Where \( GW \) [mm] is the current groundwater storage, refilled by groundwater recharge from soil moisture \( SEP \) and recharge from runoff reservoir \( GAD \) and drained by baseflow \( BAS \) [mm/d] and flow to deeper groundwater \( SNK \) [mm/d]. \( SEP \) occurs at the maximum rate \( c_{gw} \) [mm/d] if possible. \( BAS \) is a linear reservoir with time coefficient \( k_5 \) [d\(^{-1}\)]. \( SNK \) is a linear reservoir with time coefficient \( k_6 \) [d\(^{-1}\)]. Total flow \( Q_t \) [mm/d]:

\[
Q_t = SAS + SRO + RAS + BAS \quad (679)
\]
S2.46 Climate and Land-use Scenario Simulation in Catchments model (model ID: 46)

The CLASSIC model (fig. S47) is developed as a modular semi-distributed grid-based rainfall runoff model (Crooks and Naden, 2007). For comparability with other models the grid-based routing component is not included here, nor is the arable soil element because input data for this soil type is not supported. The model represents runoff from three different soil categories: permeable, semi-permeable and impermeable. It has 8 stores and 12 parameters ($f_{ap}$, $f_{dp}$, $d_p$, $c_q$, $d_1$, $f_{as}$, $f_{ds}$, $d_s$, $c_{xq}$, $c_{xs}$, $c_u$). The model aims to represent:

- Division into permeable, semi-permeable and impermeable areas;
- Infiltration into permeable soils and deficit-based soil moisture accounting;
- Infiltration into semi-permeable soils and direct runoff from semi-permeable soils (bypassing the moisture accounting);
- Fixed interception on impermeable soils;
- Linear flow routing from permeable soils;
- Fast and slow routing from semi-permeable soils;
- Linear flow routing from impermeable soils.

S2.46.1 File names

Model: m_46_classic_12p_8s
Parameter ranges: m_46_classic_12p_8s_parameter_ranges

S2.46.2 Model equations

\[
\frac{dP_x}{dt} = P_p - E_{px} = P_{px} \quad (680)
\]

\[
P_p = f_{ap} \cdot P \quad (681)
\]

\[
E_{px} = \begin{cases} 
  f_{ap} \cdot E_p, & \text{if } P_x > 0 \\
  0, & \text{otherwise}
\end{cases} \quad (682)
\]

\[
P_{px} = \begin{cases} 
  P_p, & \text{if } P_x = f_{dp} \cdot d_p \\
  0, & \text{otherwise}
\end{cases} \quad (683)
\]

Where $P_x$ [mm] is the current storage in the upper permeable layer, refilled by precipitation $P_p$ [mm/d] and drained by evaporation $E_{px}$ [mm/d] and...
excess flow $P_{px}$ [mm/d]. $P_x$ is the fraction of precipitation $P$ [mm/d] that falls on permeable area $f_{ap}$ [-]. $E_{px}$ occurs at the potential rate $E_p$ [mm/d] whenever possible, adjusted for the fraction of area that is permeable soil. $P_{px}$ only occurs when the store is at maximum capacity $f_{dp}*d_p$, where $d_p$ is the total soil depth (sum of depths X and Y) in the permeable area and $f_{dp}$ the fraction of this depth that is store X.

\[
\begin{align*}
\frac{dP_y}{dt} &= -P_{px} + E_{py} + P_{pe} \\
E_{py} &= 1.9 * \exp \left[ -0.6523 * \frac{(P_y + f_{dp} * d_p)}{f_{dp} * d_p} \right] * (f_{ap} * E_p - E_{px}) \\
P_{pe} &= \begin{cases} P_{px}, & \text{if } P_y = 0 \\ 0, & \text{otherwise} \end{cases}
\end{align*}
\] (684)

Where $P_y$ [mm] is the current deficit, which is increased by evaporation $E_{py}$ [mm/d] and decreased by inflow $P_{px}$ [mm/d]. Effective precipitation $P_{pe}$ [mm/d] is only generated when the deficit is 0. $E_{py}$ decreases exponentially with increasing deficit.

\[
\begin{align*}
\frac{dP}{dt} &= P_{pe} - q \\
qu &= c_p * P
\end{align*}
\] (685) (686)

Where $P$ [mm] is the current storage in the permeable soil routing store, refilled by effective rainfall on permeable soil $P_{pe}$ [mm/d] and drained by baseflow $q$ [mm/d]. $q$ has a linear relation with storage through time scale parameter $c_p$ [d$^{-1}$].

\[
\begin{align*}
\frac{dS_x}{dt} &= P_{si} - E_{sz} - P_{sx} \\
P_{si} &= d_1 * P_s \\
P_s &= f_{as} * P \\
E_{sz} &= \begin{cases} f_{as} * E_p, & \text{if } S_x > 0 \\ 0, & \text{otherwise} \end{cases} \\
P_{sx} &= \begin{cases} P_s, & \text{if } S_x = f_{ds} * d_s \\ 0, & \text{otherwise} \end{cases}
\end{align*}
\] (687) (688) (689) (690) (691) (692) (693)
Where $S_x \text{[mm]}$ is the current storage in the upper semi-permeable layer, refilled by infiltration $P_{sx} \text{[mm/d]}$ and drained by evaporation $E_{sx} \text{[mm/d]}$ and excess flow $P_{sx} \text{[mm/d]}$. $P_s$ is the fraction $d_1 \text{[mm]}$ of precipitation on semi-permeable area that infiltrates into the soil. The complementary fraction $1 - d_1$ of $P_s$ bypasses the soil and directly becomes effective rainfall as $P_{sd}$. $P_s$ is the fraction of precipitation $P \text{[mm/d]}$ that falls on semi-permeable area $f_{as} \text{[-]}$. $E_{sx}$ occurs at the potential rate $E_p \text{[mm/d]}$ whenever possible, adjusted for the fraction of area that is semi-permeable soil. $P_{sx}$ only occurs when the store is at maximum capacity $f_{ds} \times d_s$, where $d_s$ is the total soil depth (sum of depths X and Y) in the semi-permeable area and $f_{ds}$ the fraction of this depth that is store X.

\[
\frac{dS_y}{dt} = -P_{sx} + E_{sy} + P_{se} \tag{694}
\]

\[
E_{sy} = 1.9 \times \exp \left[ -0.6523 \times \frac{(S_y + f_{ds} \times d_s)}{f_{ds} \times d_s} \right] \times (f_{as} \times E_p - E_{sx}) \tag{695}
\]

\[
P_{pc} = \begin{cases} P_{sx}, & \text{if } S_y = 0 \\ 0, & \text{otherwise} \end{cases} \tag{696}
\]

Where $S_y \text{[mm]}$ is the current deficit, which is increased by evaporation $E_{sy} \text{[mm/d]}$ and decreased by inflow $P_{sx} \text{[mm/d]}$. Effective precipitation $P_{se} \text{[mm/d]}$ is only generated when the deficit is 0. $E_{sy}$ decreases exponentially with increasing deficit.

\[
\frac{dS_q}{dt} = P_{sq} - x_q \tag{697}
\]

\[
P_{sq} = (1 - d_2) \times (P_{se} + P_{sd}) \tag{698}
\]

\[
x_q = c_{xq} \times S_q \tag{699}
\]

Where $S_q \text{[mm]}$ is the current storage in the semi-permeable quick soil routing store, refilled by a fraction of effective rainfall on semi-permeable soil $P_{sq} \text{[mm/d]}$ and drained by quick flow $x_q \text{[mm/d]}$. $P_{sq}$ is the fraction $d_2 \text{[-]}$ of $(P_{se} + P_{sd})$ that is quick flow. $x_q$ has a linear relation with storage through time scale parameter $c_{xq} \text{[d}^{-1} \text{]}$.

\[
\frac{dS_s}{dt} = P_{ss} - x_s \tag{700}
\]

\[
P_{ss} = (1 - d_2) \times (P_{se} + P_{sd}) \tag{701}
\]

\[
x_s = c_{xs} \times S_s \tag{702}
\]

Where $S_s \text{[mm]}$ is the current storage in the semi-permeable quick soil routing store, refilled by a fraction of effective rainfall on semi-permeable soil $P_{ss} \text{[mm/d]}$ and drained by slow flow $x_s \text{[mm/d]}$. $P_{ss}$ is the fraction $1 - d_2 \text{[-]}$ of $(P_{se} + P_{sd})$ that is slow flow. $x_s$ has a linear relation with storage through time scale parameter $c_{xs} \text{[d}^{-1} \text{]}$. 

120
\[
\frac{dI}{dt} = P_{ie} - u \tag{703}
\]
\[
P_{ie} = P_i - E_i \tag{704}
\]
\[
P_i = P - P_p - P_s \tag{705}
\]
\[
u = c_u * I \tag{706}
\]

Where \(I\) [mm] is the current storage in the impermeable soil routing store, refilled by effective rainfall on impermeable soil \(P_{ie}\) [mm/d] and drained by baseflow \(u\) [mm/d]. \(P_{ie}\) is the remained of precipitation on impermeable soils \(P_i\) [mm/d], after a constant evaporation \(E_i\) has been extracted. \(E_i\) is fixed at 0.5 [mm/d]. \(xs\) has a linear relation with storage through time scale parameter \(c_{xs} [d^{-1}]\). Total flow:

\[
Q = q + xs + xq + u \tag{707}
\]
S3 Flux equations

Section S2 gives descriptions of each model and provides both Ordinary Differential Equations and the constitutive functions that describe each model’s fluxes. These constitutive functions and any relevant constraints are implemented in MARRMoT as individual flux files. Each flux file contains computer code that combines the constitutive function and constraints (if needed). Flux files are located in the folder "/MARRMoT/Models/Flux files/". The User Manual contains details on understanding, modifying and creating new flux files. Table S1 shows a complete overview of fluxes currently implemented in MARRMoT.

Table S1: Equations from model descriptions and their implementation in MARRMoT (Table starts on following page)
<table>
<thead>
<tr>
<th>Process</th>
<th>Details</th>
<th>Function name</th>
<th>Constitutive function</th>
<th>Constraints</th>
<th>MARRMoT Code</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Abstraction</td>
<td>Groundwater abstraction at a constant rate</td>
<td>abstraction_1</td>
<td>$\text{flux}_\text{out} = \theta_1$</td>
<td>None, taken from a store with possible negative depth</td>
<td>$\text{flux}_\text{out} = \theta_1$</td>
<td>25</td>
</tr>
<tr>
<td>Baseflow</td>
<td>Linear reservoir</td>
<td>baseflow_1</td>
<td>$\text{flux}_\text{out} = \theta_1 \cdot S$</td>
<td>$\text{flux}_\text{out} = \theta_1 \cdot S$</td>
<td>2, 4, 6, 8, 9, 12, 13, 15, 16, 17, 18, 20, 21, 24, 25, 26, 27, 28, 29, 30, 31, 32, 33, 34, 35, 36, 37, 38, 40, 41, 43, 44, 45, 46</td>
<td></td>
</tr>
<tr>
<td>Non-linear outflow from a reservoir</td>
<td>baseflow_2</td>
<td></td>
<td>$\text{flux}_\text{out} = \left( \frac{1}{\theta_1} \right)^{\frac{1}{\theta_2}} S$</td>
<td>$\text{flux}_\text{out} \leq \frac{S}{\Delta t}$ To prevent complex numbers, $S = [0, \infty&gt;$</td>
<td>$\text{flux}_\text{out} = \min \left( \frac{S}{\Delta t} \left( \frac{1}{\theta_1} \max(0, S) \right)^{\frac{1}{\theta_2}} \right)$</td>
<td>9, 11</td>
</tr>
<tr>
<td>Empirical exponential outflow from a reservoir</td>
<td>baseflow_3</td>
<td></td>
<td>$\text{flux}<em>\text{out} = \frac{S</em>{\max}}{4} S^5$</td>
<td>Empirical equation, so interwoven with other equations that no constraints are needed. Also implicitly assumes time step $\Delta t = 1$</td>
<td>$\text{flux}<em>\text{out} = \frac{S</em>{\max}}{4} S^5$</td>
<td>7</td>
</tr>
<tr>
<td>Exponential outflow from a deficit store</td>
<td>baseflow_4</td>
<td></td>
<td>$\text{flux}_\text{out} = \theta_1 e^{-\theta_2 S}$</td>
<td></td>
<td>$\text{flux}_\text{out} = \theta_1 e^{-\theta_2 S}$</td>
<td>14</td>
</tr>
<tr>
<td>Non-linear outflow scaled by current relative storage</td>
<td>baseflow_5</td>
<td></td>
<td>$\text{flux}<em>\text{out} = \theta_1 \left( \frac{S}{S</em>{\max}} \right)^{\theta_2}$</td>
<td>$\text{flux}_\text{out} \leq \frac{S}{\Delta t}$ To prevent complex numbers, $S = [0, \infty&gt;$</td>
<td>$\text{flux}<em>\text{out} = \min \left( \frac{S}{\Delta t} \theta_1 \left( \frac{\max(0, S)}{S</em>{\max}} \right)^{\theta_2} \right)$</td>
<td>22</td>
</tr>
<tr>
<td>Quadratic outflow from reservoir if a storage threshold is exceeded</td>
<td>baseflow_6</td>
<td></td>
<td>$\text{flux}_\text{out} = \left{ \begin{array}{ll} \theta_1 \cdot S^2, &amp; \text{if } S &gt; \theta_2 \ 0, &amp; \text{otherwise} \end{array} \right.$</td>
<td>$\text{flux}_\text{out} \leq \frac{S}{\Delta t}$</td>
<td>$\text{flux}_\text{out} = \min \left( \theta_1 \cdot S^2 \cdot \frac{S}{\Delta t} \right) \cdot \left[ 1 - \text{logisticSmother}(S, \theta_2) \right]$</td>
<td>25</td>
</tr>
<tr>
<td>Non-linear outflow from a reservoir</td>
<td>baseflow_7</td>
<td></td>
<td>$\text{flux}_\text{out} = \theta_1 S^{\theta_2}$</td>
<td>$\text{flux}_\text{out} \leq \frac{S}{\Delta t}$ To prevent complex numbers, $S = [0, \infty&gt;$</td>
<td>$\text{flux}_\text{out} = \min \left( \frac{S}{\Delta t} \theta_1 \max(0, S)^{\theta_2} \right)$</td>
<td>39, 42</td>
</tr>
</tbody>
</table>

continued …
<table>
<thead>
<tr>
<th>Process Details</th>
<th>Function name</th>
<th>Constitutive function</th>
<th>Constraints</th>
<th>MARRMoT Code</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Exponential scaled outflow from a deficit store</td>
<td>$flux_{out} = \theta_1 \left( \frac{e^{S/S_{max}}}{e^{S_{max}/S_{max}} - 1} \right)$</td>
<td>$S \leq S_{max}$</td>
<td>$f_{flux_{out}} = \theta_1 \left( \frac{e^{\min\left(S, \theta_3 \cdot S_{max}/S_{max}\right)}}{e^{S_{max}/S_{max}} - 1} \right)$</td>
<td>23</td>
<td></td>
</tr>
<tr>
<td>Linear outflow from a reservoir if a storage threshold is exceeded</td>
<td>$flux_{out} = \theta_1 (S - \theta_2)$, if $S &gt; \theta_2$ 0, otherwise</td>
<td></td>
<td>$f_{flux_{out}} = \theta_1 \ast \max(0, S - \theta_2)$</td>
<td>20</td>
<td></td>
</tr>
<tr>
<td>Capillary rise by relative deficit in receiving store</td>
<td>$flux_{out} = \theta_1 \left( 1 - \frac{S_1}{S_{1,max}} \right)$</td>
<td>$flux_{out} \leq \frac{S_1}{\Delta t}$</td>
<td>$f_{flux_{out}} = \min\left( \theta_1 \left( 1 - \frac{S_1}{S_{1,max}} \right), \frac{S_2}{S_{2}} \right)$</td>
<td>37</td>
<td></td>
</tr>
<tr>
<td>Capillary rise at a constant rate</td>
<td>$flux_{out} = \theta_1 \left( \begin{array}{l} \frac{S}{S_{max}} \quad S \geq 0 \ 0 \quad \text{otherwise} \end{array} \right)$</td>
<td>$flux_{out} \leq \frac{S}{\Delta t}$</td>
<td>$f_{flux_{out}} = \min\left( \frac{S}{\Delta t}, \theta_1 \right)$</td>
<td>13, 15</td>
<td></td>
</tr>
<tr>
<td>Capillary rise if the receiving store is below a storage threshold</td>
<td>$flux_{out} = \theta_1 \left( \begin{array}{l} \frac{S}{S_{max}} \quad S_1 &lt; \theta_1 \ 0 \quad \text{otherwise} \end{array} \right)$</td>
<td>$flux_{out} \leq \frac{S}{\Delta t}$</td>
<td>$f_{flux_{out}} = \min\left( \frac{S}{\Delta t}, \theta_1 \left( \frac{S}{S_{max}} \right) \ast \text{logSmoother}(S_1, \theta_2) \right)$</td>
<td>38</td>
<td></td>
</tr>
<tr>
<td>Depression storage</td>
<td>$flux_{out} = \theta_1 \ast \exp\left( \begin{array}{l} -\theta_2 \frac{S}{S_{max}} \quad S_{max} - S \leq \Delta t \ 0 \quad \text{otherwise} \end{array} \right) \ast flux_{in}$</td>
<td>$flux_{out} \leq \frac{S_{max} - S}{\Delta t}$</td>
<td>$f_{flux_{out}} = \min\left( \theta_1 \ast \exp\left( -\theta_2 \frac{S}{S_{max}} \right), \frac{S_{max} - S}{\Delta t} \right) \ast flux_{in}$</td>
<td>36</td>
<td></td>
</tr>
<tr>
<td>Evaporation at the potential rate</td>
<td>$E_a = \begin{array}{l} E_{p} \quad \text{if } S \geq 0 \ 0 \quad \text{otherwise} \end{array}$</td>
<td>$E_a \leq \frac{S}{\Delta t}$</td>
<td>$E_a = \min\left( E_{p}, \frac{S}{\Delta t} \right)$</td>
<td>2, 6, 12, 13, 16, 17, 18, 23, 25, 26, 27, 33, 34, 36, 38, 39, 41, 42, 44, 45, 46</td>
<td></td>
</tr>
<tr>
<td>Evaporation at scaled plant-controlled rate</td>
<td>$E_a = \left( \begin{array}{l} E_{p} \frac{S}{S_{max}} \quad S \geq 0 \ 0 \quad \text{otherwise} \end{array} \right)$</td>
<td>$E_a \leq \frac{E_{p}}{S}$</td>
<td>$E_a = \min\left( \theta_1 \frac{S}{S_{max}}, \frac{E_{p}}{\Delta t} \right)$</td>
<td>18, 36</td>
<td></td>
</tr>
<tr>
<td>Evaporation scaled by relative storage below a wilting point and at the potential rate above wilting point</td>
<td>$E_a = \left( \begin{array}{l} E_{p} \frac{S}{S_{max}} \quad S &lt; \theta_3 \cdot S_{max} \ \theta_3 \cdot S_{max} \quad E_{p} \quad \text{otherwise} \end{array} \right)$</td>
<td>$E_a \leq \frac{E_{p}}{\Delta t}$</td>
<td>$E_a = \min\left( \theta_1 \frac{E_{p}}{S_{max}}, \frac{E_{p}}{\Delta t} \right)$</td>
<td>3, 11, 14, 21, 26, 34, 37, 42</td>
<td></td>
</tr>
<tr>
<td>Scaled evaporation if storage is above the wilting point, constrained by a limitation parameter</td>
<td>$E_a = E_{p} \ast \max\left( 0, \theta_1 \frac{S - \theta_2 \cdot S_{max}}{S_{max} - \theta_2 \cdot S_{max}} \right)$</td>
<td>$E_a \leq \frac{S}{\Delta t}$</td>
<td>$E_a = \min\left( E_{p} \ast \max\left( 0, \theta_1 \frac{S - \theta_2 \cdot S_{max}}{S_{max} - \theta_2 \cdot S_{max}} \right), \frac{S}{\Delta t} \right)$</td>
<td>15</td>
<td></td>
</tr>
<tr>
<td>Evaporation from bare soil, scaled by relative storage</td>
<td>$E_a = (1 - \theta_2) \frac{S}{S_{max}} E_{p}$</td>
<td>$E_a \leq \frac{E_{p}}{\Delta t}$</td>
<td>$E_a = \min\left( (1 - \theta_2) \frac{S}{S_{max}} E_{p}, \frac{S}{\Delta t} \right)$</td>
<td>4, 8, 9, 16</td>
<td></td>
</tr>
<tr>
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</tr>
<tr>
<td>Transpiration from vegetation at the potential rate if soil moisture is above a wilting point and scaled by relative storage if not</td>
<td>evap_6</td>
<td>$E_a$</td>
<td>$\begin{cases} \theta_1 \cdot E_p &amp; \text{if } S &gt; \theta_2 \cdot S_{\text{max}} \ \frac{S}{S_{\text{max}}} \theta_2 \cdot E_p &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S}{\Delta t}$, $E_a \leq \frac{S}{S_{\text{max}}}$</td>
<td>$41, 45$</td>
</tr>
<tr>
<td>Evaporation scaled by relative storage</td>
<td>evap_7</td>
<td>$E_a$</td>
<td>$\frac{S}{S_{\text{max}}} E_p$</td>
<td>$E_a \leq \frac{S}{\Delta t}$, $E_a \leq \frac{S}{S_{\text{max}}}$</td>
<td>$1, 3, 10$, $11, 19, 22, 24, 29, 30, 31, 32, 33, 35, 45$</td>
</tr>
<tr>
<td>Transpiration from vegetation at the potential rate if soil moisture is above a wilting point and linearly decreasing if not. Also scaled by relative storage across all stores</td>
<td>evap_8</td>
<td>$E_a$</td>
<td>$\begin{cases} \frac{S_1}{S_1 + S_2} \theta_1 \cdot E_p &amp; \text{if } S_1 &gt; \theta_2 \ \frac{S_1}{S_1 + S_2} \theta_2 \cdot E_p &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S_1}{\Delta t}$, $E_a \geq 0$</td>
<td>$8$</td>
</tr>
<tr>
<td>Evaporation from bare soil scaled by relative storage and by relative water availability across all stores</td>
<td>evap_9</td>
<td>$E_a$</td>
<td>$\frac{S_1}{S_1 + S_2} \cdot (1 - \theta_1) \cdot \frac{S_1}{S_{\text{max}} - S_2} E_p$</td>
<td>$E_a \leq \frac{S_1}{\Delta t}$, $E_a \geq 0$</td>
<td>$8$</td>
</tr>
<tr>
<td>Evaporation from bare soil, scaled by relative storage</td>
<td>evap_10</td>
<td>$E_a$</td>
<td>$\frac{S_1}{S_{\text{max}}} E_p$</td>
<td>$E_a \leq E_p$, $E_a \leq \frac{S}{\Delta t}$</td>
<td>$8$</td>
</tr>
<tr>
<td>Evaporation quadratically related to current soil moisture</td>
<td>evap_11</td>
<td>$E_a$</td>
<td>$\left( 2 \cdot \frac{S}{S_{\text{max}}} - \left( \frac{S}{S_{\text{max}}} \right)^2 \right) E_p$</td>
<td>$E_a \geq 0$</td>
<td>$7$</td>
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<tr>
<td>Evaporation from deficit store, with exponential decline as deficit goes below a threshold</td>
<td>evap_12</td>
<td>$E_a$</td>
<td>$\min \left( 1, e^{\frac{-S}{S_{\text{max}}}} \right) E_p$</td>
<td>$E_a = \min \left( 1, e^{\frac{-S}{S_{\text{max}}}} \right) E_p$</td>
<td>$5$</td>
</tr>
<tr>
<td>Exponentially scaled evaporation</td>
<td>evap_13</td>
<td>$E_a$</td>
<td>$\theta_1^2 \cdot E_p$</td>
<td>$E_a \leq \frac{S}{\Delta t}$, $E_a \leq \frac{S}{S_{\text{max}}}$</td>
<td>$40$</td>
</tr>
<tr>
<td>Exponentially scaled evaporation that only activates if another store goes below a certain threshold</td>
<td>evap_14</td>
<td>$E_a$</td>
<td>$\begin{cases} \theta_1^2 \cdot E_p &amp; \text{if } S_2 \leq S_{2,\text{min}} \ 0, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S_1}{\Delta t}$, $E_a = \min \left( \theta_1^2 \cdot E_p \cdot \frac{S_1}{\Delta t} \right)$</td>
<td>$40$</td>
</tr>
<tr>
<td>Scaled evaporation if another store is below a threshold</td>
<td>evap_15</td>
<td>$E_a$</td>
<td>$\begin{cases} \frac{S_1}{S_{\text{max}}} E_p &amp; \text{if } S_2 &lt; \theta_1 \ 0, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S_1}{\Delta t}$, $E_a = \min \left( \frac{S_1}{S_{1,\text{max}}} \cdot E_p \cdot \logisticSmooth \left( S_{2,\theta_2}, \frac{S_1}{\Delta t} \right) \right)$</td>
<td>$41, 45$</td>
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### Process Details

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<td>Scaled evaporation if another store is below a threshold</td>
<td>evap_16</td>
<td>$E_a = \begin{cases} \theta_1 E_p, &amp; \text{if } S &lt; \theta_2 \ 0, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S_1}{\Delta t}$</td>
<td>$E_a = \min \left( \theta_1 \cdot E_p \cdot \logisticSmother \left(S(S_2, \theta_2) \cdot \frac{S_1}{\Delta t} \right) \right)$</td>
<td>17, 25</td>
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<tr>
<td>Scaled evaporation from a store that allows negative values</td>
<td>evap_17</td>
<td>$E_a = \frac{1}{1 + e^{-\theta_1 S_2}} E_p$</td>
<td>None, because the store is allowed to go negative</td>
<td>$E_a = \frac{1}{1 + e^{-\theta_1 S_2}} E_p$</td>
<td>39</td>
<td></td>
</tr>
<tr>
<td>Exponentially declining evaporation from deficit store</td>
<td>evap_18</td>
<td>$E_a = \theta_1 e^{-\frac{S}{\theta_2}} E_p$</td>
<td></td>
<td>$E_a = \theta_1 e^{-\frac{S}{\theta_2}} E_p$</td>
<td>46</td>
<td></td>
</tr>
<tr>
<td>Non-linear scaled evaporation</td>
<td>evap_19</td>
<td>$E_a = \theta_1 \left( \frac{S}{S_{\max}} \right)^2 E_p$</td>
<td>$E_a \leq E_p$</td>
<td>$E_a = \min \left( \theta_1 \cdot \max \left( \frac{S}{S_{\max}} \right)^2 E_p, \frac{S}{\Delta t} \right)$</td>
<td>25, 43</td>
<td></td>
</tr>
<tr>
<td>Evaporation limited by a maximum evaporation rate and scaled below a wilting point</td>
<td>evap_20</td>
<td>$E_a = \begin{cases} S \frac{S_{\max}}{E_p}, &amp; \text{if } S &lt; \theta_2 S_{\max} \ \frac{S_{\max}}{E_p} \frac{\theta_2}{\theta_1 S_{\max}}, &amp; \text{if } \theta_2 S_{\max} \leq S \leq \theta_1 \ \frac{S_{\max}}{E_p} \frac{\theta_2}{\theta_1 S_{\max}}, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq E_p$</td>
<td>$E_a = \min \left( \theta_1 \cdot \frac{S}{S_{\max}}, \frac{S}{\Delta t} \right)$</td>
<td>20</td>
<td></td>
</tr>
<tr>
<td>Threshold-based evaporation with constant minimum rate</td>
<td>evap_21</td>
<td>$E_a = \begin{cases} \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{if } S &gt; \theta_1 \ \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{if } \theta_2 \leq S \leq \theta_1 \ \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S}{\Delta t}$</td>
<td>$E_a = \min \left( \max \left( \frac{S}{\theta_1 \cdot \frac{S_{\max}}{E_p}} \right), \frac{E_p}{\Delta t} \right)$</td>
<td>28</td>
<td></td>
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<tr>
<td>Threshold-based evaporation rate</td>
<td>evap_22</td>
<td>$E_a = \begin{cases} \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{if } S &gt; \theta_1 \ \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{if } \theta_2 \leq S \leq \theta_1 \ \frac{E_p}{S_{\max}} \frac{\theta_2}{\theta_1}, &amp; \text{otherwise} \end{cases}$</td>
<td>$E_a \leq \frac{S}{\Delta t}$</td>
<td>$E_a = \min \left( \frac{S}{\theta_1 \cdot \frac{S_{\max}}{E_p}} \max \left( \frac{S}{\theta_1 \cdot \frac{S_{\max}}{E_p}} \right), \frac{S}{\Delta t} \right)$</td>
<td>44</td>
<td></td>
</tr>
<tr>
<td>Exchange</td>
<td>Water exchange between aquifer and channel</td>
<td>exchange_1</td>
<td>$f_{\text{flux out}} = \begin{cases} \theta_1 \frac{S}{\Delta t} + \theta_2 \left( 1 - \exp \left( -\theta_3 \frac{S}{\Delta t} \right) \right), &amp; \text{if } S \geq 0 \ -\theta_1 \frac{S}{\Delta t} + \theta_2 \left( 1 - \exp \left( -\theta_3 \frac{S}{\Delta t} \right) \right), &amp; \text{if } S &lt; 0 \end{cases}$</td>
<td>No constraint</td>
<td>$f_{\text{flux out}} = \max \left( \theta_1 \cdot \frac{S}{\Delta t} + \theta_2 \cdot \left( 1 - \exp \left( -\theta_3 \cdot \frac{S}{\Delta t} \right) \right), \text{sign}(S), -f_{\text{flux in}} \right)$</td>
<td>36</td>
</tr>
<tr>
<td>Water exchange based on relative storages</td>
<td>exchange_2</td>
<td>$f_{\text{flux out}} = \theta_1 \frac{S_1}{S_{\max}} - \frac{S_2}{S_{\max}}$</td>
<td></td>
<td>$f_{\text{flux out}} = \theta_1 \frac{S_1}{S_{\max}} - \frac{S_2}{S_{\max}}$</td>
<td>38</td>
<td></td>
</tr>
<tr>
<td>Water exchange with infinite size store based on threshold</td>
<td>exchange_3</td>
<td>$f_{\text{flux out}} = \theta_1 \cdot (S - \theta_2)$</td>
<td></td>
<td>$f_{\text{flux out}} = \theta_1 \cdot (S - \theta_2)$</td>
<td>36</td>
<td></td>
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</tr>
<tr>
<td>Infiltration</td>
<td>Infiltration as exponentially declining based on relative storage (taken from a flux)</td>
<td>infiltration_1</td>
<td>( f_{\text{flux}<em>\text{out}} = \theta_1 \cdot \exp \left[ -\theta_2 \frac{S}{S</em>{\text{max}}} \right] )</td>
<td>( f_{\text{flux}<em>\text{out}} = f</em>{\text{flux}_\text{in}} )</td>
<td>( f_{\text{flux}<em>\text{out}} = \min \left( \theta_1 \cdot \exp \left[ -\theta_2 \frac{S}{S</em>{\text{max}}} \right], f_{\text{flux}_\text{in}} \right) )</td>
<td>18, 36, 44</td>
</tr>
<tr>
<td></td>
<td>Delayed infiltration as exponentially declining based on relative storage (taken from a store)</td>
<td>infiltration_2</td>
<td>( f_{\text{flux}<em>\text{out}} = \theta_1 \cdot \exp \left[ -\theta_2 \frac{S_1}{S</em>{1,\text{max}}} \right] - f_{\text{flux}_\text{used}} )</td>
<td>( 0 \leq f_{\text{flux}_\text{out}} \leq \frac{S_2}{\Delta t} )</td>
<td>( f_{\text{flux}<em>\text{out}} = \max \left( \min \left( \theta_1 \cdot \exp \left[ -\theta_2 \frac{S_1}{S</em>{1,\text{max}}} \right], f_{\text{flux}_\text{used}} \right), 0 \right) )</td>
<td>36</td>
</tr>
<tr>
<td></td>
<td>Infiltration to soil moisture of liquid water stored in snow pack</td>
<td>infiltration_3</td>
<td>( f_{\text{flux}<em>\text{out}} = \begin{cases} f</em>{\text{flux}<em>\text{in}}, &amp; \text{if } S \geq S</em>{\text{max}} \ 0, &amp; \text{otherwise} \end{cases} )</td>
<td></td>
<td>( f_{\text{flux}<em>\text{out}} = f</em>{\text{flux}<em>\text{in}} \left[ 1 - \text{logisticSmother}</em>{S(S,S_{\text{max}})} \right] )</td>
<td>37</td>
</tr>
<tr>
<td></td>
<td>Constant infiltration rate</td>
<td>infiltration_4</td>
<td>( f_{\text{flux}_\text{out}} = \theta_1 )</td>
<td>( f_{\text{flux}<em>\text{out}} = f</em>{\text{flux}_\text{in}} )</td>
<td>( f_{\text{flux}<em>\text{out}} = \min (f</em>{\text{flux}_\text{in}}, \theta_1) )</td>
<td>15, 23, 40, 44</td>
</tr>
<tr>
<td></td>
<td>Maximum infiltration rate non-linearly based on relative deficit and storage</td>
<td>infiltration_5</td>
<td>( f_{\text{flux}<em>\text{out}} = \theta_1 \left( 1 - \frac{S_1}{S</em>{1,\text{max}}} \right) \left( \frac{S_2}{S_{2,\text{max}}} \right)^{\theta_2} )</td>
<td>To prevent complex numbers, ( S = [0, x] )</td>
<td>To prevent numerical issues with a theoretical infinite infiltration rate, ( f_{\text{flux}_\text{out}} &lt; 10^{-9} )</td>
<td>23</td>
</tr>
<tr>
<td></td>
<td>Infiltration rate non-linearly scaled by relative storage</td>
<td>infiltration_6</td>
<td>( f_{\text{flux}<em>\text{out}} = \theta_1 \left( \frac{S}{S</em>{\text{max}}} \right)^{\theta_2} f_{\text{flux}_\text{in}} )</td>
<td>( f_{\text{flux}<em>\text{out}} = f</em>{\text{flux}_\text{in}} )</td>
<td>( f_{\text{flux}<em>\text{out}} = \min \left( \theta_1 \cdot \max \left( 0, \frac{S}{S</em>{\text{max}}} \right)^{\theta_2}, f_{\text{flux}_\text{in}} \right) )</td>
<td>43</td>
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<tr>
<td>Interception</td>
<td>Interception excess when maximum capacity is reached</td>
<td>interception_1</td>
<td>( f_{\text{flux}<em>\text{out}} = \begin{cases} f</em>{\text{flux}<em>\text{in}}, &amp; \text{if } S \geq S</em>{\text{max}} \ 0, &amp; \text{otherwise} \end{cases} )</td>
<td></td>
<td>( f_{\text{flux}<em>\text{out}} = f</em>{\text{flux}<em>\text{in}} \left[ 1 - \text{logisticSmother}</em>{S(S,S_{\text{max}})} \right] )</td>
<td>16, 18, 22, 26, 34, 36, 39, 42, 44, 45</td>
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<td>Interception excess after a constant amount is intercepted</td>
<td>interception_2</td>
<td>( f_{\text{flux}<em>\text{out}} = \begin{cases} f</em>{\text{flux}<em>\text{in}} - \theta_1, &amp; \text{if } f</em>{\text{flux}_\text{in}} \geq 0 \ 0, &amp; \text{otherwise} \end{cases} )</td>
<td>( f_{\text{flux}_\text{out}} \geq 0 )</td>
<td>( f_{\text{flux}<em>\text{out}} = \max (f</em>{\text{flux}_\text{in}} - \theta_1, 0) )</td>
<td>2, 13, 15</td>
</tr>
<tr>
<td></td>
<td>Interception excess after a fraction is intercepted</td>
<td>interception_3</td>
<td>( f_{\text{flux}_\text{out}} = \theta_1 )</td>
<td></td>
<td>( f_{\text{flux}_\text{out}} = \theta_1 )</td>
<td>8</td>
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<tr>
<td></td>
<td>Interception excess after a time-varying fraction is intercepted</td>
<td>interception_4</td>
<td>( f_{\text{flux}<em>\text{out}} = \left( \theta_1 + (1 - \theta_1) \cdot \cos \left( \frac{\Delta t - \theta_2}{\Delta t</em>{\text{max}}} \right) \right) \cdot f_{\text{flux}_\text{in}} )</td>
<td>( f_{\text{flux}_\text{out}} \geq 0 )</td>
<td>( f_{\text{flux}<em>\text{out}} = \max \left( 0, \theta_1 + (1 - \theta_1) \cdot \cos \left( \frac{\Delta t - \theta_2}{\Delta t</em>{\text{max}}} \right) \right) \cdot f_{\text{flux}_\text{in}} )</td>
<td>32, 35</td>
</tr>
<tr>
<td></td>
<td>Interception excess after a combined absolute amount and fraction are intercepted</td>
<td>interception_5</td>
<td>( f_{\text{flux}<em>\text{out}} = \theta_1 \cdot f</em>{\text{flux}<em>\text{in}} - \theta_2, \text{if } f</em>{\text{flux}_\text{in}} \geq 0 \ 0, \text{otherwise} )</td>
<td>( f_{\text{flux}_\text{out}} \geq 0 )</td>
<td>( f_{\text{flux}<em>\text{out}} = \max (\theta_1 \cdot f</em>{\text{flux}_\text{in}} - \theta_2, 0) )</td>
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<tbody>
<tr>
<td>Interflow</td>
<td>Interflow as a scaled fraction of an incoming flux</td>
<td>interflow_1</td>
<td>( f_{\text{flux}} = \theta_1 S \frac{S}{S_{\text{max}}} )</td>
<td>( f_{\text{flux}} \leq \frac{S}{\Delta t} )</td>
<td>18, 36</td>
<td></td>
</tr>
<tr>
<td>Non-linear interflow</td>
<td></td>
<td>interflow_2</td>
<td>( f_{\text{flux}} = \theta_1 S^{1+\theta_2} )</td>
<td></td>
<td></td>
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<tr>
<td>Non-linear interflow (variant)</td>
<td></td>
<td>interflow_3</td>
<td>( f_{\text{flux}} = \theta_1 S \theta_2 )</td>
<td>( f_{\text{flux}} \leq \frac{S}{\Delta t} )</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Combined linear and scaled quadratic interflow</td>
<td></td>
<td>interflow_4</td>
<td>( f_{\text{flux}} = \theta_1 S + \theta_2 S^2 )</td>
<td>( f_{\text{flux}} \leq \frac{S}{\Delta t} )</td>
<td></td>
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</tr>
<tr>
<td>Linear interflow</td>
<td></td>
<td>interflow_5</td>
<td>( f_{\text{flux}} = \theta_1 S )</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Scaled linear interflow if a storage in the receiving store exceeds a threshold</td>
<td></td>
<td>interflow_6</td>
<td>( f_{\text{flux}} = \left{ \begin{array}{ll} \theta_1 S \frac{S}{S_{\text{max}}} - \theta_2 / (1 - \theta_2), &amp; \text{if } S &lt; S_{\text{max}} \theta_2 \ 0, &amp; \text{otherwise} \end{array} \right. )</td>
<td>( \frac{S_S}{S_{\text{max}}} \leq 1 )</td>
<td>41</td>
<td></td>
</tr>
<tr>
<td>Non-linear interflow if storage exceeds a threshold</td>
<td></td>
<td>interflow_7</td>
<td>( f_{\text{flux}} = \left{ \begin{array}{ll} \left( \frac{S - \theta_1 S_{\text{max}}}{\theta_2} \right)^{\frac{1}{\theta_1}}, &amp; \text{if } \theta_2 S_{\text{max}} &gt; S \ 0, &amp; \text{otherwise} \end{array} \right. )</td>
<td>( f_{\text{flux}} \leq \frac{S - \theta_1 S_{\text{max}}}{\Delta t} )</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Linear interflow if storage exceeds a threshold</td>
<td></td>
<td>interflow_8</td>
<td>( f_{\text{flux}} = \left{ \begin{array}{ll} \theta_1 (S - \theta_2), &amp; \text{if } \theta_2 &gt; S \ 0, &amp; \text{otherwise} \end{array} \right. )</td>
<td>( f_{\text{flux}} = \max(0, \theta_1 (S - \theta_2)) )</td>
<td>3, 12, 27, 38</td>
<td></td>
</tr>
<tr>
<td>Non-linear interflow if storage exceeds a threshold (variant)</td>
<td></td>
<td>interflow_9</td>
<td>( f_{\text{flux}} = \left{ \begin{array}{ll} \theta_1 (S - \theta_2)^{\theta_2}, &amp; \text{if } \theta_2 &gt; S \ 0, &amp; \text{otherwise} \end{array} \right. )</td>
<td>( f_{\text{flux}} \leq \frac{S - \theta_2}{\Delta t} )</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Scaled linear interflow if storage exceeds a threshold</td>
<td></td>
<td>interflow_10</td>
<td>( f_{\text{flux}} = \left{ \begin{array}{ll} \theta_1 S \frac{S}{S_{\text{max}}} - \theta_2 / \theta_3, &amp; \text{if } S &gt; \theta_2 \ 0, &amp; \text{otherwise} \end{array} \right. )</td>
<td>( f_{\text{flux}} = \max(0, S - \theta_2) / \theta_3 )</td>
<td>14</td>
<td></td>
</tr>
<tr>
<td>Constant interflow if storage exceeds a threshold</td>
<td></td>
<td>interflow_11</td>
<td>( f_{\text{flux}} = \theta_1, \text{if } \theta_2 &gt; S \ 0, \text{otherwise} )</td>
<td>( f_{\text{flux}} \leq \frac{S - \theta_2}{\Delta t} )</td>
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<tr>
<td>Misc</td>
<td>Auxiliary function to find contributing area</td>
<td>area_1</td>
<td>[ A = \begin{cases} \theta_1 &amp; \text{if } S \geq S_{\text{min}} \ \frac{S - S_{\text{min}}}{S_{\text{max}} - S_{\text{min}}} &amp; \text{if } S &lt; S_{\text{min}} \end{cases} ]</td>
<td>[ A \leq 1 ]</td>
<td>[ A = \min \left( 1, \theta_1 \left[ \frac{S - S_{\text{min}}}{S_{\text{max}} - S_{\text{min}}} \right]^\gamma_1 \right) ] * ( 1 - \text{logisticSmooth}(S, S_{\text{min}}) )</td>
<td>23</td>
</tr>
<tr>
<td></td>
<td>General effective flow (returns flux [mm/d])</td>
<td>effective_1</td>
<td>[ f_{\text{flux}} = \begin{cases} f\text{lux}<em>{\text{in},1} - f\text{lux}</em>{\text{in},2} &amp; \text{if } f\text{lux}<em>{\text{in},1} &gt; f\text{lux}</em>{\text{in},2} \ 0 &amp; \text{otherwise} \end{cases} ]</td>
<td>[ \text{flux}<em>{\text{out}} = \min(0, f</em>{\text{flux}}) ]</td>
<td>( \text{flux}<em>{\text{out}} = \max(0, f</em>{\text{flux}}) )</td>
<td>22, 23, 25, 29, 40, 42, 43, 44, 45, 46</td>
</tr>
<tr>
<td></td>
<td>Storage excess when store size changes (returns flux [mm/d])</td>
<td>excess_1</td>
<td>[ f_{\text{flux}} = \frac{S - S_{\text{max,new}}}{\Delta t} ]</td>
<td>[ \text{flux}_{\text{out}} \geq 0 ]</td>
<td>( \text{flux}<em>{\text{out}} = \max\left( \frac{S - S</em>{\text{max,new}}}{\Delta t}, 0 \right) )</td>
<td>10, 19, 22, 37, 44</td>
</tr>
<tr>
<td></td>
<td>Phenology-based correction factor for potential evapotranspiration (returns flux [mm/d])</td>
<td>phenology_1</td>
<td>[ E_p = \begin{cases} 0, &amp; \text{if } T(t) &lt; \theta_1 \ \frac{T(t) - \theta_1}{\theta_2 - \theta_1} \cdot E_p, &amp; \text{if } \theta_1 \leq T(t) &lt; \theta_2 \ E_p, &amp; \text{if } T(t) \geq \theta_2 \end{cases} ]</td>
<td>[ E_p = \min\left( 1, \max\left( 0, \frac{T(t) - \theta_1}{\theta_2 - \theta_1} \right) \right) ]</td>
<td>35</td>
<td></td>
</tr>
<tr>
<td></td>
<td>Phenology-based maximum interception capacity (returns store size [mm])</td>
<td>phenology_2</td>
<td>[ S_{\text{max}} = \theta_1 \left( 1 + \theta_2 \sin\left( \frac{2\pi t \Delta t - \theta_3}{\tau_{\text{max}}} \right) \right) ]</td>
<td>Assumes ( 0 \leq \theta_2 \leq 1 ) to guarantee ( S_{\text{max}} \geq 0 )</td>
<td>[ S_{\text{max}} = \theta_1 \left( 1 + \theta_2 \sin\left( \frac{2\pi t \Delta t - \theta_3}{\tau_{\text{max}}} \right) \right) ]</td>
<td>22</td>
</tr>
<tr>
<td></td>
<td>Split flow (returns flux [mm/d])</td>
<td>split_1</td>
<td>[ f_{\text{flux}} = \theta_1 \cdot f_{\text{flux}} ]</td>
<td>[ f_{\text{flux}} = \theta_1 \cdot f_{\text{flux}} ]</td>
<td>5, 11, 13, 17, 21, 25, 26, 28, 29, 33, 34, 40, 41, 42, 43, 45, 46</td>
<td></td>
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<tr>
<td>Percolation</td>
<td>Percolation at a constant rate</td>
<td>percolation_1</td>
<td>[ f_{\text{flux}} = \begin{cases} \theta_1 &amp; \text{if } S \geq 0 \ 0 &amp; \text{otherwise} \end{cases} ]</td>
<td>[ f_{\text{flux}} = \frac{S}{\Delta t} ]</td>
<td>[ f_{\text{flux}} = \min\left( \frac{S}{\Delta t}, \theta_1 \right) ]</td>
<td>37</td>
</tr>
<tr>
<td></td>
<td>Percolation scaled by current relative storage</td>
<td>percolation_2</td>
<td>[ f_{\text{flux}} = \theta_1 \frac{S}{S_{\text{max}}} ]</td>
<td>[ f_{\text{flux}} = \frac{S}{\Delta t} ]</td>
<td>[ f_{\text{flux}} = \min\left( \frac{S}{\Delta t}, \theta_1 \frac{S}{S_{\text{max}}} \right) ]</td>
<td>21, 26, 34</td>
</tr>
<tr>
<td></td>
<td>Non-linear percolation (empirical)</td>
<td>percolation_3</td>
<td>[ f_{\text{flux}} = \frac{S_{\text{max}}^4}{4} \left( \frac{\theta_1}{\theta_2} \right)^4 S^5 ]</td>
<td></td>
<td>[ f_{\text{flux}} = \frac{S_{\text{max}}^4}{4} \left( \frac{\theta_1}{\theta_2} \right)^4 S^5 ]</td>
<td>7</td>
</tr>
<tr>
<td></td>
<td>Demand-based percolation scaled by available moisture</td>
<td>percolation_4</td>
<td>[ f_{\text{flux}} = S \sum_{i=1}^{\text{deficiencies}} \prod_{k=1}^n \left( 1 + \theta_2 \left( \frac{\sum_{i=1}^{\text{deficiencies}} \text{capacities} \right)^\gamma_2 \right) ]</td>
<td>[ f_{\text{flux}} \leq \frac{S}{\Delta t} ]</td>
<td>[ f_{\text{flux}} = \max\left( 0, \min\left( \frac{S}{\Delta t}, \frac{S_{\text{max}} (S, 0)}{S_{\text{max}}} \right) \right) ]</td>
<td>33</td>
</tr>
<tr>
<td></td>
<td>Non-linear percolation</td>
<td>percolation_5</td>
<td>[ f_{\text{flux}} = \theta_1 \left( \frac{S}{S_{\text{max}}} \right)^\theta_2 ]</td>
<td>[ f_{\text{flux}} \leq \frac{S}{\Delta t} ]</td>
<td>[ f_{\text{flux}} = \min\left( \frac{S}{\Delta t}, \theta_1 \left( \frac{S}{S_{\text{max}}} \right)^\theta_2 \right) ]</td>
<td>22</td>
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<tr>
<td>Threshold-based percolation</td>
<td>from a store that can reach negative values</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \frac{\theta_1}{\theta_2} ) if ( S \geq \theta_2 ) [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>( \frac{\partial S}{\partial t} \leq \frac{S}{\theta_2} )[15, 16, 17, 18, 19, 20, 22, 24, 25, 30, 31, 32, 33, 35, 36, 39, 40, 41, 44, 45, 46]</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ S - \theta_2 \min \left( 1, \frac{\max(0, S)}{\theta_2} \right) ]</td>
<td>39</td>
<td></td>
</tr>
<tr>
<td>Recharge</td>
<td>Recharge as scaled fraction of incoming flux</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \frac{S}{S_{\text{max}}} ) * flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \frac{S}{S_{\text{max}}} ) * flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>( \frac{\partial S}{\partial t} \leq \frac{S}{\theta_2} )[7, 37, 45]</td>
<td>18, 36</td>
<td></td>
</tr>
<tr>
<td>Recharge</td>
<td>Linear recharge</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \frac{S}{S_{\text{max}}} ) * flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>To prevent complex numbers, ( S = [0, x])</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \frac{\min(1, S_{\text{max}})}{S_{\text{max}}} ) * flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>19, 23, 24, 27, 30, 31, 32, 35, 38, 42</td>
<td></td>
</tr>
<tr>
<td>Recharge</td>
<td>Constant recharge from a store</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq 0 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \frac{S}{S_{\text{max}}} ] * flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq 0 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>23, 44</td>
<td></td>
</tr>
<tr>
<td>Recharge</td>
<td>Recharge to fulfil evaporation demand</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \theta_2 \left[ 1 - \frac{S}{S_{\text{max}}} \right], \text{ if } S &lt; \theta_2 ) [\theta_1, \text{ if } S &gt; \theta_2 ]</td>
<td>To prevent complex numbers, ( S = [0, x])</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>20</td>
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<tr>
<td>Recharge</td>
<td>Non-linear recharge</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \theta_2 S^{\theta_2} )</td>
<td>To prevent complex numbers, ( S = [0, x])</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>44</td>
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<tr>
<td>Recharge</td>
<td>Constant recharge from a flux</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \theta_2 )</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>45</td>
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<tr>
<td>Routing</td>
<td>Threshold-based non-linear routing</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = ( \theta_2 S^{\theta_2} )</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = flux&lt;sub&gt;in&lt;/sub&gt;</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_2, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ if } S &lt; \theta_2 ]</td>
<td>39</td>
<td></td>
</tr>
<tr>
<td>Saturation excess</td>
<td>Saturation excess from a store that has reached maximum capacity</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_1, \text{ if } S \geq \theta_2 ] [\theta_1, \text{ otherwise} ]</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \theta_1, \text{ otherwise} ]</td>
<td>flux&lt;sub&gt;out&lt;/sub&gt; = [ \min(\theta_2, \text{max}(S, 0)^{\theta_2}, \theta_1 \frac{S}{S_{\text{max}}}) ]</td>
<td>1, 3, 4, 6, 8, 9, 10, 11, 12, 14, 15, 16, 17, 18, 19, 20, 22, 24, 25, 30, 31, 32, 33, 35, 36, 39, 40, 41, 44, 45, 46</td>
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<tr>
<td>Saturation excess from a store with different degrees of saturation</td>
<td>saturation_2</td>
<td>flux_out</td>
<td>( 1 - \left( 1 - \frac{S}{S_{\text{max}}} \right)^{\theta_1} ) * flux_in</td>
<td>To prevent complex numbers, ( S/S_{\text{max}} = [0,\infty) )</td>
<td>2, 13, 22, 28, 29</td>
<td>39</td>
</tr>
<tr>
<td>Saturation excess from a store with different degrees of saturation (exponential variant)</td>
<td>saturation_3</td>
<td>flux_out</td>
<td>( 1 - \frac{1}{1 + \exp \left( \frac{S - S_{\text{max}}}{\theta_1} \right) + 0.5} ) * flux_in</td>
<td></td>
<td>21, 26, 34</td>
<td></td>
</tr>
<tr>
<td>Saturation excess from a store with different degrees of saturation (quadratic variant)</td>
<td>saturation_4</td>
<td>flux_out</td>
<td>( 1 - \left( \frac{S}{S_{\text{max}}} \right)^2 ) * flux_in</td>
<td>( 0 \leq \text{flux}_{\text{out}} )</td>
<td>7</td>
<td></td>
</tr>
<tr>
<td>Deficit store exponential saturation excess based on current storage and a threshold parameter</td>
<td>saturation_5</td>
<td>flux_out</td>
<td>( 1 - \min \left( 1, \frac{S}{\theta_2} \right) ) * flux_in</td>
<td>To prevent complex numbers, ( S = [0,\infty) )</td>
<td>5</td>
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<tr>
<td>Saturation excess from a store with different degrees of saturation linear variant)</td>
<td>saturation_6</td>
<td>flux_out</td>
<td>( \theta_1 \frac{S}{S_{\text{max}}} ) * flux_in</td>
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<td>40</td>
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<tr>
<td>Saturation excess from a store with different degrees of saturation (gamma function variant)</td>
<td>saturation_7</td>
<td>flux_out</td>
<td>( \frac{1}{\theta_1 \Gamma(\theta_2)} \int_{\theta_2}^{\infty} (x - \theta_2)^{\theta_2 - 1} e^{-\frac{x - \theta_2}{\theta_1}} ) * flux_in</td>
<td>To prevent numerical problems, ( S = [0,\infty) )</td>
<td>14</td>
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<tr>
<td>Saturation excess flow from a store with different degrees of saturation (min-max linear variant)</td>
<td>saturation_8</td>
<td>flux_out</td>
<td>( \theta_1 + (\theta_2 - \theta_1) \frac{S}{S_{\text{max}}} ) * flux_in</td>
<td>( \text{flux}<em>{\text{out}} \leq \text{flux}</em>{\text{in}} )</td>
<td>45</td>
<td></td>
</tr>
<tr>
<td>Deficit store saturation excess from a store that has reached maximum capacity</td>
<td>saturation_9</td>
<td>flux_out</td>
<td>( \text{flux}_{\text{out}} ) if ( S = 0 ) ( 0, \text{otherwise} )</td>
<td></td>
<td>17, 25, 43, 46</td>
<td></td>
</tr>
<tr>
<td>Saturation excess flow from a store with different degrees of saturation (min-max exponential variant)</td>
<td>saturation_10</td>
<td>flux_out</td>
<td>( \min (\theta_1, \theta_2 + \theta_2 e^{\theta_3}) ) * flux_in</td>
<td></td>
<td>39</td>
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<tr>
<td>Saturation excess flow from a store with different degrees of saturation (min exponential variant)</td>
<td>saturation_11</td>
<td>$\text{flux}<em>{\text{out}} = \left( \frac{S - S</em>{\text{min}}}{S_{\text{max}} - S_{\text{min}}} \right)^{\theta_2} \text{flux}<em>{\text{in}}$ if $S &gt; S</em>{\text{min}}$ $0$, otherwise</td>
<td>$\text{flux}<em>{\text{out}} \leq f\text{lux}</em>{\text{in}}$</td>
<td>$\text{flux}<em>{\text{out}} = \min \left( 1, \frac{S - S</em>{\text{min}}}{S_{\text{max}} - S_{\text{min}}} \right) \text{flux}_{\text{in}} \times \left[ 1 - \text{logisticSmother}<em>S(S, S</em>{\text{min}}) \right]$</td>
<td></td>
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<tr>
<td>Saturation excess flow from a store with different degrees of saturation (min-max linear variant)</td>
<td>saturation_12</td>
<td>$\text{flux}<em>{\text{out}} = \frac{\theta_1 - \theta_2}{1 - \theta_2} \text{flux}</em>{\text{in}}$</td>
<td>$\text{flux}_{\text{out}} \geq 0$</td>
<td>$\text{flux}<em>{\text{out}} = \max \left( 0, \frac{\theta_1 - \theta_2}{1 - \theta_2} \right) \text{flux}</em>{\text{in}}$</td>
<td></td>
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</tr>
<tr>
<td>Saturation excess flow from a store with different degrees of saturation (normal distribution variant)</td>
<td>saturation_13</td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \int_0^\infty \frac{1}{\sqrt{2\pi} \xi} \exp \left[ -\frac{\xi^2}{2} \right] d\xi$, with $\xi = \frac{\log(S/\theta_1)}{\log(\theta_1/\theta_2)}$</td>
<td></td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \text{normcdf} \left( \frac{\log(\max(0.5)/\theta_1)}{\log(\theta_1/\theta_2)} \right)$</td>
<td></td>
<td>42</td>
</tr>
<tr>
<td>Saturation excess flow from a store with different degrees of saturation (two-part exponential variant)</td>
<td>saturation_14</td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \left{ \begin{array}{ll} (0.5 - \theta_1)^{1-\theta_2} \left( \frac{S}{S_{\text{max}}} \right)^{\theta_1} &amp; \text{if } \frac{S}{S_{\text{max}}} \leq 0.5 - \theta_1 \ 1 - (0.5 - \theta_1)^{1-\theta_2} \left( 1 - \frac{S}{S_{\text{max}}} \right)^{\theta_2} &amp; \text{otherwise} \end{array} \right.$</td>
<td></td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \left( \left( 0.5 - \theta_1 \right)^{1-\theta_2} \left( \frac{0.5}{S_{\text{max}}} \right)^{\theta_1} \right)^* \times \text{flux}_{\text{in}}$</td>
<td></td>
<td>28</td>
</tr>
<tr>
<td>Snowfall based on temperature threshold</td>
<td>snowfall_1</td>
<td>$\text{flux}<em>{\text{out}} = {\text{flux}</em>{\text{in}}, \text{if } T \leq T_{\text{threshold}} }$ $0$, otherwise</td>
<td></td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \left[ \text{logisticSmother}<em>T(T, T</em>{\text{threshold}}) \right]$</td>
<td></td>
<td>6, 12, 30, 31, 32, 34, 35, 41, 43, 44, 45</td>
</tr>
<tr>
<td>Snowfall based on a temperature threshold interval</td>
<td>snowfall_2</td>
<td>$\text{flux}<em>{\text{out}} = {\text{flux}</em>{\text{in}}, \text{if } T \leq \theta_1 - \frac{1}{2} \theta_2 \ \frac{\theta_1 + \frac{1}{2} \theta_2 - T}{\theta_2}, \text{if } \theta_1 - \frac{1}{2} \theta_2 &lt; T &lt; \theta_1 + \frac{1}{2} \theta_2 \ 0, \text{if } T \geq \theta_1 + \frac{1}{2} \theta_2 }$</td>
<td></td>
<td>$\text{flux}<em>{\text{out}} = \min \left( \text{flux}</em>{\text{in}} \text{max} \left( 0, \text{flux}_{\text{in}} \times \frac{\theta_1 + \frac{1}{2} \theta_2 - T}{\theta_2} \right) \right)$</td>
<td></td>
<td>37</td>
</tr>
<tr>
<td>Rainfall based on temperature threshold</td>
<td>rainfall_1</td>
<td>$\text{flux}<em>{\text{out}} = {\text{flux}</em>{\text{in}}, \text{if } T &gt; T_{\text{threshold}} }$ $0$, otherwise</td>
<td></td>
<td>$\text{flux}<em>{\text{out}} = \text{flux}</em>{\text{in}} \times \left[ 1 - \text{logisticSmother}<em>T(T, T</em>{\text{threshold}}) \right]$</td>
<td></td>
<td>6, 12, 30, 31, 32, 34, 35, 41, 43, 44, 45</td>
</tr>
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continued ...
<table>
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<tr>
<th>Process</th>
<th>Details</th>
<th>Function name</th>
<th>Constitutive function</th>
<th>Constraints</th>
<th>MARRMoT Code</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Snowfall based on a temperature threshold interval</td>
<td>rainfall_2</td>
<td>fluxout</td>
<td>0, if ( T \leq \theta_1 - \frac{1}{2} \theta_2 ) [ \begin{align*} \text{fluxout} &amp;= \frac{\theta_1 + \frac{1}{2} \theta_2 - T}{\theta_2}, \text{if } \theta_1 - \frac{1}{2} \theta_2 &lt; T &lt; \theta_1 + \frac{1}{2} \theta_2 \ &amp;= \text{fluxin}, \text{if } T \geq \theta_1 + \frac{1}{2} \theta_2 \end{align*} ]</td>
<td>( \text{fluxout} = \min \left( \text{fluxin, max} \left( 0, \frac{T - \theta_1 + \frac{1}{2} \theta_2}{\theta_2} \right) \right) )</td>
<td>37</td>
<td></td>
</tr>
<tr>
<td>Refreezing of stored melted snow</td>
<td>refreeze_1</td>
<td>fluxout</td>
<td>( \theta_1 \times \theta_2 \times (T_{\text{threshold}} - T), \text{if } T \leq T_{\text{threshold}} ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} \leq \frac{S}{\Delta t} ) ( \text{fluxout} = \min \left( \frac{S}{\Delta t}, \text{max} \left( 0, \theta_1 \times \theta_2 \times (T_{\text{threshold}} - T) \right) \right) )</td>
<td>37, 44</td>
<td></td>
</tr>
<tr>
<td>Snowmelt from degree-day-factor</td>
<td>melt_1</td>
<td>fluxout</td>
<td>( \theta_1 \times (T - T_{\text{threshold}}), \text{if } T \geq T_{\text{threshold}} ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} \leq \frac{S}{\Delta t} ) ( \text{fluxout} = \min \left( \frac{S}{\Delta t}, \text{max} \left( 0, \theta_1 \times (T - T_{\text{threshold}}) \right) \right) )</td>
<td>6, 12, 30, 31, 32, 34, 35, 37, 43, 44, 45</td>
<td></td>
</tr>
<tr>
<td>Snowmelt at a constant rate</td>
<td>melt_2</td>
<td>fluxout</td>
<td>( \theta_1 \times S \geq 0 ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} \leq \frac{S}{\Delta t} ) ( \text{fluxout} = \min \left( \frac{S}{\Delta t}, \theta_1 \right) )</td>
<td>44</td>
<td></td>
</tr>
<tr>
<td>Glacier melt provided no snow is stored on the ice layer</td>
<td>melt_3</td>
<td>fluxout</td>
<td>( \theta_1 \times (T - T_{\text{threshold}}), \text{if } T \geq T_{\text{threshold}}, S_2 = 0 ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} \leq \frac{S}{\Delta t} ) ( \text{fluxout} = \min \left( \frac{S}{\Delta t}, \text{max} \left( 0, \theta_1 \times (T - T_{\text{threshold}}) \right) \right) ) ( \text{logisticSmoothing}_{S}(S_2, 0) )</td>
<td>43</td>
<td></td>
</tr>
<tr>
<td>Soil moisture</td>
<td>soilmoisture_1</td>
<td>fluxout</td>
<td>( \frac{S_2 \Delta S_{\text{max}} - S_1 S_{\text{max}}}{S_{\text{max}} + S_{\text{max}}}, \text{if } \frac{S_1}{S_{\text{max}}} &lt; \frac{S_2}{S_{\text{max}}} ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} = \frac{S_2 \Delta S_{\text{max}} - S_1 S_{\text{max}}}{S_{\text{max}} + S_{\text{max}}} \times \text{logisticSmoothing}_{S}(S_1, S_2) )</td>
<td>33</td>
<td></td>
</tr>
<tr>
<td>Water balance to equal relative storage (2 stores)</td>
<td>soilmoisture_2</td>
<td>fluxout</td>
<td>( \frac{S_2 (S_2 \Delta S_{\text{max}} + S_3 \Delta S_{\text{max}}) + S_1 \Delta S_{\text{max}} (S_2 + S_3)}{S_{\text{max}} + S_{\text{max}}}, \text{if } \frac{S_1}{S_{\text{max}}} &lt; \frac{S_2}{S_{\text{max}}} + \frac{S_3}{S_{\text{max}}} ) ( 0, \text{otherwise} )</td>
<td>( \text{fluxout} = \frac{S_2 (S_2 \Delta S_{\text{max}} + S_3 \Delta S_{\text{max}}) + S_1 \Delta S_{\text{max}} (S_2 + S_3)}{S_{\text{max}} + S_{\text{max}}} \times \text{logisticSmoothing}_{S}(S_1, S_2, S_3) )</td>
<td>33</td>
<td></td>
</tr>
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## S4 Unit Hydrographs

This section provides details on the implementation of various Unit Hydrographs. An overview of the 7 UHs is given in Table S2. Computational implementation of each UH is given in sections S4.1 to S4.7. Unit Hydrograph files can be found in "./MARRMoT/Models/Unit Hydrograph files/".

<table>
<thead>
<tr>
<th>File name</th>
<th>Inputs</th>
<th>Diagram</th>
<th>Description</th>
<th>In model ...</th>
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<tbody>
<tr>
<td>uh_1_half</td>
<td>1: amount to be routed&lt;br&gt;2: time base&lt;br&gt;3: Δt</td>
<td><img src="image1.png" alt="Diagram" /></td>
<td>Exponentially increasing scheme</td>
<td>7</td>
</tr>
<tr>
<td>uh_2_full</td>
<td>1: amount to be routed&lt;br&gt;2: time base (time is doubled inside the function)&lt;br&gt;3: Δt</td>
<td><img src="image2.png" alt="Diagram" /></td>
<td>Exponential triangular scheme</td>
<td>7</td>
</tr>
<tr>
<td>uh_3_half</td>
<td>1: amount to be routed&lt;br&gt;2: time base&lt;br&gt;3: Δt</td>
<td><img src="image3.png" alt="Diagram" /></td>
<td>Triangular scheme: 13, 15, 21, 26 linearly increasing 34</td>
<td></td>
</tr>
<tr>
<td>uh_4_full</td>
<td>1: amount to be routed&lt;br&gt;2: time base&lt;br&gt;3: Δt</td>
<td><img src="image4.png" alt="Diagram" /></td>
<td>Triangular scheme: 0 (template), linearly increasing 16, 37, and decreasing nn (example)</td>
<td></td>
</tr>
<tr>
<td>uh_5_half</td>
<td>1: amount to be routed&lt;br&gt;2: time base&lt;br&gt;3: Δt</td>
<td><img src="image5.png" alt="Diagram" /></td>
<td>Exponentially decreasing scheme</td>
<td>5</td>
</tr>
<tr>
<td>uh_6_gamma</td>
<td>1: amount to be routed&lt;br&gt;2: gamma parameter [-]&lt;br&gt;3: time for flow to reduce by factor e [d]&lt;br&gt;4: length of time series</td>
<td><img src="image6.png" alt="Diagram" /></td>
<td>Gamma function-based</td>
<td>40</td>
</tr>
<tr>
<td>uh_7_uniform</td>
<td>1: amount to be routed&lt;br&gt;2: time base&lt;br&gt;3: Δt</td>
<td><img src="image7.png" alt="Diagram" /></td>
<td>Uniform distribution</td>
<td>39</td>
</tr>
</tbody>
</table>
S4.1 Code: uh_1_half

This section provides the computational implementation of a unit hydrograph with an increasing exponential distribution of flows.

File location ./MARRMoT/Models/Unit Hydrograph files/uh_1_half

References E.g. GR4J Perrin et al. (2003)

```matlab
function [ out, UH ] = uh_1_half( in, d_base, delta_t )
%uh_1_half Unit Hydrograph [days] with half a bell curve.
% GR4J-based

% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY WARRANTY. See <https://www.gnu.org/licenses/> for details.

% Inputs
% in - volume to be routed
% d_base - time base of routing delay [d]
% delta_t - time step size [d]

% Unit hydrograph spreads the input volume over a time period x4.
% Percentage of input returned only increases.
% I.e. d_base = 3.8 [days], delta_t = 1:
% UH(1) = 0.04 [% of inflow]
% UH(2) = 0.17
% UH(3) = 0.35
% UH(4) = 0.45

% Inputs
if any(size(in)) > 1; error('UH input should be a single value.'); end

% Time step size
delay = d_base/delta_t;
if delay == 0; delay = 1; end % any value below t = 1 means no delay,
% but zero leads to problems
tt = 1:ceil(delay);
% Time series for which we need UH
% ordinates [days]
```

135
%%EMPTIES
SH = zeros(1,length(tt)+1); SH(1) = 0;
UH = zeros(1,length(tt));

%%UNIT HYDROGRAPH
for t = tt
    if t < delay; SH(t+1) = (t./delay).^(5./2);
    elseif t >= delay; SH(t+1) = 1;
    end
    UH(t) = SH(t+1)-SH(t);
end

%%DISPERSE VOLUME
out = in.*UH;
end

S4.2 Code: uh_2_full

This section provides the computational implementation of a unit hydrograph with an exponential triangular distribution of flows.

File location  ./MARRMoT/Models/Unit Hydrograph files/uh_2_full
References  E.g. GR4J Perrin et al. (2003)

function [ out , UH ] = uh_2_full( in , d_base , delta_t )
%uh_2_full Unit Hydrograph [days] with a full bell curve. 
%Copyright (C) 2018 W. Knoben
%This program is free software (GNU GPL v3) and 
distributed WITHOUT ANY
%WARRANTY. See <https://www.gnu.org/licenses/> for details
%
%Inputs 
% in - volume to be routed 
% d_base - time base of routing delay [d] 
% delta_t - time step size [d] 
% Unit hydrograph spreads the input volume over a time period 2*x4.
Knoben et al., 2018

14  % Percentage of input returned goes up (till x4), then
down again.
15  % I.e. d_base = 3.8 [days], delta_t = 1:
16  % UH(1) = 0.02 [% of inflow]
17  % UH(2) = 0.08
18  % UH(3) = 0.18
19  % UH(4) = 0.29
20  % UH(5) = 0.24
21  % UH(6) = 0.14
22  % UH(7) = 0.05
23  % UH(8) = 0.00
24
25  % INPUTS
26  if any(size(in)) > 1; error('UH input should be a single
value.'); end
27
28  % TIME STEP SIZE
29  delay = d_base/delta_t;
30  tt = 1:2*ceil(delay); % time series for which we need UH ordinates [days]
31
32  % EMPTIES
33  SH = zeros(1,length(tt)+1); SH(1) = 0;
34  UH = zeros(1,length(tt));
35
36  % UNIT HYDROGRAPH
37  for t = tt
38      if (t <= delay)
39          SH(t+1) = 0.5*(t./delay).^(5./2);
40      elseif (t > delay) && (t < 2*delay);
41          SH(t+1) = 1-0.5*(2-t./delay).^(5./2);
42      elseif (t >= 2*delay);
43          SH(t+1) = 1;
44      end
45
46      UH(t) = SH(t+1)-SH(t);
47  end
48
49  % DISPERSE VOLUME
50  out = in.*UH;
51
52  end
S4.3 Code: uh_3_half

This section provides the computational implementation of a unit hydrograph with an linearly increasing distribution of flows.

File location ./MARRMoT/Models/Unit Hydrograph files/uh_3_half
References E.g. FLEX-Topo Savenije (2010)

function [ out, UH ] = uh_3_half( in, d_base, delta_t )

% uh_3_half Unit Hydrograph [days] with half a triangle ( linear )
% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY WARRANTY. See <https://www.gnu.org/licenses/> for details.

% Inputs
% in - volume to be routed
% d_base - time base of routing delay [d]
% delta_t - time step size [d]
% Unit hydrograph spreads the input volume over a time period delay.
% Percentage of input returned only increases.
% I.e. d_base = 3.8 [days], delta_t = 1:
% UH(1) = 0.04 [% of inflow]
% UH(2) = 0.17
% UH(3) = 0.35
% UH(4) = 0.45

%% INPUTS
if any(size(in)) > 1; error('UH input should be a single value.'); end

%% TIME STEP SIZE
delay = d_base/delta_t;
if delay == 0; delay = 1; end % any value below t = 1 means no delay,
% but zero leads to problems
tt = 1:ceil(delay);
  need UH

% ordinates [days]
% The area under the unit hydrograph by definition sums to 1. Thus the area
% is S(t=0 to t = delay) t*[ff: fraction of flow to move per time step] dt
% Analytical solution is [1/2 * t^2 + c]*ff, with c = 0.
% Thus the fraction
% of flow step size is:
ff = 1/(0.5*delay^2);

%% EMPTIES
UH = zeros(1,length(tt));

%% UNIT HYDROGRAPH
for t = 1:length(tt)
    if t <= delay
        UH(t) = ff.*(0.5*t^2 - 0.5*(t-1)^2);
    else
        UH(t) = ff.*(0.5*delay^2 - 0.5*(t-1)^2);
    end
end

%% DISPERSE VOLUME
out = in.*UH;

S4.4 Code: uh_4_full
This section provides the computational implementation of a unit hydrograph with
an linear triangular distribution of flows.

File location ./MARRMoT/Models/Unit Hydrograph files/uh_4_full
References E.g. HBV-96
citeLindstrom1997

function [ out, UH ] = uh_4_full( in, d_base, delta_t )
%uh_4_half Unit Hydrograph [days] with a triangle (linear)
% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY
% WARRANTY. See <https://www.gnu.org/licenses/> for details.
% Inputs
% in - volume to be routed
% d_base - time base of routing delay [d]
% delta_t - time step size [d]
% Unit hydrograph spreads the input volume over a time period delay.
% Percentage runoff goes up, peaks, and goes down again.
% I.e. d_base = 3.8 [days], delta_t = 1:
% UH(1) = 0.14 [% of inflow]
% UH(2) = 0.41
% UH(3) = 0.36
% UH(4) = 0.09

%% INPUTS
if any(size(in)) > 1; error('UH input should be a single value.'); end

%% TIME STEP SIZE
delay = d_base/delta_t;
if delay == 0; delay = 1; end % any value below t = 1 means no delay,
% but zero leads to problems
% time series for which we need UH
% ordinates [days]

tt = 1:ceil(delay);

%% UNIT HYDROGRAPH
% The area under the unit hydrograph by definition sums to 1. Thus the area
% is S(t=0 to t = delay) t*[ff: fraction of flow to move per time step] dt
% Analytical solution is [1/2 * t^2 + c]*ff, with c = 0.
% Here, we use two half triangles t make one big one, so the area of half a
% triangle is 0.5. Thus the fraction of flow step size is:
% ff = 0.5/(0.5*(0.5*delay)^2);
d50 = 0.5*delay;

%% TRIANGLE FUNCTION
tri = @(t) max(ff.*(t-d50).*sign(d50-t)+ff.*d50,0);

%% EMPTIES
S4.5 Code: uh_5_half

This section provides the computational implementation of a unit hydrograph with an decreasing exponential distribution of flows.

File location ./MARRMoT/Models/Unit Hydrograph files/uh_5_half
References E.g. IHACRES Littlewood et al. (1997); Croke and Jakeman (2004)

```
function [ out, UH ] = uh_5_half( in, d_base, delta_t )
% uh_5_half Unit Hydrograph [days] with decreasing exponential distribution of flows.
% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY WARRANTY. See <https://www.gnu.org/licenses/> for details.
% Inputs
% in - volume to be routed
% d_base - time base of routing delay [d]
% delta_t - time step size [d]
% Unit hydrograph spreads the input volume over a time period x4.
% I.e. d_base = 3.8 [days], delta_t = 1;

UH = zeros(1,length(tt));

%% UNIT HYDROGRAPH
for t = 1:length(tt)
    UH(t) = integral(tri,t-1,t);
end

%% ENSURE UH SUMS TO 1
tmp_diff = 1-sum(UH);
tmp_weight = UH./sum(UH);
UH = UH + tmp_weight.*tmp_diff;

%% DISPERSE VOLUME
out = in.*UH;
```
% UH(1) = 0.04 [% of inflow]
% UH(2) = 0.17
% UH(3) = 0.35
% UH(4) = 0.45

%% INPUTS
if any(size(in)) > 1; error('UH input should be a single value. '); end

%% TIME STEP SIZE
delay = d_base/delta_t;
if delay == 0; delay = 1; end % any value below t = 1 means no delay,
% but zero leads to problems

tt = 1:ceil(delay); % Time series for which we need UH
% ordinates [days]

%% EMPTIES
SH = zeros(1,length(tt)+1); SH(1) = 0;
UH = zeros(1,length(tt));

%% UNIT HYDROGRAPH
for t = tt
    if t < delay; SH(t+1) = (t./delay).^(5./2);
    elseif t >= delay; SH(t+1) = 1;
    end
    UH(t) = SH(t+1) - SH(t);
end

%% DISPERSE VOLUME
out = in.*UH;
end

S4.6 Code: uh_6_gamma
This section provides the computational implementation of a unit hydrograph with a gamma distribution of flows.

File location ./MARRMoT/Models/Unit Hydrograph files/uh_6_gamma
References E.g. SMAR O’Connell et al. (1970); Tan and O’Connor (1996)
function [ out,UH,frac_routing_beyond_time_series ] = uh_6_gamma( in,n,k, t_end, delta_t )

%uh_6_gamma Unit Hydrograph [days] from gamma function.
%
% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY WARRANTY. See <https://www.gnu.org/licenses/> for details.
%
% Inputs
% n = shape parameter [-]
% k = time delay for flow reduction by a factor e [d]
% t_end = length of time series [d]
% delta_t = time step size [d]
%
% Unit hydrograph spreads the input volume over a time period delay.
% Percentage of input returned only decreases.
% I.e. n = 1, k = 3.8 [days], delta_t = 1:
% UH(1) = 0.928 [% of inflow]
% UH(2) = 0.067
% UH(3) = 0.005
% UH(4) = 0.000
%
% INPUTS
if any(size(in)) > 1; error('UH input should be a single value.'); end
%
% TIME STEP SIZE
tt = 1:tmax; % time series for which we need UH ordinates [days]
%
% EMPTIES
UH_full = zeros(1,length(tt));
frac_routing_beyond_time_series = 0;
%
% UNIT HYDROGRAPH
% The Unit Hydrograph follows a gamma distribution. For a given
% delay time, the fraction of flow per time step is thus
the integral of
% t-1 to t of the gamma distribution. The curve has range
[0, Inf].
% We need to choose a point at which to cap the integration,
% but this
% depends on the parameters n & k, and the total time step.
We choose the
% cutoff point at the time step where less than 0.1% of the
% peak flow
% is still on route.

%% Unit hydrograph
for t = 1:length(tt)
    UH_full(t) = integral(@(x) 1./(k.*gamma(n)).*(x./k).^(n -1).* ...
                        exp(-1.*x./k),(t-1)*
                        delta_t,t*delta_t);
end

%% Find cutoff point where less than 0.1% of the peak flow
% is being routed
[max_val, max_here] = max(UH_full);
end_here = find(UH_full(max_here:end)./max_val<0.001,1) +
            max_here;

%% Take action depending on whether the distribution
% function exceeds the
% time limit or not
if ~isempty(end_here)
    %% Construct the Unit Hydrograph
    UH = UH_full(1:end_here);
    % Account for the truncated part of the full UH.
    % find probability mass to the right of the cut-off
    point
    tmp_excess = 1-sum(UH);
    % find relative size of each time step
    tmp_weight = UH_full(1:end_here)./sum(UH_full(1:
                                            end_here));
    % distribute truncated probability mass proportionally
to all elements
    % of the routing vector
    UH = UH+tmp_weight.*tmp_excess;
end
else
    
    UH = UH_full;

    frac_routing_beyond_time_series = 1-sum(UH);

end

%% DISPERSE VOLUME
out = in.*UH;

end

S4.7 Code: uh_7_uniform

This section provides the computational implementation of a unit hydrograph with a uniform distribution of flows.

File location ./.MARRMoT/Models/Unit Hydrograph files/uh_7_uniform
References E.g. MCRM?Moore and Bell (2001)

function [ out,UH ] = uh_7_uniform( in, d_base, delta_t )
%uh_7_uniform Unit Hydrograph [days] with uniform spread
% Copyright (C) 2018 W. Knoben
% This program is free software (GNU GPL v3) and distributed WITHOUT ANY
% WARRANTY. See <https://www.gnu.org/licenses/> for details.
% Inputs
% in - volume to be routed
% d_base - time base of routing delay [d]
% delta_t - time step size [d]
% Unit hydrograph spreads the input volume over a time period delay.
% I.e. d_base = 3.8 [days], delta_t = 1:
% UH(1) = 0.26 [% of inflow]
%% INPUTS
if any(size(in)) > 1; error('UH input should be a single value.'); end

%% TIME STEP SIZE
delay = d_base/delta_t;
tt = 1:ceil(delay); % time series for which we need UH ordinates [days]

%% EMPTIES
UH = NaN.*zeros(1,length(tt));

%% FRACTION FLOW
ff = 1/delay; % fraction of flow per time step

%% UNIT HYDROGRAPH
for t=1:ceil(delay)
    if t < delay
        UH(t) = ff;
    else
        UH(t) = mod(delay,t-1)*ff;
    end
end

%% DISPERSE VOLUME
out = in.*UH;
end
S5 Parameter ranges

Each model function in MARRMoT is accompanied by a file that specifies suitable sampling ranges for each parameter used in the model, that could be applied if the user chooses to pair MARRMoT with a calibration or parameter sampling procedure. This section gives the reasoning behind our choices of parameter ranges used within MARRMoT.

S5.1 Model-specific ranges versus generalised process-specific ranges

There are two different approaches to determining parameter ranges for model calibration or parameter sampling studies: (1) make a choice for appropriate parameter ranges per model, based on previous applications of the model, or (2) try to make consistent choices for all models based on literature (e.g. ensure that all 'slow' linear reservoirs, regardless of which model they are part of, have the same limits for the drainage time scale parameter). Generalization of parameter ranges across models is difficult because models use different flux formulations and thus different parameter values might be appropriate, even if the fluxes are intended to represent the same hydrologic process. On the other hand, using model-specific parameter ranges based on earlier studies might limit a model’s potential. Especially if the model has only been applied to a small number of places, published 'appropriate' parameter ranges might also reflect the climate or catchment characteristics of the few study catchments the model has been applied to. MARRMoT is intended as a model comparison framework. We thus attempt to generalize parameter ranges across all models in the framework, to facilitate fair comparison of different models. We try to err on the side of caution and intentionally set these ranges wide. Table S3 shows the parameter ranges used in MARRMoT and specifies in which model(s) each parameter range is used.
| Description                                                                 | Min(lit) | Max(lit) | Min(used) | Max(used) | Reference(s)                                                                 | Notes                                                                 | Model              |
|----------------------------------------------------------------------------|----------|----------|-----------|-----------|------------------------------------------------------------------------------|                                                                     |                   |
| Snow                                                                       |          |          |           |           | [°C]                                                                         |                                                                     |                   |
| **Threshold temperature for snowfall (and melt, if not specified otherwise)** | Table S4 | Table S4 | -3        | 5         | Kienzle (2008); Kollat et al. (2012)                                         |                                                                     | 6, 12, 30, 31, 32, 34, 35, 37, 43, 44, 45 |
| **Threshold interval width for snowfall [°C]**                              | 0        | 7        | 0         | 17        | Kienzle (2008)                                                               | 0 is a physical limit                                                | 37                |
| **Threshold temperature for melt [°C]**                                     | -3       | 3        | -3        | 3         |                                                                                  | Not easy to find any interval. Temperature for melt tends to be treated as constant at 0 | 37, 43, 44        |
| **Degree-day-factor for snow or ice melt [mm/°C/d]**                       | 0        | Table S5 | 0         | 20        | Kollat et al. (2012)                                                         | 0 is a physical limit                                                | 6, 12, 30, 31, 32, 34, 35, 37, 41, 43, 44, 45 |
| **Water holding content of snow pack [-]**                                 | 0        | 0.8      | 0         | 1         | Kollat et al. (2012)                                                         | [0,1] are physical limits                                            | 37, 44            |
| **Refreezing factor of retained liquid water [-]**                          | 0        | 1        | 0         | 1         |                                                                                  | [0,1] are physical limits                                            | 37, 44            |
| **Maximum melt rate due to ground-heat flux [mm/d]**                        | 0        | 2        | 0         | 2         | Schaeffli et al. (2014)                                                      |                                                                     | 44                |

continued...
<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
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<tr>
<td><strong>Interception</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Maximum store depth [mm]</td>
<td>0</td>
<td>Table S6</td>
<td>0</td>
<td>5</td>
<td>Chiew and McMahon (1994); Gerrits (2010)</td>
<td>0 is a physical limit. Gerrits (2010) reports 3.8mm as maximum value used out of 15 studies. Chiew and McMahon (1994) (table 3) report 5.6mm as maximum value for 28 catchments 2, 13, 15, 16, 18, 22, 23, 26, 34, 36, 39, 42, 44, 45</td>
<td>2, 13, 15, 16, 18, 22, 23, 26, 34, 36, 39, 42, 44, 45</td>
</tr>
<tr>
<td>Maximum intercepted fraction of precipitation [-]</td>
<td>0</td>
<td>0.42</td>
<td>0</td>
<td>1</td>
<td>Gerrits (2010)</td>
<td>0,1 are physical limits. Gerrits (2010) reports 42% as maximum intercepted fraction out of 15 studies 8, 23, 32, 35, 45</td>
<td>8, 23, 32, 35, 45</td>
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<tr>
<td>Seasonal variation in LAI as fraction of mean [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>0 is a physical limit 22</td>
<td>22</td>
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<tr>
<td>Timing of maximum Leaf Area Index [d]</td>
<td>1</td>
<td>365</td>
<td></td>
<td></td>
<td></td>
<td>Refers to days in a normal calendar year 22</td>
<td>22, 32, 35</td>
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<tr>
<td><strong>Surface depression</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Maximum surface area contributing to store [-]</td>
<td>0</td>
<td>1</td>
<td>0</td>
<td>1</td>
<td></td>
<td>[0,1] are physical limits 36, 45</td>
<td>36, 45</td>
</tr>
<tr>
<td>Maximum store depth [mm]</td>
<td>0</td>
<td>Table S7</td>
<td>0</td>
<td>50</td>
<td>Chiew and McMahon (1994)</td>
<td>0 is physical limit. 50 is recommended in Chiew and McMahon (1994) 36, 45</td>
<td>36, 45</td>
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<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
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<th>Min(used)</th>
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<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Filling parameter [-]</strong></td>
<td>1</td>
<td>1</td>
<td>0.99</td>
<td>1</td>
<td>Chiew (1990); Porter and McMahon (1971)</td>
<td>Controls the shape of the depression store inflow flux but is usually set at 1 because no studies are (were?) available about how a depression store fills</td>
<td></td>
</tr>
<tr>
<td><strong>Infiltration</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Maximum loss [mm]</td>
<td>0</td>
<td>400</td>
<td>0</td>
<td>600</td>
<td>Chiew et al. (2002)</td>
<td>Fig 11.11a shows calibrated parameter values for 339 catchments. Pattern indicates that limit was set at 400</td>
<td>18, 36</td>
</tr>
<tr>
<td>Loss exponent [-]</td>
<td>0</td>
<td>12</td>
<td>0</td>
<td>15</td>
<td>Chiew et al. (2002)</td>
<td>Fig 11.11a shows calibrated parameter values for 339 catchments. Pattern indicates that limit was set at 10</td>
<td>18, 36</td>
</tr>
<tr>
<td>Maximum infiltration rate [mm/d]</td>
<td>Table S8</td>
<td>Table S8</td>
<td>0</td>
<td>200</td>
<td></td>
<td>Infiltration rates can be very high. However, to have a practical effect on modelling, (i.e. generate infiltration excess flow), Inf_rate &lt; P(t). In the context of a follow-up study, Inf_rate is capped at 200mm/d because the maximum daily P in the study area is 200mm/d.</td>
<td>15, 20, 23, 40, 44</td>
</tr>
</tbody>
</table>

continued ...
<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Infiltration decline non-linearity parameter [-]</td>
<td>0</td>
<td>5</td>
<td></td>
<td></td>
<td>Sivapalan et al. (1996)</td>
<td>Very difficult to find information for (original paper mentions nothing)</td>
<td>23, 43</td>
</tr>
<tr>
<td>Evaporation</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Plant-controlled maximum rate [mm/d]</td>
<td>5</td>
<td>24.5</td>
<td>0</td>
<td>20</td>
<td>Chiew and McMahon (1994)</td>
<td>Although the study reports an upper value of 24.5, the recommended range is capped at 20 (paper appendix)</td>
<td>20, 36</td>
</tr>
<tr>
<td>Wilting point as fraction of Soil moisture capacity [-]</td>
<td>0.1</td>
<td>0.25</td>
<td>0.05</td>
<td>0.95</td>
<td>Son and Sivapalan (2007)</td>
<td>0 is a physical limit but can break model equations through &quot;divide-by-zero&quot; errors. 1 is a physical limit</td>
<td>3, 4, 8, 9, 10, 12, 14, 15, 16, 19, 20, 21, 26, 31, 32, 34, 35, 37, 44</td>
</tr>
<tr>
<td>Moisture constrained rate parameter [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td>15</td>
</tr>
<tr>
<td>Forest fraction for separate soil/vegetation evap [-]</td>
<td>0</td>
<td>1</td>
<td>0.05</td>
<td>0.95</td>
<td></td>
<td>[0,1] are physical limits, but using these limits can result in divide-by-zero-errors in certain fluxes</td>
<td>3, 4, 8, 16</td>
</tr>
<tr>
<td>Phenology: minimum temperature where transpiration stops [°C]</td>
<td>-5</td>
<td>-5</td>
<td>0</td>
<td>-10</td>
<td>Ye et al. (2012)</td>
<td></td>
<td>35</td>
</tr>
</tbody>
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<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
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<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Phenology: maximum temperature above which transpiration fully utilizes Ep [°C]</td>
<td>10</td>
<td>10</td>
<td>1</td>
<td>20</td>
<td>Ye et al. (2012)</td>
<td>The setup of minimum and maximum temperature used in Ye et al. (2012) is here changed to a minimum temperature + temperature range (Tmax = Tmin + Trange) to avoid overlap in parameter values</td>
<td>35</td>
</tr>
<tr>
<td>Evaporation reduction with depth coefficient [-]</td>
<td>0.083</td>
<td>1</td>
<td>0</td>
<td>1</td>
<td>Penman (1950); Tan and O’Connor (1996)</td>
<td>[0,1] are physical limits</td>
<td>17, 23, 25, 40</td>
</tr>
<tr>
<td>Shape parameter for evaporation reduction in a deficit store [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td>Moore and Bell (2001)</td>
<td>This uses a sigmoid function to determine a fraction of Ep to evaporate. Values &gt;1 make the transition very steep</td>
<td>39</td>
</tr>
<tr>
<td>Evaporation non-linearity coefficient [-]</td>
<td>0</td>
<td>10</td>
<td></td>
<td></td>
<td>Sivapalan et al. (1996)</td>
<td>Very difficult to find information for. Assumption made to be in line with other non-linearity coefficients.</td>
<td>23, 43</td>
</tr>
<tr>
<td>Soil moisture</td>
<td></td>
<td></td>
<td>1</td>
<td>2000</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Maximum store depth [mm]</td>
<td>1</td>
<td>Table S9</td>
<td>1</td>
<td>2000</td>
<td></td>
<td>0 is a physical limit</td>
<td>1, 2, 3, 4, 5, 6, 7, 8, 9, 10, 11, 12, 13, 14, 15, 16, 17, 18, 19, 20, 21, 22, 23, 24, 26, 27, 28, 29, 30, 31, 32, 33, 34, 35, 36, 37, 38, 39, 40, 41, 42, 43, 44, 45, 46</td>
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<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
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</thead>
<tbody>
<tr>
<td>Capillary rise [mm/d]</td>
<td>0</td>
<td>Table S10</td>
<td>0</td>
<td>4</td>
<td></td>
<td>0 is a physical limit</td>
<td>13, 15, 37, 38</td>
</tr>
<tr>
<td>Percolation rate [mm/d]</td>
<td>0</td>
<td>Table S11</td>
<td>0</td>
<td>20</td>
<td>Bethune et al. (2008)</td>
<td>Some modelling studies report very large percolation rates (100 mm/d). Bethune et al. (2008) report 11 mm/d from field observations.</td>
<td>21, 26, 34, 37, 39, 44, 45</td>
</tr>
<tr>
<td>Percolation fraction [-]</td>
<td>0.013</td>
<td>0.533</td>
<td>0</td>
<td>1</td>
<td>Ye et al. (2012) (Table 1)</td>
<td>[0,1] are physical limits</td>
<td>14, 22, 23, 24, 27, 30, 31, 32, 35, 45</td>
</tr>
<tr>
<td>Recharge nonlinearity [-]</td>
<td>0</td>
<td>7</td>
<td>0</td>
<td>10</td>
<td>Kollat et al. (2012)</td>
<td>Also seen as a soil depth distribution</td>
<td>5, 22, 33, 37</td>
</tr>
<tr>
<td>Soil depth distribution [-]</td>
<td>0</td>
<td>Table S12</td>
<td>0</td>
<td>10</td>
<td></td>
<td>For cases where the soil depth is not considered constant. Most studies limit this to 0-2.5 but this seems based on a single source (Wagener et al., 2004) which is UK only. Thus we use a wider range here</td>
<td>2, 13, 15, 21, 22, 26, 28, 29, 34</td>
</tr>
<tr>
<td>Porosity [-]</td>
<td>0.35</td>
<td>0.5</td>
<td>0.05</td>
<td>0.95</td>
<td>Son and Sivapalan (2007)</td>
<td>[0,1] are theoretical physical limits, but no (0) porosity and full (1) porosity are not sensible: there would be no soil moisture or soil respectively</td>
<td>10, 19</td>
</tr>
</tbody>
</table>

continued ...
<table>
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<tr>
<th>Description</th>
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<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
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<tbody>
<tr>
<td>Gamma distribution for topographic indices - phi [-]</td>
<td>0.4</td>
<td>3.5</td>
<td>0.1</td>
<td>5</td>
<td>Clark et al. (2008)</td>
<td></td>
<td>14</td>
</tr>
<tr>
<td>Gamma distribution for topographic indices - chi [-]</td>
<td>2</td>
<td>5</td>
<td>1</td>
<td>7.5</td>
<td>Clark et al. (2008)</td>
<td></td>
<td>14</td>
</tr>
<tr>
<td>Fraction area with permeable soils [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td>Crooks and Naden (2007)</td>
<td>[0,1] are physical limits</td>
<td>46</td>
</tr>
<tr>
<td>Fraction area with semi-permeable soils [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td>Crooks and Naden (2007)</td>
<td>[0,1] are physical limits</td>
<td>46</td>
</tr>
<tr>
<td>Fraction area with impermeable soils [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td>Crooks and Naden (2007)</td>
<td>[0,1] are physical limits</td>
<td>46</td>
</tr>
<tr>
<td>Variable contributing area scaling [-]</td>
<td>0</td>
<td>5</td>
<td></td>
<td></td>
<td>Sivapalan et al. (1996)</td>
<td>Very difficult to find information about this. Assumption made</td>
<td>23</td>
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<tr>
<td>Variable contributing area non-linearity [-]</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Sivapalan et al. (1996)</td>
<td>See: <strong>Soil depth distribution</strong> above</td>
<td>23</td>
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<tr>
<td>Fraction of D50 that is D16 [-]</td>
<td>0.01</td>
<td>0.99</td>
<td></td>
<td></td>
<td></td>
<td>Note: re-writing of D16 parameter in Fukushima (1988)</td>
<td>42</td>
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<tr>
<td>Variable contributing area equation inflection point [-]</td>
<td>-0.5</td>
<td>0.5</td>
<td>-0.5</td>
<td>0.5</td>
<td>Jayawardena and Zhou (2000)</td>
<td></td>
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<tr>
<td><strong>Groundwater</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Leakage coefficient [-]</td>
<td>0.07</td>
<td>0.13</td>
<td>0</td>
<td>0.5</td>
<td>Chiew and McMahon (1994)</td>
<td>0 is physical limit, 0.5 is recommended in the paper’s appendix</td>
<td>36</td>
</tr>
<tr>
<td>Leakage rate [mm/d]</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td>See: <strong>Percolation rate</strong> above</td>
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<thead>
<tr>
<th>Description</th>
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<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
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<tbody>
<tr>
<td>Level compared to channel level [mm]</td>
<td>-2.8</td>
<td>3.9</td>
<td>-10</td>
<td>10</td>
<td>Chiew and McMahon (1994)</td>
<td>Range recommended in appendix of the paper</td>
<td>36</td>
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<tr>
<td>Base flow rate at no deficit [mm/d]</td>
<td>0</td>
<td>201.6</td>
<td>0.1</td>
<td>200</td>
<td>Beven (1997)</td>
<td>Based on Table 2 (Beven, 1997)</td>
<td>14, 23</td>
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<tr>
<td>Basflow deficit scaling parameter [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td>14, 23</td>
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<tr>
<td><strong>Flow distribution</strong></td>
<td></td>
<td></td>
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<td></td>
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<td></td>
<td></td>
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<tr>
<td>Interflow and saturation excess [-]</td>
<td>0</td>
<td>1</td>
<td>0</td>
<td>1</td>
<td></td>
<td>[0,1] are physical limits</td>
<td></td>
</tr>
<tr>
<td>Preferential recharge [-]</td>
<td>0</td>
<td>2</td>
<td>0</td>
<td>1</td>
<td>Chiew and McMahon (1994)</td>
<td>0 is a physical limit. Later paper sets max limit to 1</td>
<td></td>
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<tr>
<td>Surface/groundwater division [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td></td>
</tr>
<tr>
<td>Fast and slow flow [-]</td>
<td>0</td>
<td>1</td>
<td>0</td>
<td>1</td>
<td>Son and Sivapalan (2007)</td>
<td>[0,1] are physical limits</td>
<td></td>
</tr>
<tr>
<td>Groundwater recharge and interflow [-]</td>
<td>0.05</td>
<td>0.3</td>
<td>0</td>
<td>1</td>
<td></td>
<td>[0,1] are physical limits</td>
<td></td>
</tr>
<tr>
<td>Infiltration and direct runoff [-]</td>
<td>0.161</td>
<td>0.422</td>
<td>0</td>
<td>1</td>
<td>Tan and O’Connor (1996)</td>
<td>[0,1] are physical limits</td>
<td></td>
</tr>
<tr>
<td>Impervious and infiltration area [-]</td>
<td>0</td>
<td>1</td>
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<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td></td>
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<tr>
<td>Contributing area to overland flow [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
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<td>[0,1] are physical limits</td>
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<tr>
<td>Tension water and free water [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td></td>
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<tr>
<td>Threshold for overland flow generation [-]</td>
<td>0</td>
<td>&lt;1</td>
<td>0.99</td>
<td></td>
<td>Nielsen and Hansen (1973)</td>
<td>[0,1] are physical limits</td>
<td>41</td>
</tr>
</tbody>
</table>

continued ...
<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Threshold for overland flow generation [-]</td>
<td>0</td>
<td>&lt;1</td>
<td>0</td>
<td>0.99</td>
<td>Nielsen and Hansen (1973)</td>
<td>[0,1] are physical limits</td>
<td>41</td>
</tr>
<tr>
<td>Channel and land division [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td>42</td>
</tr>
<tr>
<td>Throughfall/stem flow division [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td>42</td>
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<tr>
<td>Glacier/non-glacier precipitation [-]</td>
<td>0</td>
<td>1</td>
<td></td>
<td></td>
<td></td>
<td>[0,1] are physical limits</td>
<td>43</td>
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<tr>
<td><strong>Flow time scale and shape</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
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<tr>
<td>Fast reservoir time scale [d⁻¹]</td>
<td>0.05</td>
<td>Table</td>
<td>0</td>
<td>1</td>
<td>0 is a physical limit</td>
<td></td>
<td>12, 21, 24, 26, 28, 29, 30, 31, 32, 33, 34, 35, 37, 39, 41, 42, 43, 44, 46</td>
</tr>
<tr>
<td>Slow reservoir time scale [d⁻¹]</td>
<td>0.01</td>
<td>Table</td>
<td>0</td>
<td>1</td>
<td>0 is a physical limit</td>
<td></td>
<td>2, 3, 4, 6, 8, 10, 13, 15, 16, 17, 18, 19, 20, 21, 22, 24, 25, 26, 28, 29, 30, 31, 32, 33, 34, 35, 37, 39, 40, 41, 42, 43, 44, 46</td>
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<tr>
<td>Flow non-linearity Ṡx [-]</td>
<td>0</td>
<td>Table</td>
<td>1</td>
<td>5</td>
<td>0 is a physical limit</td>
<td></td>
<td>4, 9, 10, 11, 16, 19, 22, 23, 37, 39, 42, 44, 45</td>
</tr>
<tr>
<td>Flow reduction (S/X) [mm]</td>
<td>5</td>
<td>40</td>
<td>1</td>
<td>50</td>
<td>Son and Sivapalan (2007)</td>
<td></td>
<td>9</td>
</tr>
<tr>
<td>Exponential shape parameter [mm⁻¹]</td>
<td>0</td>
<td>2</td>
<td></td>
<td></td>
<td>Moore and Bell (2001)</td>
<td>Very difficult to find documentation for</td>
<td>39</td>
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</table>

continued ...
### Routing

<table>
<thead>
<tr>
<th>Description</th>
<th>Min(lit)</th>
<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
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<tbody>
<tr>
<td>Routing delay to fast flow [d]</td>
<td>0</td>
<td>1</td>
<td>1</td>
<td>5</td>
<td>Fenicia et al. (2008)</td>
<td></td>
<td>5, 21, 26, 34</td>
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<tr>
<td>Routing delay to slow flow [d]</td>
<td>0</td>
<td>8</td>
<td>1</td>
<td>15</td>
<td></td>
<td></td>
<td>5, 7, 21, 26, 34</td>
</tr>
<tr>
<td>Routing delay [d]</td>
<td>1</td>
<td>120</td>
<td>1</td>
<td>120</td>
<td>Kollat et al. (2012)</td>
<td>1 is the limit (water shouldn't speed up). 120 because it seems very high</td>
<td>13, 15, 16, 21, 37, 39, 40</td>
</tr>
<tr>
<td>Routing store depth [mm]</td>
<td>1</td>
<td>300</td>
<td>1</td>
<td>300</td>
<td>Perrin et al. (2003)</td>
<td>0 would mean no routing, so slightly above that</td>
<td>7, 20, 39, 45</td>
</tr>
<tr>
<td>Gamma function, number of Nash cascade reservoirs [-]</td>
<td>0.75</td>
<td>9.79</td>
<td>1</td>
<td>10</td>
<td></td>
<td></td>
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### Water exchange parameters

<table>
<thead>
<tr>
<th>Coefficient 1 [-]</th>
<th>0.005</th>
<th>0.54</th>
<th>0</th>
<th>1</th>
<th>Chiew and McMahon (1994)</th>
<th>Although the study only reports values up to 0.54, an upper range of 1 is recommended in the study's appendix</th>
<th>36</th>
</tr>
</thead>
<tbody>
<tr>
<td>Coefficient 2 [-]</td>
<td>0.01</td>
<td>0.29</td>
<td>0</td>
<td>1</td>
<td>Chiew and McMahon (1994)</td>
<td>Although the study only reports values up to 0.29, an upper range of 1 is recommended in the study's appendix</td>
<td>36</td>
</tr>
<tr>
<td>Coefficient 3 [-]</td>
<td>0</td>
<td>13</td>
<td>0</td>
<td>100</td>
<td>Chiew and McMahon (1994)</td>
<td>Although the study only reports values up to 13, an upper range of 100 is recommended in the study's appendix</td>
<td>36</td>
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</table>

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<thead>
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<th>Description</th>
<th>Min(lit)</th>
<th>Max(lit)</th>
<th>Min(used)</th>
<th>Max(used)</th>
<th>Reference(s)</th>
<th>Notes</th>
<th>Model</th>
</tr>
</thead>
<tbody>
<tr>
<td>Water exchange coefficient [mm/d]</td>
<td>-10</td>
<td>14</td>
<td>-10</td>
<td>15</td>
<td>Perrin et al. (2003); Santos et al. (2017)</td>
<td>Parameter x2 in GR4J model</td>
<td>7</td>
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</table>
Table S4: Literature-based ranges for snowmelt parameter "threshold temperature for snowfall"

<table>
<thead>
<tr>
<th>Threshold temperature for snowfall [°C]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>-2.5</td>
<td>2.5</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>-3</td>
<td>3</td>
</tr>
<tr>
<td>Table 2 in Kienzle (2008) Note: always coupled with a snow interval [10,17]</td>
<td>1.1</td>
<td>4.5</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>-1.5</td>
<td>2.5</td>
</tr>
</tbody>
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Table S5: Literature-based ranges for snowmelt parameter "degree-day-factor"

<table>
<thead>
<tr>
<th>Degree-day factor for snowmelt [mm/°C/d]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>1</td>
<td>10</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>0</td>
<td>20</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>1</td>
<td>10</td>
</tr>
</tbody>
</table>

Table S6: Literature-based ranges for interception parameter "maximum interception capacity"

<table>
<thead>
<tr>
<th>Interception bucket [mm]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Figure 11.11a in Chiew et al. (2002)</td>
<td>0</td>
<td>5</td>
</tr>
<tr>
<td>Table 3 in Chiew and McMahon (1994)</td>
<td>0.5</td>
<td>5.6</td>
</tr>
<tr>
<td>Table 1.1 in Gerrits (2010)</td>
<td>0</td>
<td>3.8</td>
</tr>
<tr>
<td>Table 2 in Son and Sivapalan (2007)</td>
<td></td>
<td>0.4</td>
</tr>
</tbody>
</table>
Table S7: Literature-based ranges for depression parameter "maximum depression capacity"

<table>
<thead>
<tr>
<th>Depression bucket [mm]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 3 in Chiew and McMahon (1994)</td>
<td>1</td>
<td>100</td>
</tr>
<tr>
<td>Table 1 in Amoah et al. (2013)</td>
<td>5</td>
<td>110</td>
</tr>
</tbody>
</table>

Table S8: Literature-based ranges for infiltration parameter "maximum infiltration rate"

<table>
<thead>
<tr>
<th>Infiltration rate</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Figure 2 in Assouline (2013) [mm/d]</td>
<td>40</td>
<td>100</td>
</tr>
<tr>
<td>Table 3.3 in Jones (1997) [mm/h]</td>
<td>6</td>
<td>76</td>
</tr>
<tr>
<td>Table 3 in Cerdà (1996) [mm/h]</td>
<td>50</td>
<td>770</td>
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</tbody>
</table>

Table S9: Literature-based ranges for soil moisture parameter "maximum soil moisture capacity"

<table>
<thead>
<tr>
<th>Soil moisture bucket [mm]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Figure 11.11b in Chiew et al. (2002)</td>
<td>0</td>
<td>500</td>
</tr>
<tr>
<td>Table 3 in Chiew and McMahon (1994)</td>
<td>65</td>
<td>400</td>
</tr>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>50</td>
<td>500</td>
</tr>
<tr>
<td>Table 1 in Rusli et al. (2015)</td>
<td>100</td>
<td>800</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>0</td>
<td>2000</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>50</td>
<td>500</td>
</tr>
<tr>
<td>Table 3 in Sun et al. (2015)</td>
<td>1</td>
<td>500</td>
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</table>

Table S10: Literature-based ranges for capillary rise parameter "maximum capillary rise rate"

<table>
<thead>
<tr>
<th>Capillary rise [mm/d]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 1 in Rusli et al. (2015)</td>
<td>0.1</td>
<td>1</td>
</tr>
<tr>
<td>Default value in SMHI (2004)</td>
<td>1</td>
<td>1</td>
</tr>
<tr>
<td>Figure 3 in Bethune et al. (2008)</td>
<td>0</td>
<td>0.06</td>
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</tbody>
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Table S11: Literature-based ranges for percolation parameter "maximum percolation rate"

<table>
<thead>
<tr>
<th>Percolation rate [mm/d]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>0</td>
<td>6</td>
</tr>
<tr>
<td>Table 1 in Rusli et al. (2015)</td>
<td>0.1</td>
<td>5</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>0</td>
<td>100</td>
</tr>
<tr>
<td>Figure 3 in Bethune et al. (2008)</td>
<td>0</td>
<td>10.4</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>0</td>
<td>3</td>
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Table S12: Literature-based ranges for soil moisture parameter "soil depth distribution non-linearity"

<table>
<thead>
<tr>
<th>Soil depth distribution [-]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 3 in Sun et al. (2015)</td>
<td>0</td>
<td>2</td>
</tr>
<tr>
<td>Figure 9 in Lamb (1999)</td>
<td>0</td>
<td>2.5</td>
</tr>
<tr>
<td>Table 4 in Bulygina et al. (2009)</td>
<td>0</td>
<td>2.5</td>
</tr>
<tr>
<td>Figure 4.12 in Wagener et al. (2004)</td>
<td>0</td>
<td>2</td>
</tr>
<tr>
<td>Page 700 in Sivapalan and Woods (1995)</td>
<td>4.03</td>
<td></td>
</tr>
<tr>
<td>Figure 4 in Huang et al. (2003) Note: estimated values, 97% &lt; 6</td>
<td>0</td>
<td>11.5</td>
</tr>
</tbody>
</table>

Table S13: Literature-based ranges for flow parameter "fast flow time scale"

<table>
<thead>
<tr>
<th>Fast flow time scale [day^{-1}]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>0.05</td>
<td>0.5</td>
</tr>
<tr>
<td>Table 1 in Rusli et al. (2015)</td>
<td>0.05</td>
<td>0.8</td>
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<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>0.01</td>
<td>1</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>0.01</td>
<td>0.4</td>
</tr>
<tr>
<td>Table 3 in Sun et al. (2015)</td>
<td>0.5</td>
<td>1.2</td>
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Table S14: Literature-based ranges for flow parameter "slow flow time scale"

<table>
<thead>
<tr>
<th>Slow flow time scale [day^{-1}]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Figure 11.11b in Chiew et al. (2002)</td>
<td>0</td>
<td>0.3</td>
</tr>
<tr>
<td>Table 2 in Son and Sivapalan (2007)</td>
<td>2.40E-05</td>
<td>0.1</td>
</tr>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>0.001</td>
<td>0.1</td>
</tr>
<tr>
<td>Table 1 in Rusli et al. (2015)</td>
<td>0.0005</td>
<td>0.1</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>0.00005</td>
<td>0.05</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>0.001</td>
<td>0.15</td>
</tr>
<tr>
<td>Table 3 in Sun et al. (2015)</td>
<td>0.001</td>
<td>0.5</td>
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</table>
Table S15: Literature-based ranges for flow parameter "flow non-linearity"

<table>
<thead>
<tr>
<th>Flow non-linearity</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 3 in Lidén and Harlin (2000) $S^{1+var}$</td>
<td>0</td>
<td>3</td>
</tr>
<tr>
<td>Table 1 in Son and Sivapalan (2007) $S^{1/var}$</td>
<td>0.45</td>
<td>0.5</td>
</tr>
<tr>
<td>Table 3 in Jothityangkoon et al. (2001)</td>
<td>0.5</td>
<td>0.5</td>
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</tbody>
</table>

Table S16: Literature-based ranges for routing parameter "routing delay"

<table>
<thead>
<tr>
<th>Routing delay [d]</th>
<th>Min</th>
<th>Max</th>
</tr>
</thead>
<tbody>
<tr>
<td>Table 2 in Seibert (1997)</td>
<td>1</td>
<td>5</td>
</tr>
<tr>
<td>Table 1 in Kollat et al. (2012)</td>
<td>24</td>
<td>120</td>
</tr>
<tr>
<td>Table 3 in Lidén and Harlin (2000)</td>
<td>1</td>
<td>4</td>
</tr>
<tr>
<td>Table 1 in Perrin et al. (2003)</td>
<td>0.5</td>
<td>4</td>
</tr>
<tr>
<td>Table A3 in Seibert and Vis (2012)</td>
<td>1</td>
<td>7</td>
</tr>
<tr>
<td>Table 2 in Atkinson et al. (2003) Note: converted from a flow speed of 0.5m/s and catchment area of 47km$^2$</td>
<td>&lt;1</td>
<td></td>
</tr>
<tr>
<td>Table 3 in Goswami and O’Connor (2010)</td>
<td>12</td>
<td>36</td>
</tr>
<tr>
<td>Table 2 in Vinogradov et al. (2011) Note: approximated from flow velocities and catchment sizes</td>
<td>0.01</td>
<td>4</td>
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</table>
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