



1 **Carbon-nitrogen coupling under three schemes of model representation:**
2 **Traceability analysis**

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4 Zhenggang Du^{1,2}, Ensheng Weng³, Jianyang Xia^{1,2*}, Lifen Jiang⁴, Yiqi Luo^{4,5}, Xuhui Zhou^{1,2,6*}

5
6 ¹*Center for Global Change and Ecological Forecasting, School of Ecological and*
7 *Environmental Sciences, East China Normal University, Shanghai 200062, China*

8 ²*Tiantong National Field Observation Station for Forest Ecosystem, School of Ecological and*
9 *Environmental Sciences, East China Normal University, Shanghai 200062, China*

10 ³*Department of Ecology & Evolutionary Biology, Princeton University, Princeton, NJ, USA*

11 ⁴*Center for Ecosystem Science and Society, Northern Arizona University, AZ, USA*

12 ⁵*Department for Earth System Science, Tsinghua University, Beijing 100084, China*

13 ⁶*Shanghai Institute of Pollution Control and Ecological Security, 1515 North Zhongshan Rd,*
14 *Shanghai 200437, China*

15

16 ***For correspondence:**

17 Xuhui Zhou

18 *School of Ecological and Environmental Sciences*

19 *East China Normal University*

20 *500 Dongchuan Road, Shanghai 200062, China*

21 **Email:** xhzhou@des.ecnu.edu.cn

22 **Tel/Fax:** +86 21 54341275

23 Jianyang Xia

24 *School of Ecological and Environmental Sciences*



25 *East China Normal University*

26 *500 Dongchuan Road, Shanghai 200062, China*

27 **Email:** jyxia@des.ecnu.edu.cn



28 **Abstract** The interaction between terrestrial carbon (C) and nitrogen (N) cycles has been
29 incorporated into more and more land surface models. However, the scheme of C-N coupling
30 differs greatly among models, and how these diverse representations of C-N interactions will
31 affect C-cycle modeling remains unclear. In this study, we explored how the simulated
32 ecosystem C storage capacity in the terrestrial ecosystem (TECO) model varies with three
33 different commonly-used schemes of C-N coupling. The three schemes (SM1, SM2, and SM3)
34 have been used in three different coupled C-N models (i.e., TECO-CN 2.0, CLM 4.5, and O-CN,
35 respectively). They differ mainly in the stoichiometry of C and N in vegetation and soils, plant N
36 uptake strategies, pathways of N import, and the competition between plants and microbes for
37 soil mineral N. We incorporated them into the C-only version of TECO model, and evaluated
38 their impacts on the C cycle with a traceability framework. Our results showed that all of the
39 three C-N schemes resulted in significant reductions in steady-state C storage capacity compared
40 with the C-only version, but the magnitude varied with -23%, -30% and -54% for SM1, SM2,
41 SM3, respectively. The reduced C storage capacity is the combination of decreases in net
42 primary productivity (NPP) by -29%, -15% and -45% with changes of mean C residence time
43 (MRT) by 9%, -17% and -17% for SM1, SM2, and SM3, respectively. The divergent NPP are
44 mainly attributed to the different assumptions on plant N uptake, plant tissue C:N ratio, down-
45 regulation photosynthesis, and biological N fixation. In comparison, the alternative
46 representations of the plant and microbe competition strategy and the plant N uptake, combining
47 with the flexible C:N ratio in vegetation and soils, led to a notable spread MRT. These results
48 highlight that the diverse assumptions on N process representation among different C-N coupled
49 models could cause additional uncertainty to land surface models. Understanding their difference
50 can help us to improve the capability of models to predict future biogeochemical cycles on land.

51 **Keywords:** carbon-nitrogen coupling, traceability analysis, carbon storage capacity, nitrogen
52 limitation, carbon residence time

53



54 1. Introduction

55 The terrestrial ecosystem carbon (C) storage is jointly determined by ecosystem C input (i.e., net
56 primary productivity, NPP) and mean residence time (MRT), which are modulated by the
57 availability of nitrogen (N) for plant and microbial growth (Vitousek et al., 1991; Wieder et al.,
58 2015; Luo et al., 2017). N is an essential component of enzymes, proteins, and secondary
59 metabolites (van Oijen and Levy, 2004). Plant production require N to meet the stoichiometric
60 demands (Cleveland et al., 2013). Although there is abundant N in the atmosphere, it is difficult
61 to make it available for biological systems (Houlton et al., 2008). As a consequence, the
62 biological N availability, which strongly affects C storage in ecosystems, is often highly
63 correlated with key metabolic rates such as photosynthesis (Field and Mooney, 1986; Du et al.,
64 2017) and respiration (Sprugel et al., 1996). N thus plays an important role in governing the C
65 balance and turnover of terrestrial ecosystem (García-Palacios et al., 2013; Shi et al., 2015).

66 Given the importance of N availability on C sink projections (Wang and Houlton 2009,
67 Zaehle et al., 2015, Wieder et al., 2015), N processes are increasingly incorporated into
68 biogeochemical models. The representation of N cycling and their feedback to C cycling in
69 models reflects what has been established in the ecosystem research community. Early C-N
70 coupled models demonstrated that the N availability limits C storage capacity and can lead to
71 growth enhancement when N mineralization increases in many terrestrial ecosystems (i.e.,
72 progressively increasing N limitation) (Melillo et al., 1993; Luo et al., 2004). Evidences from
73 more recently studies have largely confirmed these results and have generated multiple
74 hypotheses for improving C-N coupling models (Zhou et al., 2013; Zaehle et al., 2014; Thomas
75 et al., 2015). They include the plant down-regulation productivity based on N required for cell
76 construction or N availability for plant absorption (Thornton et al., 2009; Gerber et al., 2010),
77 constant or flexible stoichiometry for allocation and tissue (Wang et al., 2001; Shevliakova et al.,
78 2009; Zaehle et al., 2010), competition between plants and microbes for soil nutrients (Zhu et al.,
79 2017), Evapotranspiration- or NPP-driven empirical functions to generate spatial estimates of
80 biological N fixation (BNF) (Wieder et al., 2015), and respiration of excess C to obtain N from
81 environment and/or to prevent the accumulation of C beyond the storage capacity (Zaehle et al.,
82 2010). These knowledge have significantly helped improve our understanding of the terrestrial
83 C-N coupling and are an important basis to develop comprehensive terrestrial process-based



84 models (Thornton et al., 2007; Thomas et al., 2013). However, simulated results of the terrestrial
85 C cycle illustrated considerable spread among models, and much of uncertainty arose from
86 predictions of N effects on C dynamic. The contradictory results were largely from different
87 representations of fundamental N processes (e.g., the degree of flexibility of C:N ratio in
88 vegetation and soils, plant N uptake strategies, pathways of N export, decomposition, and the
89 representations of the competition between plants and microbes for mineral N). Furthermore, the
90 methodology used to derive the C-N coupling schemes among models varies largely, which may
91 be invalid for the model intercomparisons to provide insight into the underlying mechanism of N
92 status for terrestrial C cycle projection.

93 In the past decades, terrestrial models integrated more and more processes to improve model
94 performance. While the more processes incorporated, the more difficult it becomes to understand
95 or evaluate model behavior (Luo et al., 2015). Xia et al (2013) developed a traceability analysis
96 framework that helped improve the comparability of models and data, evaluated impacts of
97 additional model components, facilitated benchmark analyses, model intercomparisons, and data-
98 model fusion, and improved model predictive power. Based on the traceability analysis
99 framework, key traceable elements, including fundamental properties of the terrestrial C cycle
100 and their representation in shared structures among existing models, can be identified and
101 characterized under different sources of variation (e.g., external forcing and uncertainty in
102 processes) compared to the achieved predictive ability. The traceability analysis framework
103 enables diagnosis of where models are clearly lacking predictive ability and evaluation of the
104 relative benefit when more or alternative components are added to the models (Luo et al., 2015).

105 The present study is designed to examine the effects of C-N coupling under different schemes
106 of model representation on ecosystem C storage in the Terrestrial Ecosystem (TECO) model
107 with traceability analysis framework. Three schemes of model representation were conducted
108 mainly based on TECO-CN 2.0 (SM1), CLM 4.5 (SM2), and O-CN (SM3, Table 1). The three
109 C-N schemes differ in degrees of flexibility of C:N ratio in vegetation and soils, plant N uptake
110 strategies, pathways of N import, and the representations of the competition between plants and
111 microbes for soil available N. Based on the forcing data of ambient CO₂ concentration, N
112 deposition and meteorological data (i.e., air temperature, soil temperature, relative humidity,
113 vapour pressure deficit, precipitation, wind speed, photosynthetically active radiation) obtained
114 from Duke Forest during the period of 1996-2007, we conduct three alternative C-N coupling



115 schemes (i.e., SM1, SM2 and SM3) as well as C-only in TECO model framework to compare
116 their effects on the ecosystem C storage capacity using traceability analysis framework. The N-
117 processes sensitivity analysis was carried out to evaluate the variability in estimated ecosystem C
118 storage caused by the process-related parameters at steady state.

119

120 **2. Materials and methods**

121 **2.1 Data sources**

122 The forcing data used in this study were taken from the AmeriFlux site at Duke Forest, located in
123 the Blackwood Division, North Carolina, USA (35.97° N, 79.08° W). The flux tower lies on a
124 15-year-old loblolly pine (*Pinus taeda* L.) plantation. The meteorological forcing data were
125 downloaded from the AmeriFlux database at <http://ameriflux.lbl.gov>, including ambient CO₂
126 concentration ([CO₂]), air temperature at the top canopy (*T_a*), soil temperature (*T_s*),
127 photosynthetically active radiation (*PAR*), relative humidity (*RH*), vapor pressure deficit (*VPD*),
128 precipitation, wind speed [*W_s*], and N deposition. All forcing data sets are available from 1996 to
129 2007. Standing biomass and biomass production data at each plot for plant compartments (i.e.,
130 foliage, fine root and woody biomass, including branches and coarse roots) were taken from
131 McCarthy et al. (2010). The C and N concentration data for each plant compartment based on
132 Finzi et al. (2007) were used to estimate C and N stocks and fluxes. Plant N demand and uptake
133 were calculated from these data following Finzi et al. (2007). The C and N concentrations of
134 litter and SOM were obtained from Lichter et al. (2008).

135

136 **2.2 Model description and C-N schemes**

137 **2.2.1 TECO-CN 2.0**

138 The terrestrial ecosystem C-N coupling model (TECO-CN, version 2.0) used in the present study
139 is a variant of the TECO-Carbon-only version (TECO-C) by incorporating additional key N
140 processes (Figure 1). TECO-C model is a process-based ecosystem model designed to examine
141 critical processes regulating interactive responses of plants and ecosystems to climate change. It
142 has four major components: canopy photosynthesis module, plant growth module, soil water
143 dynamic module, and soil C dynamic module. The canopy photosynthesis and soil water
144 dynamic modules run at hourly time step while the plant growth and soil C dynamic modules run



145 at the daily time step. The detailed description of the TECO-C model can be found in Weng and
146 Luo (2008).

147 The N cycle added to the TECO model for this study is simplified following the structure of
148 Luo & Reynolds (1999), Gerber et al. (2010), and Wang et al. (2010). It has a similar structure to
149 the TECO-C model (Figure 1). There are nine organic N pools and one inorganic soil N pool,
150 including plant, litter and soil N pools. The plant N pools include leaves, wood, roots, and
151 mineral N in plant tissues. The litter and soil N pools include metabolic and structural litter N,
152 fast, slow, and passive soil organic N (SON), and soil mineral N pools. The total plant N demand
153 on each time step is calculated following the NPP allocation to new tissue growth based on their
154 C:N ratios. To meet the demand, the plant N supply is calculated from three parts, including the
155 retranslocated N from senescing tissues, plant uptake from soil mineral N pool, and external N
156 sources from atmospheric deposition and biological N fixation. The N absorbed by roots enters
157 into the mineral N pool in plant tissues, and then is allocated to the remaining plant pools with
158 plant growth. The N in leaves and fine roots is reabsorbed before senescence. Plant litters will
159 enter metabolic or structural pools depending on their C:N ratios.

160 Allocation of assimilated C among the leaves, stems and roots depends on their growth rates,
161 and varies with phenology (Luo et al., 1995; Denison and Loomis, 1989; Shevliakova et al., 2009;
162 Weng and Luo, 2008):

$$163 \quad b_l = \frac{1}{1+c_1+c_2} \quad (1)$$

$$164 \quad b_s = \frac{c_2}{1+c_1+c_2} \quad (2)$$

$$165 \quad b_r = \frac{c_1}{1+c_1+c_2} \quad (3)$$

166 where b_l , b_s and b_r are the partitioning coefficient of newly assimilated C to leaves, stems and
167 roots, respectively. Parameters c_1 and c_2 are calculated as:

$$168 \quad c_1 = \frac{bm_l}{bm_r} * \frac{CN_l^i}{CN_l^0} \quad (4)$$

$$169 \quad c_2 = 0.5 * 250e^3 * SLA * 0.00021 * h^2 \quad (5)$$

170 where bm_l and bm_r are the leaf and root biomass; CN_l^i and CN_l^0 represent the C:N ratio of the
171 leaf pool at 0 and current time step, respectively; SLA is specific leaf area; h is plant height,
172 which is calculated as:

$$173 \quad h = h_{max}(1 - \exp(-h_1 * bmP)) \quad (6)$$



174 where h_{max} is the maximum canopy height; h_1 is an empirical parameter and bmP is plant
175 biomass.

176

177 **2.2.2 C-N coupling schemes**

178 We conducted four experiments including three simulations with their representations of C-N
179 coupling schemes (SM1, SM2 and SM3) and an additional C-only simulation in TECO model
180 framework. The three C-N interaction simulations include one original scheme in TECO-CN2.0
181 model and the other two schemes representing CLM4.5-BGC and O-CN. The three C-N
182 coupling schemes differ in the representation of down-regulation of photosynthesis, the degree of
183 flexibility of C:N ratio in vegetation and soils, plant N uptake strategies, pathways of N import to
184 the plant reserves, and the competition between plants, and microbes for soil mineral N (Table1,
185 Figure 2).

186

187 **SM1 (TECO-CN2.0)**

188 The N down-regulation of photosynthesis in SM1 is determined by the comparison between
189 plant N demand and actual supply of N:

$$190 \quad f_{dreg} = \min\left(\frac{N_{sup}}{N_{demand}}, 1\right) \quad (7)$$

191 where N_{demand} is plant N demand, and N_{sup} is actual supply of N obtained from re-translocated
192 N, plant N uptake, and biological N fixation.

193 The re-translocated N is calculated as:

$$194 \quad N_{retrans} = \sum_{i=leaf,wood,root} r_i \times outC_i / CN_i \quad (8)$$

195 where r_i is the N resorption coefficient and $outC_i$ is the value of C leaving plant pool i in each
196 time step.

197 The plant N uptake from soil mineral N pool is a function of root biomass density ($Root_{total}$, g
198 C m⁻²) and N demand of plants, following McMurtrie *et al.* (2012)

$$199 \quad N_{uptake} = \min(\max(0, N_{demand} - N_{retrans}), f_{U,max} \times SN_{mine} \times \frac{Root_{total}}{Root_{total} + Root_0}) \quad (9)$$

200 where N_{demand} is the N demand of plants; SN_{mine} is the soil mineral N (gN m⁻²); $f_{U,max}$ is the
201 maximum rate of N absorption per step when $Root_{total}$ approaches infinity; and $Root_0$ is a
202 constant of root biomass (g C m⁻²) at which the N-uptake rate is half of the parameter $f_{U,max}$.

203 The biological N fixation is calculated as:



204
$$N_{BNF} = \min(\max(0, N_{demand} - N_{retrans} - N_{uptake}), n_{fix} \times f_{nsc} \times NSC) \quad (10)$$

205 where $n_{fix} = 0.0167$ is the maximum N fixation ratio and f_{nsc} is the nutrient concentration

206 limiting factor. f_{nsc} is calculated as

207
$$f_{nsc} = \begin{cases} 0, & NSC < NSC_{min} \\ \frac{NSC - NSC_{min}}{NSC_{max} - NSC_{min}}, & NSC_{min} < NSC < NSC_{max} \\ 1, & NSC > NSC_{max} \end{cases} \quad (11)$$

208 where NSC_{min} and NSC_{max} are the minimal and maximal sizes of nonstructural C pool,

209 respectively.

210 Two pathways of N loss are modeled. One is gaseous loss and another is leaching. They both

211 are proportional to the availability of soil mineral N (SN_{min}). The equations are:

212
$$N_{gas_loss} = f_{ngas} \times e^{\frac{T_{soil} - 25}{10}} \times SN_{min} \quad (12)$$

213
$$N_{leach} = f_{nleach} \times \frac{V_{runoff}}{h_{depth}} \times SN_{min} \quad (13)$$

214 where $f_{ngas} = 0.001$ and $f_{nleach} = 0.5$, T_{soil} is the soil temperature, V_{runoff} is the value of

215 runoff, and h_{depth} is the soil depth.

216

217 **SM2 (CLM4.5bgc)**

218 The N down-regulation of photosynthesis in SM2 is calculated as:

219
$$f_{dreg} = \frac{CF_{allo} - CF_{avail_alloc}}{CF_{GPP_{pot}}} \quad (14)$$

220 where CF_{allo} is the total flux of allocated C, which is determined by available mineral N.

221 CF_{avail_alloc} is the potential C flux from photosynthesis, which can be allocated to new growth.

222 $CF_{GPP_{pot}}$ is the potential gross primary productivity (GPP) when there no N limitation.

223 The re-translocated N is calculated as:

224
$$N_{retrans} = \min(N_{demand} \times \frac{N_{retrans_{ann}}}{N_{demand_{ann}}}, N_{retrans_{avail}}) \quad (15)$$

225 where $N_{retrans_{ann}}$ is the previous year's annual sum of re-translocated N obtained from

226 senescing tissues, $N_{demand_{ann}}$ is the previous year's annual sum of plant N demand.

227 $N_{retrans_{avail}}$ is the available re-translocated N in senescing tissues, which is calculated by the

228 proportional of senescing tissues.

229 The plant N uptake is described as:



$$N_{uptake} = (N_{demand} - N_{retrans}) \times f_{plant_demand} \quad (16)$$

where f_{plant_demand} is the fraction (from 0 to 1) of the plant N demand, which can be met given the current soil mineral N supply and competition with heterotrophs. f_{plant_demand} is set equal to the fraction of potential immobilization demand (f_{immob_demand}) that is calculated as:

$$f_{plant_demand} = f_{immob_demand} = \frac{SN_{min}}{N_{plant_demand} + N_{immob_demand}} \quad (17)$$

where N_{immob_demand} is the total potential N immobilization demand (i.e., total potential microbial N demand).

The biological N fixation is calculated as:

$$N_{BNF} = 1.8(1 - e^{-0.03 \times NPP_{py}}) / (86400 \times 365) \quad (19)$$

where NPP_{py} is the previous year NPP.

240

241 SM3 (O-CN)

The N downregulation of photosynthesis in SM3 is calculated as:

$$f_{dreg} = a + b \times N_{leaf/LAI} \quad (20)$$

where a and b are empirical constants, and $N_{leaf/LAI}$ is foliage N per unit leaf area.

The re-translocated N is calculated as:

$$N_{retrans} = \sum_{i=leaf,root} \tau_i \times f_{trans,i} \quad (21)$$

where τ is the foliage or roots shed each step. $f_{trans,leaf} = 0.5$ and $f_{trans,root} = 0.2$ are the fractions of N re-translocated when the tissue dying off.

The plant N uptake is calculated as:

$$N_{uptake} = v_{max} \times SN_{min} \times \left(k_{Nmin} + \frac{1}{N_{min} \times K_{Nmin}} \right) \times f(T_{soil}) \times f(NC_{plant}) \times C_{root} \quad (22)$$

where v_{max} is maximum N uptake capacity per unit fine root mass, k_{Nmin} is the rate of N uptake not associated with Michaelis-Menten Kinetics, K_{Nmin} is the half saturation concentration of fine root N uptake. $f(T_{soil})$ is calculated as:

$$f(T_{soil}) = \exp \left(308.56 * \left(\frac{1}{56.02} - \frac{1}{T_{soil} + 46.02} \right) \right) \quad (23)$$

where T_{soil} is soil temperature.



256 C_{root} is fine root mass. $f(NC_{plant})$ is the dependency of N uptake on plant N status, and is
257 calculated as:

$$258 \quad f(NC_{plant}) = \max\left(\frac{NC_{plant} - nc_{leaf,max}}{nc_{leaf,min} - nc_{leaf,max}}, 0\right) \quad (24)$$

259 where $nc_{leaf,min}$ and $nc_{leaf,max}$ are the minimum and maximum foliage N concentrations,
260 respectively. NC_{plant} is taken as the mean N concentration of foliage, fine roots, and labile N
261 pools, representing the active and easily translocatable portion of plant N:

$$262 \quad NC_{plant} = \frac{N_{leaf} + N_{root} + N_{labile}}{C_{leaf} + C_{root} + C_{labile}} \quad (25)$$

263 The biological N fixation is calculated as:

$$264 \quad N_{BNF} = 0.1 \times \max(0.0234 \times 30 \times AET + 0.172, 0) \quad (26)$$

265 where AET is the mean annual evapotranspiration.

266

267 **2.3 Traceability analysis framework**

268 The traceability analysis framework was used to evaluate the variation of the modeled ecosystem
269 C storage capacity under different C-N schemes (Figure S1). According to the traceability
270 analysis framework (Xia et al., 2013), the modeled C storage capacity can be traced to (i) a
271 product of NPP and ecosystem residence time (τ_E). The latter τ_E can be further traced to (ii)
272 baseline C residence time (τ'_E), which is usually preset in a model according to vegetation
273 characteristics and soil types, (iii) N scalar (ξ_N), (iv) environmental scalars (ξ) including
274 temperature (ξ_T) and water (ξ_W) scalars, and (v) the external climate forcing. The framework for
275 decomposing modeled C storage capacity into a few traceable components is built upon a pool-
276 and flux- structure, which is adopted in all of the terrestrial C models. The structure can well be
277 represented by a matrix equation (Luo et al., 2003; Luo and Weng, 2011):

$$278 \quad \frac{dX(t)}{dt} = BU(t) - A\xi CX(t) \quad (27)$$

279 where $X(t) = (X_1(t), X_2(t), \dots, X_8(t))^T$ is an 8×1 vector describing eight C pool sizes in leaf,
280 root, wood, metabolic litter, structural litter, fast, slow, and passive soil organic C, respectively,
281 in the TECO model (Weng and Luo, 2008). $B = (b_1, b_2, b_3, 0, \dots, 0)^T$ represents the partitioning
282 coefficients of the photosynthetically fixed C into different plant pools. $U(t)$ is the input of fixed
283 C via plant photosynthesis. A is an 8×8 matrix representing the C transfer between pools. ξ is an
284 8×8 diagonal matrix of control of plant N status and environmental scalars on C decay rate at



285 each time step. C is an 8×8 diagonal matrix representing the C exit rates from a pool at each
 286 time step.

287 The C storage capacity equals the sum of C in all pools at steady state (X_{SS}), which can be
 288 obtained by making Eqn.(27) equal zero as described in Xia et al. (2013):

$$289 \quad X_{SS} = (A\xi C)^{-1} B U_{SS} \quad (28)$$

290 The vector U_{SS} is the ecosystem C influx at steady state. The partitioning (B vector), transfer
 291 coefficients (A matrix) and exit rates (C matrix) in Eqn.(27) together determine the baseline C
 292 residence time (τ'_E):

$$293 \quad \tau'_E = (AC)^{-1} B \quad (29)$$

294 The baseline C residence time (τ'_E) in Eqn.(29), N scalars (ξ_N) and environmental scalars (ξ_E)
 295 values together determine the C residence time (τ_E):

$$296 \quad \tau_E = \xi^{-1} \tau'_E = (\xi_N \times \xi_E)^{-1} \tau'_E \quad (30)$$

297 Thus, the C storage capacity is jointly determined by the ecosystem residence time (τ_E) and
 298 steady state C influx (U_{SS}):

$$299 \quad X_{SS} = \tau_E U_{SS} \quad (31)$$

300 The environmental scalar is further separated into the temperature (ξ_T) and water (ξ_W) scalar
 301 components, which can be represented as:

$$302 \quad \xi_E = \xi_T \times \xi_W \quad (32)$$

303 The N scalar is given by vector $\xi_N = (\xi_{N1}(t), \xi_{N2}(t), \dots, \xi_{N8}(t))^T$. The component $\xi_{Ni}(t)$
 304 quantifies the changes of N content at each time step compared with initial condition in the plant
 305 pool i . It is calculated as:

$$306 \quad \xi_{Ni} = \exp\left(-\frac{CN_i^0 - CN_i^n}{CN_i^0}\right) \quad (33)$$

307 where CN_i^0 and CN_i^n are the $C:N$ ratio of the pool i at 0 and n time step, respectively.

308

309 **2.4 Model simulations and sensitivity analysis**

310 To obtain the modeled C storage capacity, we spun up the TECO model with the C -only and
 311 three $C-N$ coupling schemes to steady state using the semi-analytical solution method developed
 312 by Xia et al. (2012). Once the simulations are spun up to steady state, C and N fluxes and state
 313 variables as well as the matrix elements A , C , B , and ξ in Eqn.(28) from all time steps in the last
 314 recycle of the climate forcing were saved for traceability analysis.



315 The sensitivities of both NPP and MRT to each main N process in three schemes were
316 calculated as:

$$317 \quad S_i^{NPP}(P) = \frac{NPP_i^+(P) - NPP_i^-(P)}{NPP_i^0} \quad (34)$$

$$318 \quad S_i^{MRT}(P) = \frac{MRT_i^+(P) - MRT_i^-(P)}{MRT_i^0} \quad (35)$$

319 where $S_i^{NPP}(P)$ and $S_i^{MRT}(P)$ ($i = 1, 2, 3$) represent the sensitivities of the NPP and MRT to the
320 N-process P in the scheme i , respectively. $NPP_i^+(P)$ and $NPP_i^-(P)$ are the annual mean values
321 of NPP that were simulated in scheme i based on the value of the N-process P (ie., DRP, PS,
322 PUN, PMC, BNF, RtrN and SS) increasing 50% and decreasing 50%, respectively. $MRT_i^+(P)$
323 and $MRT_i^-(P)$ are the annual mean values of MRTs that were simulated in the same way as NPP
324 and calculated using Eqn.(29) and Eqn.(30). NPP_i^0 and MRT_i^0 are the annual mean values of
325 NPP and MRT at the steady state in the scheme i .

326

327 3. Results

328 3.1 Simulations of C and N dynamics at steady state

329 At the steady state, the dynamics of N fluxes and soil mineral N showed different patterns among
330 three C-N schemes in the TECO model (Fig 3). The simulated soil N mineralization and plant N
331 uptake fluxes in SM2 displayed the largest daily variations (0.0015 and $0.00086 \text{ g N m}^{-2}\text{d}^{-1}$,
332 respectively) and annual mean values (1.26 and $0.23 \text{ g N m}^{-2}\text{yr}^{-1}$, respectively) among three C-N
333 schemes. For the N leaching flux, SM1 showed the largest daily variation ($0.04 \text{ g N m}^{-2}\text{d}^{-1}$) and
334 annual mean value ($0.36 \text{ g N m}^{-2}\text{yr}^{-1}$). However, the biological N fixation (BNF) flux in SM1
335 showed the largest daily variation ($0.028 \text{ g N m}^{-2}\text{d}^{-1}$) but the smallest annual value ($0.04 \text{ g N m}^{-2}\text{yr}^{-1}$)
336 among three C-N schemes. For the N immobilization fluxes, SM3 simulated the largest
337 daily variation ($0.0013 \text{ g N m}^{-2}\text{d}^{-1}$) and SM1 showed the largest annual mean value ($1.15 \text{ g N m}^{-2}\text{yr}^{-1}$).
338 The dynamics of soil mineral N in SM2 and SM3 displayed the similar patterns on the
339 daily and annual dynamics.

340 Compared with the TECO-C model, the three C-N coupling schemes introduced significant
341 signs of N limitation on forest growth at the steady state but with varying strength (Fig 4).
342 Specifically, the three N schemes resulted in significant reductions in GPP (10%, 10% and 12%
343 for SM1, SM2 and SM3, respectively) compared to the C-only TECO model. Similar response
344 patterns were also found on NPP, ecosystem respiration, and heterotrophic respiration. Among



345 the three schemes, SM3 had the strongest effect (45%, 12% and 45% reduction for NPP,
346 ecosystem respiration, and heterotrophic respiration, respectively), SM2 had the weakest effect
347 (15%, 8% and 13%, respectively) and the effect in SM1 was moderate (29%, 10% and 29%,
348 respectively). However, both the SM1 and SM2 schemes increased the autotrophic respiration
349 (R-auto) by 12% and 27%, respectively, and SM2 scheme increased the NEE by 32%. Due to the
350 NSC pool of TECO model, NEE were positive in all the experiments at the steady state (Weng
351 and Luo, 2008). The NPP and plant N uptake (PNU) joint determine the N use efficiency (NUE).
352 The divergent effects of three C-N schemes on NPP and PNU lead to different NUE (Fig. 5).
353 SM1 had the highest NUE (159.1 g C g⁻¹ N), mainly resulting from its lowest PNU. In contrast,
354 SM3 had the lowest NUE (67.3 g C g⁻¹ N) as a result of its smallest NPP.

355

356 **3.2 Simulation of C storage capacity**

357 The ecosystem C storage capacity differed greatly among the three C-N coupling schemes as
358 well as with the C-only version of TECO model (Fig. 6). The C-only version had the largest C
359 storage capacity (19.5 Kg C m⁻²) among the four simulations, resulting from its highest NPP
360 (879.9 g C m⁻² yr⁻¹). The C storage capacity in SM1 (15.1 Kg C m⁻²) was close to that in SM2
361 (13.7 Kg C m⁻²). The SM3 had the lowest C storage capacity (8.9 Kg C m⁻²) among the four
362 simulations as a result of its smallest NPP (483.9 g C m⁻² yr⁻¹) and relative short MRT (18.6
363 years). By comparison with the C-only version, the three C-N schemes all induced different
364 reductions on NPP (-29%, -15% and -45% for SM1, SM2, SM3, respectively) and further
365 reduced their ecosystem C storage capacity. For the MRT, the three C-N schemes exhibited
366 contrasting effects between SM1 (+9%) and another two schemes (i.e., -16.9% in SM2 and -16.7%
367 in SM3) compared with the C-only TECO model.

368

369 **3.3. Ecosystem C residence time**

370 Ecosystem C residence time (τ_E) is collectively determined by baseline residence time, N scalar
371 and environmental scalar as shown in Eqn (30). Specifically, differences in τ_E among three C-N
372 coupling schemes and C-only TECO model are determined by baseline residence time and the
373 effects of N scalar on eight plant C pools (Fig. 7). For example, SM1 had the longest τ_E because
374 the N scalar had very strong control on passive SOM. The baseline residence time were further
375 determined by the C allocation pattern (Fig. 8). Overall, compared with C-only version, the



376 additional N processes enhanced the partitioning coefficient of NPP to roots (33%, 82% and
377 53%, respectively for SM1, SM2 and SM3), while it decreased the partitioning coefficient to
378 wood (-25%, -45% and -34%, respectively). Furthermore, the decreased partitioning coefficient
379 to wood (b2) regulated the variations of the baseline residence time of wood, structural litter,
380 slow and passive SOM. However, the increased partitioning coefficient to roots (b3) determined
381 the variations of the baseline residence time of root and metabolic litter.

382

383 **3.4. Sensitivity of N processes to NPP and MRT**

384 For either NPP or MRT, the N processes had different sensitivities among the three C-N schemes
385 of TECO model (Fig. 9). For NPP, plant C:N ratio had the highest sensitivities in both SM1
386 (0.32) and SM2 (0.53). However, plant N uptake in SM3 had the highest sensitivity (0.87) for
387 NPP. For MRT, competition between plants and microbes, down-regulation photosynthesis and
388 plant C:N had the highest sensitivities in SM1 (0.27), SM2(0.19) and SM3 (0.56), respectively.

389

390 **4. Discussions**

391

392 **4.1 Underlying N processes and plant production**

393 Gross or net primary production (i.e., NPP or GPP) is regulated by the amount of N available for
394 growth through the N demand, which is set by the relative proportion of biomass growth in the
395 different plant components and their C:N stoichiometry (Zaehle et al., 2014; Thomas et al., 2015).
396 The limitation of equilibrium N on plant production reflects the effects from multiple processes
397 of the C-N interaction, mainly including down-regulation of photosynthetic capacity (DRP) by N
398 availability, the ecosystem's balance of N inputs and losses (net ecosystem N exchange, NNE),
399 plant N uptake (PNU), soil N mineralization (SNM), and the C:N stoichiometry of vegetation
400 and soils. However, due to a lack of consensus on the nature of the mechanisms, the
401 representation of these processes varies greatly among models (Zaehle et al., 2014).

402 There are two common alternative assumptions of the DRP that have been implemented in
403 models: (1) the change in photosynthetic capacity is directly associated with the magnitude of
404 plant available N (e.g., SM2), and (2) N limitation is associated with foliage N, which feeds back
405 to limit photosynthetic capacity (e.g., SM1 and SM3). Our results showed that both assumptions
406 had significant limitations with similar effects on GPP (Fig.3). The probable reason is that the



407 TECO model calculates photosynthesis by light availability vs. carboxylation rate based on the
408 Farquhar model (Farquhar et al., 1980). The effects of N stress under TECO framework, either
409 associated with plant available N or associated with foliage N concentration, are estimated
410 according to limiting factors of photosynthetic biochemistry (the maximum rate of carboxylation,
411 V_{cmax} , and the maximum rate of electron transport at saturating irradiance, V_{jmax}).

412 At or near the steady state, NNE is driven by the processes of N input via deposition and
413 fixation and N loss via leaching and volatilization stoichiometry (Zaehle et al., 2014; Thomas et
414 al., 2015). Previous studies have stated that analyzing the steady-state condition is useful to
415 understand N effects because the balance between external N sources and N losses determine
416 whether an ecosystem is N limited (Rastetter et al., 1997; Menge et al., 2009; Thomas et al.,
417 2015). In this study, divergent NPP responses among three schemes might partly result from
418 their different representations of BNF. For example, SM2 and SM3 simulated BNF explicitly,
419 which used modified empirical relationships to calculate BNF based on evapotranspiration (ET)
420 and NPP, respectively. These phenomenological relationships generally captured
421 biogeographical observations of higher rates of BNF in humid environments with high solar
422 radiation (Wieder et al., 2015). However, the highest response of NPP in only ET-driven BNF
423 (i.e., SM3) may illustrate that not only energetic but also C costs of ‘fixing’ atmospheric di-N
424 (N_2) into a biologically usable form (NH_3) broadly affect NPP (Gutschick 1981, Rastetter et al.,
425 2001). This was because SM3 considered C investments in BNF while SM2 did not. On the other
426 hand, SM1 applied a different strategy, which represents BNF as a complement to the plant N
427 uptake in terms of C investment, leading to the highest plant NUE but the lowest response of
428 BNF to NPP. Another driving factor of NNE is the N loss, which depends on the rate of leaching
429 and volatilization. Using the same formulation as proportional to the size of soil mineral N pool
430 among three schemes, the divergent annual mean magnitude of N leaching was more correlated
431 to soil mineral N.

432 The processes of PNU and net N mineralization determine how N moves through the plant-
433 soil system, thereby triggering N limitation on plant growth. However, to our knowledge,
434 exploring those processes exactly in models is limited by inadequate representation of above-
435 and below-ground interactions that control the patterns of N allocation and whole-plant
436 stoichiometry (Zaehle et al., 2014; Thomas et al., 2015). Plant tissue, litter, and SOM are the
437 primary sinks of N in terrestrial ecosystems, while N in these forms is not directly available for



438 PNU, leading to an increase in N demand due to plant growth. On the other hand, these N must
439 turn over to become available for plant growth. Therefore, the time for N to stay in these
440 unavailable pools controls the transactional delay between the incorporation of N into plant
441 unavailable pool and becomes available for plant uptake. In this way, the residence time of N in
442 SOM appears to be an important factor for governing plant growth (see next section). In the
443 present study, SM1 had the highest NUE from the combined effects of PNU based on C
444 investment strategy (as described above) and flexible tissue C:N ratio. N stress increased tissue
445 C:N ratio, leading to a high microbial N immobilization and then a lower net N mineralization,
446 which allowed plant cell construction with a lower N requirement. The inclusion of flexible C:N
447 stoichiometry (i.e., PS&SS) appeared to be an important feature allowing models to capture the
448 ecosystem response to climate variability through adjusting the C:N ratio of nonphotosynthetic
449 tissues or the whole-plant allocation among tissues with different C:N ratios (Zaehle & Friend,
450 2010). However, it is unclear whether those regulatory mechanisms exist in reality. Further
451 modelling approaches need more reliable framework to predict stoichiometric flexibility.

452

453 **4.2 Ecosystem N status and C residence time**

454 Ecosystem N status in models, including plant-available and unavailable N forms, is set by N
455 inputs from N fixation and N deposition, N losses from leaching and denitrification, and N gain
456 from the turnover of litter and SOM through tissue senescence and decomposition. As noted
457 above, external N cycle (i.e., N inputs and N losses) couples the N processes within the plant-
458 litter-SOM system, being mainly associated with the limitation of plant production. The effects
459 of ecosystem N status on C mean residence time (MRT), however, has been much less studied
460 than N limitation on productivity of plants and soil organisms, largely because these effects
461 involve various impacts on C transfer among pools and release from each pool via
462 decomposition and respiration (Thompson & Randerson, 1999; Xia et al., 2013). Therefore, the
463 different impacts of ecosystem N status induce oscillating N limitation on MRT due to the
464 inherently different assumptions of C-N interactions among three C-N coupling schemes.

465 At the steady state, the different effects of N status on changes in modelled MRT can be
466 attributed to: the different rate of soil N mineralization dependent on the total amount of N in
467 SOM and its turnover time, immobilization based on the competition strategy between plants and
468 microbes and their stoichiometry, and different deployment of reabsorbed N. The traceability



469 framework in this study can trace those different effects into three components (i.e., climate
470 forcing, N scalar and baseline MRT) based on three alternative C-N coupling schemes under the
471 TECO model framework. Since the forcing data are identical, we assumed the same effects for
472 this component in all four experiments, which is thus not discussed further in this section.

473 In our study, the N scalar was based on the dynamics of C:N ratios (Eqn. 33). Therefore, N
474 scalar had no effect on MRT in SM2, resulting from the assumption of fixed C:N ratio in all C
475 pools (Fig. 6c). In both SM1 and SM3, however, the N scalar had large effects on the SOM pool,
476 which is probably related to different mechanisms. Specifically, N scalar in the SM1 had the
477 contrasting effects on MRT of fast and passive SOM pools (i.e., negative vs. positive,
478 respectively), which may largely be attributed to the plant and microbe competition strategy
479 combining with a much larger passive SOM pool in TECO-CN2.0 model (Du et al., 2017; Zhu et
480 al., 2017). Under N stress, the competition between plants and microbes is expected to be
481 intensified, resulting in increasing C:N ratio of nonphotosynthetic tissues (e.g., wood and root)
482 and the total C:N ratio. This effectively prevents N limitation of cell construction and
483 corresponds to an increase in whole-plant NUE (Thomas et al., 2015). In this case, higher C:N
484 ratio in those tissues lowers structural litter quality, leading to soil microbes to immobilize more
485 N to maintain their stoichiometric balance (Hu et al., 2001; Manzoni et al., 2010). However, in
486 the SM3, increased respiration acted as a mechanism to remove the excess accumulated C, which
487 is a stoichiometry-based implementation to prevent the accumulation of labile C under N stress
488 (Zaehle & Friend, 2010; Thomas et al., 2015). This mechanism promotes absorption and
489 respiration of the faster turnover pools (fast and slow SOM pools), leading to decrease in MRT
490 in these two pools.

491 In the traceability framework, the baseline MRT is determined by the potential decomposition
492 rates of C pools (C matrix), coefficients of C partitioning of NPP (B vector), and transfer
493 coefficients between C pools (A matrix, Eqn. [29]. Xia et al., 2013). The matrices A and C are
494 preset in the TECO model according to vegetation characteristics and soil textures (Weng and
495 Luo., 2008). Therefore, the notable spread in baseline MRT across the C-N schemes was induced
496 by the B vector, which was modified by different N-limitation assumptions. Conceptually, in
497 order to meet the N demand, plants adjust NPP allocation to N absorption tissues (e.g., roots). In
498 this study, three schemes all had similar trends of adjusting allocation C from wood to roots (Fig.
499 7), but with different mechanisms. For both SM1 and SM3, increased root C allocation was



500 mainly driven by N uptake capacity, which is associated with plant competitiveness in SM1 and
501 the respiration of excess labile C in SM3, respectively. However, for SM2, increasing root C
502 allocation may occur in spin-up stage from plant adjustment to whole-plant allocation among
503 tissues to fit fixed C:N ratio.

504

505 **5. Conclusions**

506 The C-N coupling has been represented in ecosystem and land surface models with different
507 schemes, generating great uncertainties in model predictions. The most striking difference
508 among terrestrial C-N coupling models occurs with the degree of flexibility of C:N ratio in
509 vegetation and soils, plant N uptake strategies, pathways of N import, and the representations of
510 the competition between plants and microbes for soil mineral N. In this study, we evaluated
511 alternative representations of C-N interactions and their impacts on C cycle using the TECO
512 model framework. Our traceability analysis showed that different representations of C-N
513 coupling processes lead to divergent effects on both plant production and C residence time, and
514 thus the ecosystem C storage capacity. The plant production are mainly affected by the different
515 assumptions on net ecosystem N exchange, plant N uptake, net N mineralization, and the C:N
516 ratio of vegetation and soil. In comparison, the alternative representations of the plant and
517 microbe competition strategy, combining with the flexible C:N ratio in vegetation and soils, led
518 to a notable spread effects on C residence time. Identifying the representations of main C-N
519 processes under different schemes can help us to improve the N-limitation assumptions
520 employed in terrestrial ecosystem models and forecasting future C sink dynamic in response to
521 climate change.

522

523 *Code availability.* The code for TECO-CN2.0 and the three C-N coupling schemes is available at
524 <https://github.com/zgdu/TECO-CN-2.0-new>.

525 *Data availability.* The data for this paper are available upon request to the corresponding author.

526 *Competing interests.* The authors declare that they have no conflict of interest.

527



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534

535



536 Figure legends

537 **Figure 1.** Schematic diagram of the terrestrial ecosystem carbon (C) and nitrogen (N) coupling
538 model (TECO-CN2.0). (A) Canopy module, (B) Plant growth module, (C) Soil water dynamics
539 module, (D) Soil carbon-nitrogen coupling module. Rectangles represent the carbon and nitrogen
540 pools. R_a , autotrophic respiration. R_h , heterotrophic respiration. Retr., re-translocation. NSC,
541 nonstructural carbohydrate. MNP, mineral N in plant tissues. SOM, soil organic matter.

542

543 **Figure 2.** Schematic diagram illustrating the major carbon (C) and nitrogen (N) flows and stores
544 in a terrestrial ecosystem. Light-blue arrows indicate C-cycle processes and red arrows show N-
545 cycle processes.^{1,2,3} alternative assumptions of N processes represent in scheme 1, 2 and 3,
546 respectively. Met./Str. Litter, metabolic and/or structural litters; SOM, soil organic matter.

547 **Figure 3.** Simulated nitrogen fluxes and soil mineral nitrogen from three carbon-nitrogen
548 coupling schemes (SM1, SM2 and SM3) in TECO-CN model for 1996 to 2007 at Duke Forest.

549 **Figure 4.** Simulated annual (a-f) and mean (g-l) carbon fluxes from carbon-only version and
550 carbon-nitrogen coupled with three schemes (SM1, SM2 and SM3) of TECO model for 1996 to
551 2007 at Duke Forest. GPP, gross primary productivity; NPP, net primary productivity; NEE, net
552 ecosystem exchange of CO_2 ; R-eco, ecosystem respiration; R-heter, heterotrophic respiration; R-
553 auto, autotrophic respiration.

554 **Figure 5.** The nitrogen use efficiency (NUE) in three C-N schemes of TECO model (SM1, SM2
555 and SM3).

556 **Figure 6.** Simulation of annual ecosystem carbon storage capacity for 1996 to 2006 at Duke
557 Forest by carbon in flux (NPP, x axis) and ecosystem residence time (τ_E , y axis) in TECO model
558 framework with three carbon-nitrogen coupling schemes (SM1, SM2 and SM3) and in TECO C-
559 only model (C). Inset (a), ecosystem carbon residence time (τ_E) in SM1, SM2, SM3 and C-only
560 model; inset (b), mean ecosystem carbon storage simulated among SM1, SM2, SM3 and C-only
561 model; inset (c), relative change of NPP and ecosystem residence time simulated among three
562 schemes compared with in C-only model.

563 **Figure 7.** Determination of carbon-pool residence times based on traceability framework in
564 TECO C-N model with three C-N coupling schemes (SM1, SM2 and SM3) and TECO C-only
565 model (C). Panel (a), baseline residence time; panel (b), mean residence time, and panel (c),
566 nitrogen scalar.

567 **Figure 8.** Coefficients of partitioning of NPP to nonstructural C (NSC), root, woody and leaf in
568 C-only model (C) and C-N coupling model with three schemes (SM1, SM2 and SM3).

569 **Figure 9.** The sensitivity of nitrogen processes to NPP (panel a) and ecosystem residence time
570 (τ_E , panel b) among three carbon-nitrogen coupling schemes (SM1, SM2 and SM3). DRP, down-
571 regulation photosynthesis; PS, plant tissue C:N ratio; PNU, plant N uptake; PMC: plant and
572 microbe competition; BNF, biological N fixation; RtrN, re-tranlocation N; SS, soil pool C:N
573 ratio.



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697 resolves inconsistencies between observations and model predictions. *Ecological Applications*,
698 27(3), 875-886.
- 699



700 **Table1.** Summary of the nitrogen-carbon coupling schemes used and the representation of key
 701 processes in the carbon-nitrogen cycle.

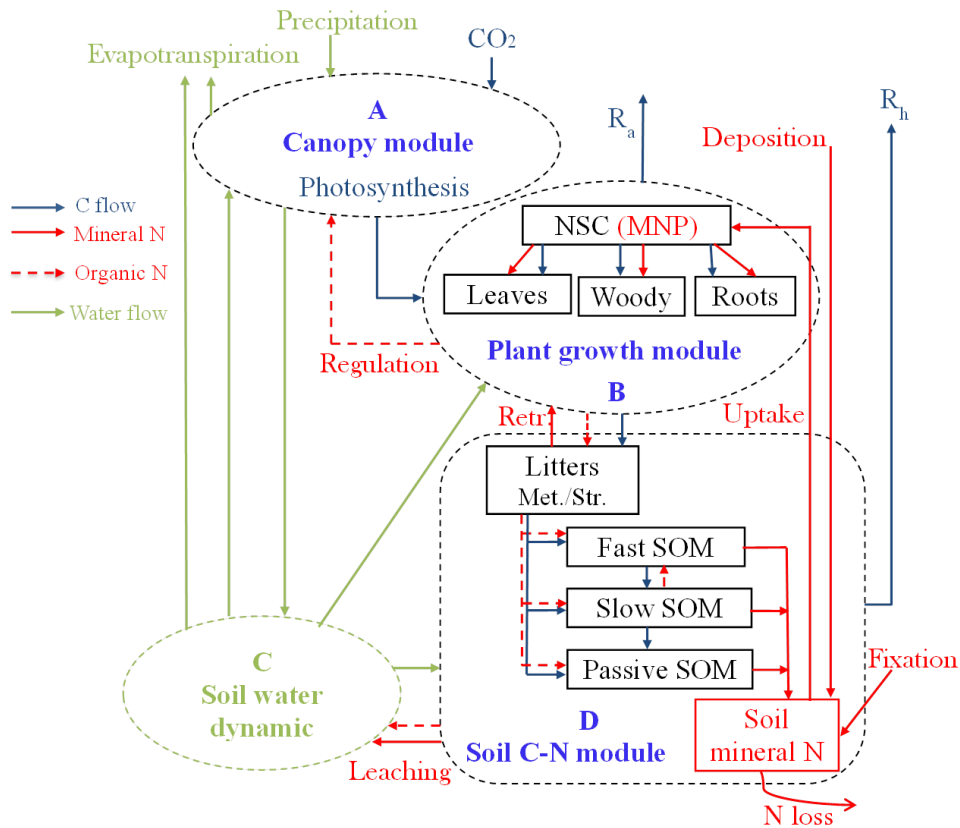
| | SM1 (TECO-CN2.0)^a | SM2 (CLM4.5)^{b,c} | SM3 (O-CN)^{d,e} |
|---|--|--|--|
| Photosynthesis down regulation by N availability (DRP) | Based on the comparison between plant N demand and actual supply | Based on the available soil mineral N relative to the N demanded to allocate photosynthate to tissue | Based on foliage N concentration, which varies with N deficiency |
| Plant tissue stoichiometry (PS) | Flexible plant C:N ratio | Fixed plant C:N ratio | Flexible plant C:N ratio |
| Plant N uptake (PNU) | Based on fine root biomass, soil mineral N and N demand of plant. Plants itself choose the strategy between uptake from soil mineral N and fix N ₂ by comparing C investment | Based on N required to allocate NPP to tissue. Plants uptake N for free | Combining active and passive uptake of mineral N based on fine root C, soil mineral N, plant transpiration flux, increases with increased plant N demand |
| N competition between plants and microbes (PMC) | Microbes have first access to soil mineral N | Based on demand by both microbial immobilization and plant N uptake | Microbes have first access to soil mineral N, the competitive strength of plants increases under nutrient stress |
| Biological N fixation (BNF) | Based on the nitrogen demand of plants and maximum N fixing ratio considering nutrient concentration | $f(NPP)$ | $f(ET)$ |
| Deployment of re-translocated N (RtrN) | Fixed fraction of litter | Based on available N in the tissue and the previous year's annual sum of plant N demand | Fixed fraction of dying leaf and root tissue |
| Soil organic matter stoichiometry (SS) | Flexible soil C:N ratio | Fixed soil C:N ratio | Flexible soil C:N ratio |
| N leaching | Function of soil mineral N pool and runoff | Function of soil mineral N pool and runoff | Function of soil mineral N and runoff |

702 ^aSee this study; ^bThornton et al. (2007), ^cThornton et al. (2009); ^dZaehle & Friend (2010),
 703 ^eZaehle et al. (2011).

704 C, carbon; N, nitrogen; NPP, net primary productivity; ET, evapotranspiration.



705 **Figure 1. TECO-CN 2.0**



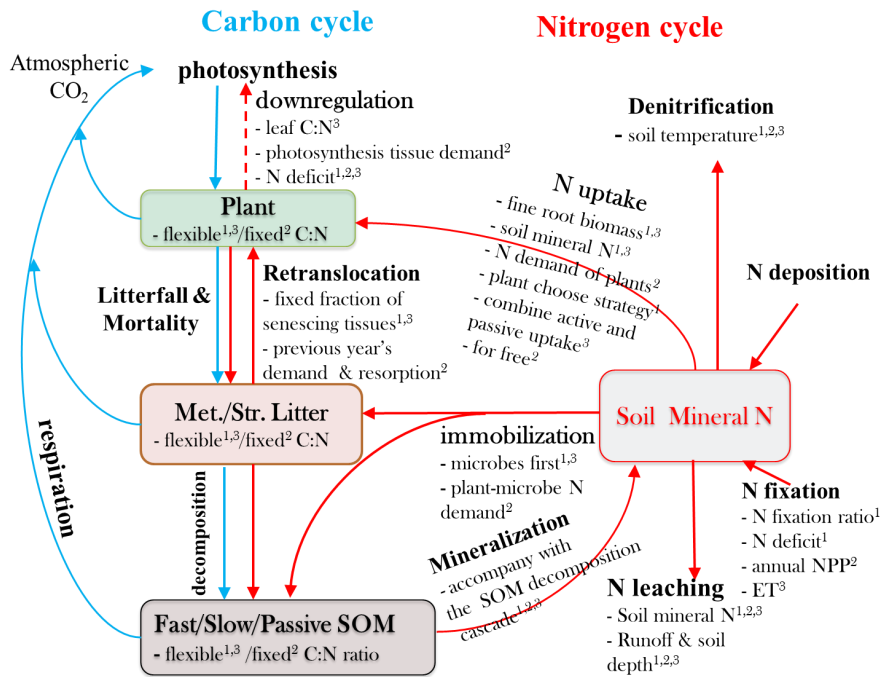
706

707 **Figure 1.** Schematic diagram of the terrestrial ecosystem carbon (C) and nitrogen (N) coupling
 708 model (TECO-CN2.0). (A) Canopy module, (B) Plant growth module, (C) Soil water dynamics
 709 module, (D) Soil carbon-nitrogen coupling module. Rectangles represent the carbon and nitrogen
 710 pools. R_a , autotrophic respiration. R_h , heterotrophic respiration. Retr., re-translocation. NSC,
 711 nonstructural carbohydrate. MNP, mineral N in plant tissues. SOM, soil organic matter.

712



713 **Figure 2**



714

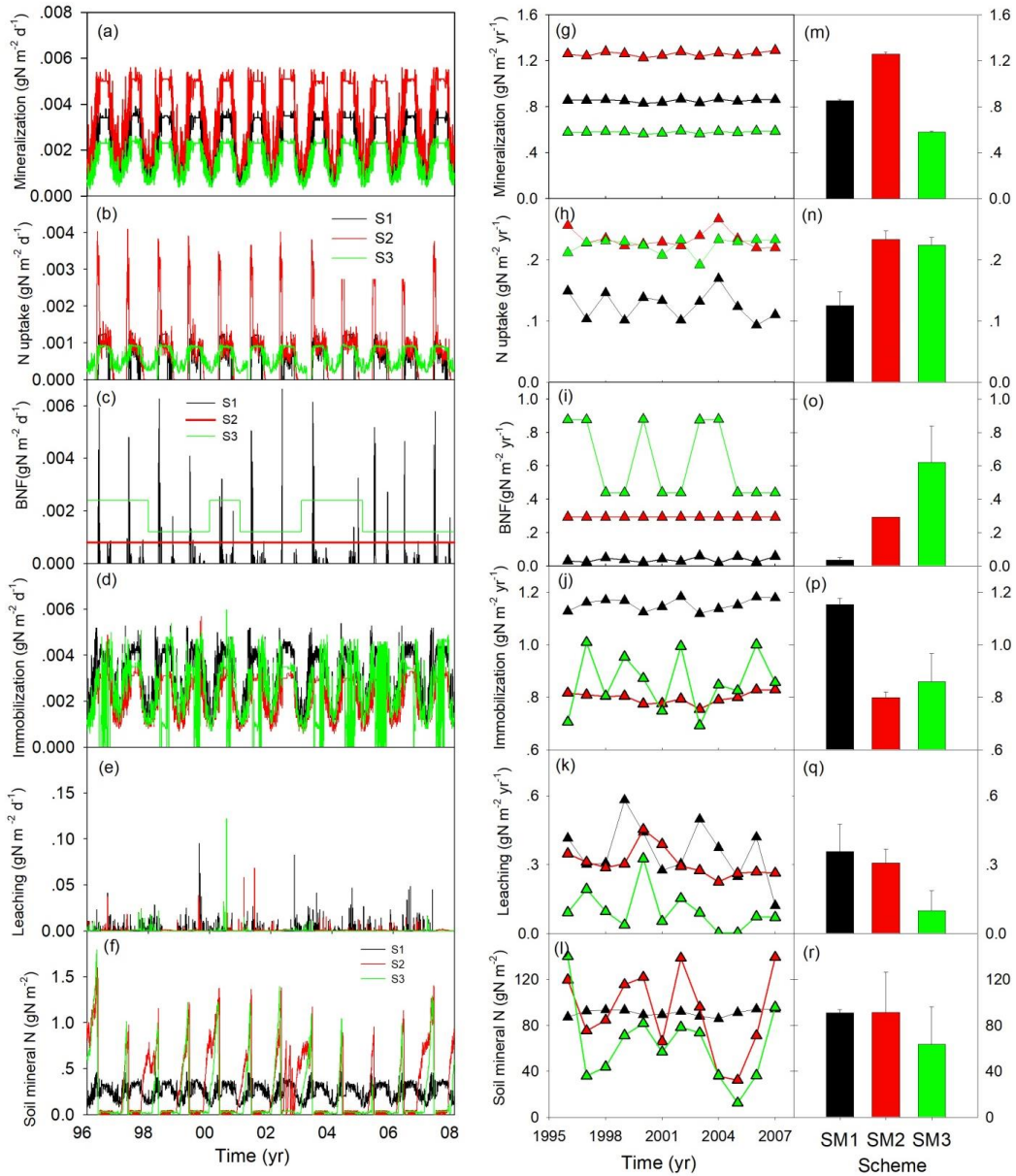
715

716 **Figure 2.** Schematic diagram illustrating the major carbon (C) and nitrogen (N) flows and stores
 717 in a terrestrial ecosystem. Light-blue arrows indicate C-cycle processes and red arrows show N-
 718 cycle processes. ^{1,2,3} alternative assumptions of N processes represent in scheme 1, 2 and 3,
 719 respectively. Met./Str. Litter, metabolic and/or structural litters; SOM, soil organic matter.

720



721 **Figure 3**



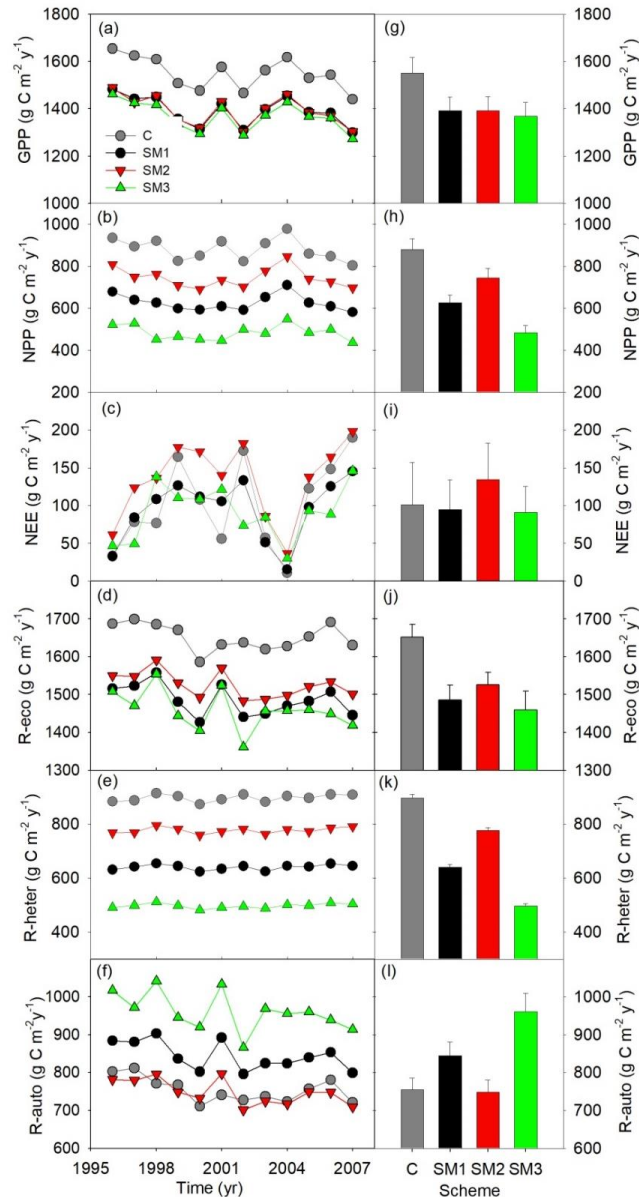
722

723 **Figure 3.** Simulated nitrogen fluxes and soil mineral nitrogen from three carbon-nitrogen
 724 coupling schemes (SM1, SM2 and SM3) in TECO-CN model for 1996 to 2007 at Duke Forest.

725



726 **Figure 4**

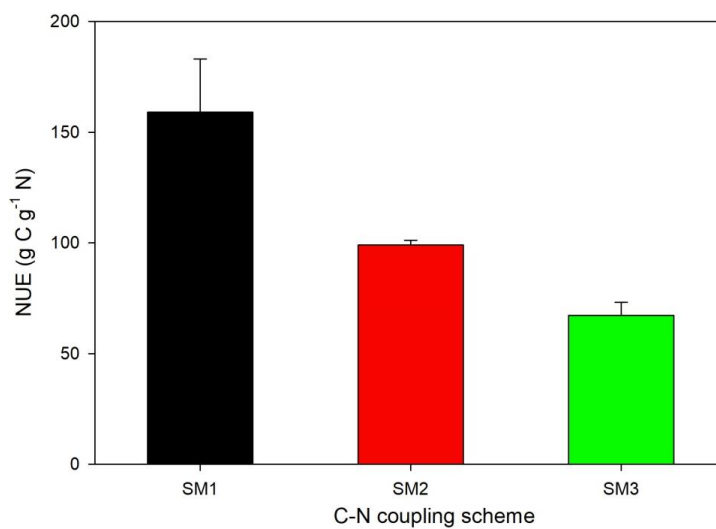


727

728 **Figure 4.** Simulated annual (a-f) and mean (g-l) carbon fluxes from carbon-only version and
 729 carbon-nitrogen coupled with three schemes (SM1, SM2 and SM3) of TECO model for 1996 to
 730 2007 at Duke Forest. GPP, gross primary productivity; NPP, net primary productivity; NEE, net
 731 ecosystem exchange of CO₂; R-eco, ecosystem respiration; R-heter, heterotrophic respiration; R-
 732 auto, autotrophic respiration.



733 **Figure 5**



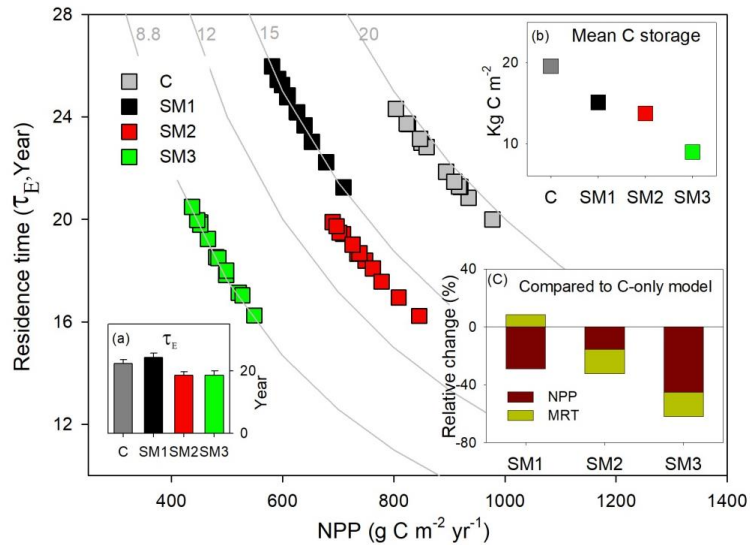
734

735 **Figure 5.** The nitrogen use efficiency (NUE) in three C-N schemes of TECO model (SM1, SM2
736 and SM3).

737



738 **Figure 6**



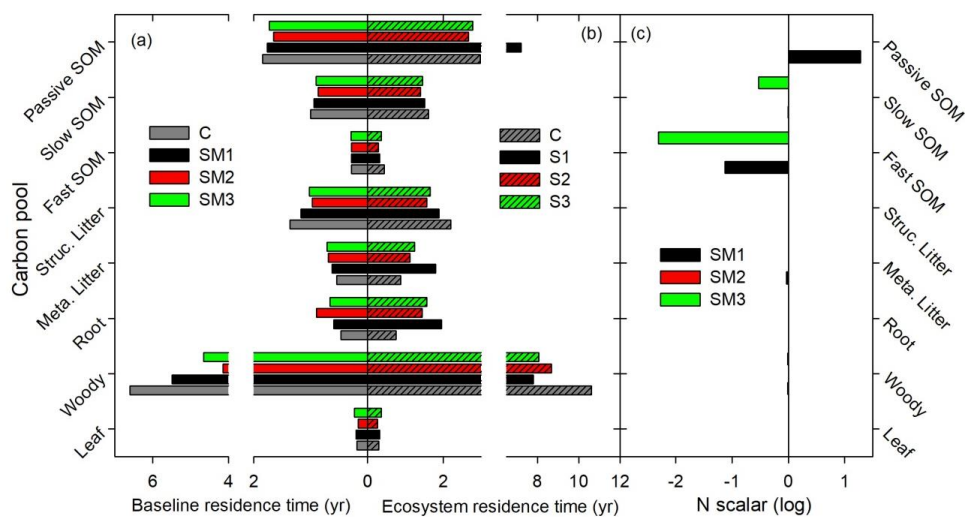
739

740 **Figure 6.** Simulation of annual ecosystem carbon storage capacity for 1996 to 2006 at Duke
 741 Forest by carbon in flux (NPP, x axis) and ecosystem residence time (τ_E , y axis) in TECO model
 742 framework with three carbon-nitrogen coupling schemes (SM1, SM2 and SM3) and in TECO C-
 743 only model (C). The hyperbolic curves represent constant values (shown across the curves) of
 744 ecosystem carbon storage capacity. Inset (a), ecosystem carbon residence time (τ_E) in SM1, SM2,
 745 SM3 and C-only model; inset (b), mean ecosystem carbon storage simulated among SM1, SM2,
 746 SM3 and C-only model; inset (c), relative change of NPP and ecosystem residence time
 747 simulated among three schemes compared with in C-only model.

748



749 **Figure 7**



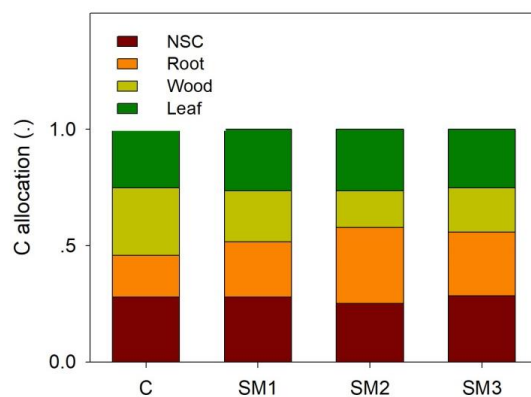
750

751 **Figure 7.** Determination of carbon-pool residence times based on traceability framework in
 752 TECO C-N model with three C-N coupling schemes (SM1, SM2 and SM3) and TECO C-only
 753 model (C). Panel (a), baseline residence time; panel (b), mean residence time, and panel (c),
 754 nitrogen scalar.

755



756 **Figure 8**



757

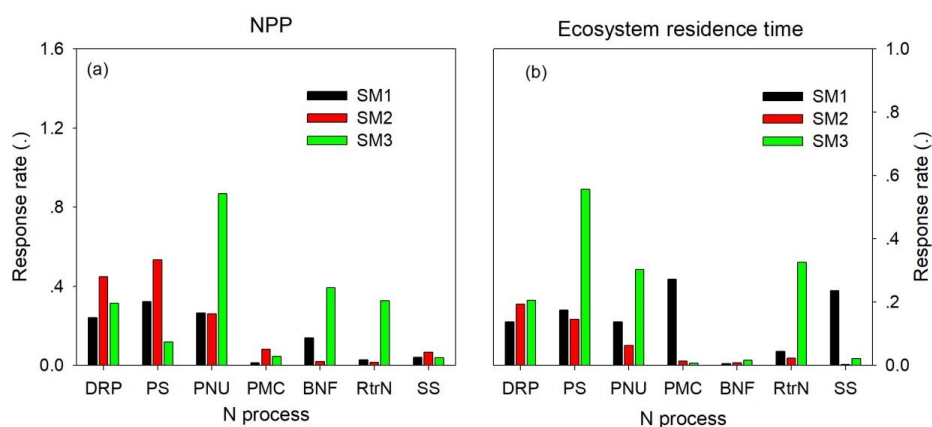
758 **Figure 8.** Coefficients of partitioning of NPP to nonstructural C (NSC), root, woody and leaf in
759 C-only model (C) and C-N coupling model with three schemes (SM1, SM2 and SM3).

760

761



762 **Figure 9**



763

764 **Figure 9.** The sensitivity of nitrogen processes to NPP (panel a) and ecosystem residence time
 765 (τ_E , panel b) among three carbon-nitrogen coupling schemes (SM1, SM2 and SM3). DRP, down-
 766 regulation photosynthesis; PS, plant tissue C:N ratio; PNU, plant N uptake; PMC: plant and
 767 microbe competition; BNF, biological N fixation; RtrN, re-tranlocation N; SS, soil pool C:N
 768 ratio.

769