



1 **Optimizing a dynamic fossil fuel CO₂ emission model with**
2 **CTDAS (v1.0) for an urban area using atmospheric**
3 **observations of CO₂, CO, NO_x, and SO₂**

4 Ingrid Super^{1,2}, Hugo A.C. Denier van der Gon¹, Michiel K. van der Molen², Stijn N.C.
5 Dellaert¹, Wouter Peters^{2,3}

6 ¹Department of Climate, Air and Sustainability, TNO, P.O. Box 80015, 3508 TA Utrecht, Netherlands

7 ²Meteorology and Air Quality Group, Wageningen University, P.O. Box 47, 6700 AA Wageningen, Netherlands

8 ³Centre for Isotope Research, Energy and Sustainability Research Institute Groningen, University of Groningen,
9 Nijenborgh 4, 9747 AG Groningen, Netherlands

10 *Correspondence to:* Ingrid Super (ingrid.super@tno.nl)



11 **Abstract.** We present a modelling framework for fossil fuel CO₂ emissions in an urban environment, which allows
12 constraints from emission inventories to be combined with atmospheric observations of CO₂ and its co-emitted
13 species CO, NO_x, and SO₂. Rather than a static assignment of average emission rates to each unit-area of the urban
14 domain, the fossil fuel emissions we use are dynamic: they vary in time and space in relation to data that describe
15 or approximate the activity within a sector, such as traffic density, power demand, 2m temperature (as proxy for
16 heating demand), and sunlight and wind speed (as proxies for renewable energy supply). Through inverse
17 modelling, we optimize the relationships between these activity data and the resulting emissions of all species
18 within the dynamic fossil fuel emission model, based on atmospheric mole fraction observations. The advantage
19 of this novel approach is that the optimized parameters (emission factors and emission ratios, N=44) in this
20 dynamic model (a) vary much less over space and time, (b) allow a physical interpretation of mean and uncertainty,
21 and (c) have better defined uncertainties and covariance structure. This makes them more suited to extrapolate,
22 optimize, and interpret than the gridded emissions themselves. The merits of this approach are investigated using
23 a pseudo-observation-based ensemble Kalman filter inversion setup for the Dutch Rijnmond area at 1x1 km
24 resolution.

25 We find that the dynamic fossil fuel model approximates the gridded emissions well (annual mean differences <
26 2 %, hourly temporal $r^2 = 0.21-0.95$), while reported errors on the underlying parameters allow a full covariance
27 structure to be created readily. Propagating this error structure into atmospheric mole fractions shows a strong
28 dominance of a few large sectors and a few dominant uncertainties, most notably the emission ratios of the various
29 gases considered. If these are either sufficiently well-known a-priori, or well-constrained from a dense observation
30 network, we find that including observations of co-emitted species improves our ability to estimate emissions per
31 sector relative to using CO₂ mole fractions only. Nevertheless, the total CO₂ emissions can be well-constrained
32 with CO₂ as only tracer in the inversion. Because some sectors are sampled only sparsely over a day, we find that
33 propagating solutions from day-to-day leads to largest uncertainty reduction and smallest CO₂ residuals over the
34 14 consecutive days considered. Although we can technically estimate the temporal distribution of some emission
35 categories like shipping separate from their total magnitude, the controlling parameters are difficult to distinguish.
36 Overall, we conclude that our new system looks promising for application in verification studies, provided that
37 reliable urban atmospheric transport fields and reasonable a-priori emission ratios for CO₂ and its co-emitted
38 species can be produced.



39 1 Introduction

40 Within the 2015 Paris Agreement, 195 nations agreed with a climate action plan in which each nation sets its own
41 targets for carbon emission reductions and reports all efforts regularly to the UNFCCC (UNFCCC, 2015). An
42 important role in reaching emission reduction targets is laid out for cities, which emit a large portion of the global
43 fossil fuel CO₂ emissions (about 70 % according to the International Energy Agency (IEA, 2008)). The Paris
44 Agreement also states that parties should strengthen their cooperation, also with respect to the sharing of
45 information and good practices. Within this context it becomes increasingly important to map fossil fuel emissions
46 and to quantify emission trends, both at the country- and city-level.

47 Most country-level greenhouse gas emission estimates reported to the UNFCCC are currently based on yearly fuel
48 consumption data (bottom-up method), and are often spatiotemporally disaggregated using activity data and
49 proxies to create spatially explicit emission inventories (Kuenen et al., 2014; Hutchins et al., 2017). Although the
50 yearly national estimates are reasonably accurate (estimated uncertainty for developed countries is less than 8 %
51 for CO₂ (Monni et al., 2004; Fauser et al., 2011; Andres et al., 2014)), their uncertainty increases rapidly when
52 disaggregating them towards finer spatiotemporal resolutions (Ciais et al., 2010; Nassar et al., 2013; Andres et al.,
53 2016). A method to improve emission estimates is using transport models in combination with independent
54 observations of atmospheric mole fractions (Palmer et al., 2018), called data assimilation (DA) or inverse
55 modelling (a top-down method). Recently, efforts have been made to apply DA techniques to the urban
56 environment (McKain et al., 2012; Brioude et al., 2013; Lauvaux et al., 2013; Bréon et al., 2015; Boon et al., 2016;
57 Lauvaux et al., 2016; Stauffer et al., 2016; Brophy et al., 2018), but several challenges and unexploited opportunities
58 remain.

59 First, urban DA studies have tried to constrain the total fossil fuel flux to validate bottom-up CO₂ inventories, often
60 without considering the underlying emission process that caused the mismatch between observed and modelled
61 concentrations. As one of very few exceptions, Lauvaux et al. (2013) used the CO:CO₂ concentration ratio to
62 conclude that the emission reduction in Davos during the World Economic Forum 2012 was likely related to
63 reduced traffic emissions, but without a quantification. However, emission reduction policies usually target
64 specific source sectors. Therefore, an increase in fossil fuel emissions from one source sector can cause the total
65 CO₂ emissions to appear stable, although a policy targeting another source sector can be effective in itself. To
66 monitor the effect of each measure independently it becomes essential to attribute changes in the total CO₂
67 emissions to these policies and thus to specific source sectors. It is, therefore, not sufficient to constrain the total
68 CO₂ flux, but we need to differentiate the total CO₂ signal into signals from the different source sectors. One way
69 to accomplish this is by using additional measurements of co-emitted species and isotopes. Such measurements
70 have previously been used in modelling studies to differentiate between biogenic and anthropogenic emissions or
71 between fuel types (Djuricin et al., 2010; LaFranchi et al., 2013; Lopez et al., 2013; Turnbull et al., 2015; Fischer
72 et al., 2017; Super et al., 2017b; Brophy et al., 2018; Graven et al., 2018), but also to separate between different
73 fossil fuel sources (Lindenmaier et al., 2014; Super et al., 2017a; Nathan et al., 2018).

74 Second, for urban DA the fine scales (less than 1km and less than 1 hour) need to be resolved, therefore putting a
75 higher demand on the atmospheric transport models. For example, Boon et al. (2016) mentioned that sources with
76 a small spatial extent (point sources) are not correctly represented on a 2x2 km² grid, while these sources have a
77 significant impact on the locally observed mole fractions. Concurrently, we have previously shown that a plume
78 model improves the representation of sources with a limited spatial extent. Moreover, we found that the description



79 of short-term variations in the wind direction by the Eulerian WRF model in the vicinity of an urban area is poor
80 (Super et al., 2017a).

81 Third, the prior emissions also need to have a higher resolution for urban-scale studies to resolve the dominant
82 spatiotemporal variations. Previous studies have often used high-resolution emission maps developed specifically
83 for that region, using local data as much as possible (Zhou and Gurney, 2011; Bréon et al., 2015; Boon et al., 2016;
84 Lauvaux et al., 2016; Rao et al., 2017; Gurney et al., 2019). Yet such emission maps are only available for a few
85 data-rich regions. For other regions, continental or global emission maps (such as MACC or EDGAR) can be used
86 if downscaling is applied to reach the high resolution required for urban-scale inversions. For example, the
87 temporal downscaling can be done using typical daily, weekly and monthly profiles for each source sector (Denier
88 van der Gon et al., 2011), which are based on activity data (e.g. traffic counts) averaged over several years and/or
89 a large region. Spatial downscaling often involves proxies like population density. This spatiotemporal
90 downscaling introduces a large additional uncertainty due to uncertainties in the proxies. For example, Hogue et
91 al. (2016) have found an uncertainty of 150 % in the $1 \times 1^\circ$ fossil fuel CO₂ emissions for the US, whereas Ciais et
92 al. (2010) estimated the uncertainty of regional European emissions at 100 km resolution to be about 50 %.
93 Quantification of the uncertainty at an even higher resolution for urban applications has so far been limited (Andres
94 et al., 2016) (Super et al., 2019), also for most local inventories, while a correct definition of the prior error
95 covariance matrix for an inversion is important to get reliable output (Chevallier et al., 2006; Boschetti et al.,
96 2018). This currently complicates the application of DA studies to urban areas.

97 Here, we describe the development of an urban-scale DA framework (based on the CarbonTracker Data
98 Assimilation Shell (CTDAS)) which uses a dynamic fossil fuel emission model as a prior and optimizes the
99 parameters of this model. The dynamic fossil fuel emission model uses a wide range of (statistical) data to calculate
100 CO₂ emissions per source sector at high spatiotemporal resolution (1×1 km² and hourly). The emission model is
101 dynamic in the sense that its formulation allows emissions to change as a function of rapidly varying conditions
102 in the emission landscape, such as the outside temperature, the traffic density, or availability of wind and solar
103 radiation for sustainable power generation. Using such information enables the calculation of dynamic emissions
104 in near real-time, as opposed to the construction of a static emission map based on statistical downscaling.
105 Moreover, the emission model can supply spatiotemporal emission uncertainties and error correlations between
106 source sectors, based on the estimated uncertainty of its model parameters. Since many of these parameters are
107 also used in bottom-up accounting of emissions, their uncertainty is often better established than the uncertainty
108 in the total emissions themselves. Finally, we use the dynamic emission model to calculate emissions of other co-
109 emitted species (CO, NO_x and SO₂) from the CO₂ emissions using source sector specific emission ratios. These
110 co-emitted species are included in the DA system to facilitate source attribution, which is possible due to the
111 distinct emission ratios of different source sectors. The overall aim of this study is to explore how our dynamic
112 fossil fuel emission model and additional tracers can be used to overcome the known limitations in anthropogenic
113 CO₂ inverse modelling described above. The research questions are:

- 114 1. Can our dynamic fossil fuel emission model represent the spatiotemporal structure of a high-resolution
115 emission inventory, and what does it add to that on small scales?
- 116 2. Is the addition of co-emitted species beneficial for the attribution of CO₂ signals to specific source sectors,
117 and which observations help most in that effort?

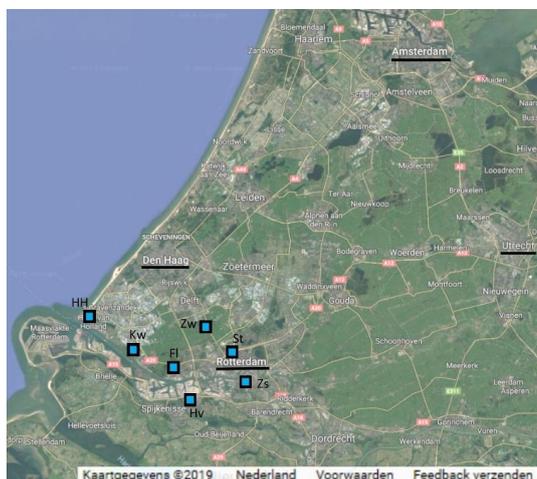


118 3. Does the prior error covariance structure that we build with the dynamic emissions model help the
119 optimization, and what can we learn from the posterior error covariance estimate?

120 In the inverse part of this study we use observing system simulation experiments (OSSEs, experiments using
121 pseudo-observations), applied to the urban-industrial complex of Rotterdam (Netherlands). This choice allows us
122 to test our new approach, while with real observations the errors in non-fossil and background fluxes, model
123 structure, and model transport will likely dominate the results (Tolk et al., 2008; Super et al., 2017a; He et al.,
124 2018) and reduce the ability to evaluate the methodology. First, we give an overview of the dynamic fossil fuel
125 emission model and demonstrate its applicability to the domain, followed by an introduction to the DA system
126 components and the model settings. Then we discuss the different experiments in which we start with the
127 comparison of different network configurations, one with only CO₂ and one including co-emitted species to
128 examine the ability to attribute CO₂ emissions to specific source sectors, and different state vectors. Another
129 experiment is used to examine the importance of propagating posterior parameters values and covariances. Finally,
130 we address the effect of cross-correlations.

131 2 Methods

132 2.1 The dynamic emission model



133
134 **Figure 1.** Map of the domain covered (Randstad area, the Netherlands) within this study, including major cities
135 Amsterdam, Rotterdam, The Hague, and Utrecht (underlined). The squares show the locations of the measurement
136 sites within the urban network configuration. The area of this domain is approximately 77x88km. Source: © Google Maps.

137 Although generally applicable, the dynamic emission model is initially developed for the Netherlands and focused
138 on Rotterdam (Fig. 1). This is one of the major cities in the Netherlands (about 625,000 inhabitants) with the
139 largest sea port of Europe to its west. It is located in a larger urbanized area (Randstad, about 7 million inhabitants)
140 with The Hague, Amsterdam and Utrecht being other major cities. A large area to the southwest of The Hague is
141 covered with glasshouses. The Rotterdam area is characterized by a complex mixture of residential and industrial
142 activities and therefore we distinguish five source sectors and a total of ten sub-sectors to construct its emissions
143 (see Table 1). Note that, for simplicity, only the largest source sectors are included, which are responsible for >95



144 % of the CO₂ emissions in the area. The main goal is to get a reasonable first estimate of the emission landscape
 145 using readily available data.

146 **Table 1. Overview of source sectors and subsectors distinguished in the dynamic emission model, including their short**
 147 **name used in the figures, whether they are represented as point or area sources, and their approximate contribution to**
 148 **the total CO₂ emission in Rotterdam. Crosses indicate which emission factors (EF), and tracer ratios of CO, NO_x or**
 149 **SO₂ (R_{CO}, R_{NO_x}, R_{SO₂}) are part of the state vector and circles indicate whether they are also part of the short state**
 150 **vector (see Sect. 2.3).**

Source sector	Subsector	Short name	Source type	Contribution	EF	R _{CO}	R _{NO_x}	R _{SO₂}
Power plants	Gas-fired power plants	1A	Point	37 %	XO	X	X	
	Coal-fired power plants	1B			XO	X	X	X
Non-industrial combustion	Households	2A	Area	15 %	XO	XO	X	X
	Glasshouses	2B			XO	X	X	
Industry		3	Point	39 %	XO	XO	XO	XO
Road traffic	Cars	7A	Area	6 %	XO	XO	XO	
	Heavy duty vehicles	7B			XO	XO	XO	
Shipping	Ocean shipping	8A	Area	3 %	XO	X	XO	XO
	Inland shipping	8B			XO	X	XO	XO
	Recreational shipping	8C						

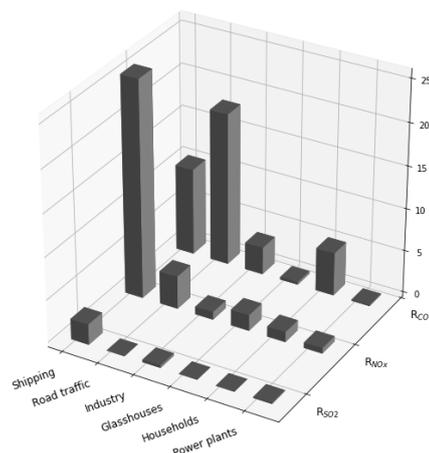
151
 152 The emissions are calculated in four steps. First, the yearly, national emission is calculated per sector using
 153 reported annual activity data and CO₂ emission factors. Second, we apply temporal disaggregation to hourly
 154 emissions using time profiles based on a combination of default temporal profiles, and environmental conditions.
 155 Third, we downscale the national totals to 1x1 km² resolution using statistical data, such as population density.
 156 Finally, our approach also allows uncertainties to be described in detail based on parameters in Eq. (2).

157 2.1.1 Step 1: Sectorial total emission calculations

158 Total annual emissions (F_X in kg yr⁻¹) per sector and species ($X=CO_2, CO, NO_x, SO_2$) are calculated as a function
 159 of the economic activity and an emission factor (adapted from Raupach et al. (2007)):

$$160 F_X = A \left(\frac{E}{A} \right) \left(\frac{F}{E} \right) R_X \quad (1)$$

161 where A is the amount of activity, such as vehicle kilometres driven or generated power, and E is the primary
 162 energy consumption (petajoule (PJ)). R_X is the emission ratio needed to calculate emissions of co-emitted species
 163 X from the CO₂ emissions, which is specific for each economic sector (R_{CO_2} is always 1, others are illustrated in
 164 Fig. 2). In this equation the term F/E is the emission factor (EF), i.e. the amount of CO₂ emitted per amount of
 165 energy consumed. The term E/A can be seen as a measure of energy efficiency, in which technological
 166 development plays an important role (Nakicenovic et al., 2000).



167
 168 **Figure 2. Emission ratios of CO:CO₂ (R_{CO}), NO_x:CO₂ (R_{NO_x}) and SO₂:CO₂ (R_{SO₂}) for specific source sectors based on**
 169 **the Dutch Pollution Release and Transfer Register (Netherlands PRTR, 2014). Units are in ppb ppm⁻¹. A value of 10 on**
 170 **the y-axis thus implies that for each 1000 moles of CO₂, 10 moles of the auxiliary tracer is emitted.**

171 The information needed in Eq. (1) comes from various inventories and national information sources. For example,
 172 changes in annual activity can be approximated based on national statistics such as the GDP (Gross Domestic
 173 Product), which can be a proxy for industrial activity. Or A can be based on environmental data such as the annual
 174 degree day sum based on the outside temperature, as proxy for the need for household heating in a particular year.
 175 The second term in Eq. (1) (E/A , the energy efficiency) can be estimated from energy consumption statistics, such
 176 as available from the International Energy Agency. Note that this term can show a large trend in case of
 177 technological development. The last terms in Eq. (1) (F/E and R_x , the emission factors) are the most uncertain
 178 ones, because the emission factor is dependent on the fuel mix and the energy efficiency, which itself can vary
 179 with environmental conditions (e.g. a cold engine on a winter day burns less efficiently). It can therefore differ
 180 significantly between countries. Emission factor values that are generally valid can be gathered from the
 181 Intergovernmental Panel on Climate Change (IPCC) or the European Environmental Agency (EEA), while
 182 country-specific values are typically less easily accessible. For our study area, we have access to both EEA data,
 183 as well as to Netherlands-specific numbers and even to Rijnmond-specific values (PRTR). See Appendix A for a
 184 full overview of the data used.

185

186 2.1.2 Step 2: Temporal profiles and parameterizing activity

187 The second step is to disaggregate the yearly emissions to hourly emissions by calculating time profiles, such that
 188 Eq. (1) becomes "dynamic":

$$189 F_{X,t} = A \left(\frac{E}{A} \right) \left(\frac{F}{E} \right) R_X T_t \quad (2)$$

190 where T_t is the hourly time factor. Averaged over a year the value of T_t is 1.0, so that it only alters the temporal
 191 evolution and not the total emissions. Energy use is often specifically linked to an activity (A in Eq. (1) and Eq.
 192 (2)) on which temporal information is more readily available than on the resulting emissions. Therefore, T_t can be
 193 calculated in two ways: 1) by directly using temporally explicit activity data or 2) by parameterizing temporal
 194 variations from environmental and/or economic conditions. When activity data is available the first option is



195 preferable. However, in data-sparse regions the second option might be necessary to implement, which is still an
 196 improvement compared to long-term average profiles as commonly used as we will discuss next for several sectors
 197 represented in our dynamic emission model.

198 Non-industrial combustion is dominated by households' natural gas consumption to heat houses, for cooking, and
 199 for warm water supply. A Dutch energy provider has a dataset publicly available from about 80 smart meters for
 200 the year 2013 with hourly gas consumption (Liander, 2018). It clearly shows a seasonal cycle, but also more small-
 201 term variations (daily data are shown in Fig. 3). We also see higher gas consumption in the beginning of the year,
 202 where the first three months of 2013 had some long, cold spells.

203 The use of energy for household heating is connected to the outside temperature. Previous studies have therefore
 204 used the concept of heating degree days to describe the temporal variability in emissions from households (Mues
 205 et al., 2014; Terrenoire et al., 2015). This concept assumes that heating only takes place below a certain temperature
 206 threshold (here 18°C) and the hourly time factor can be defined as:

$$207 \quad T_t = H / \bar{D} \quad (3)$$

208 where H is the heating degree day factor ($H = \max(291.15 - \bar{T}_{2m}, 0)$) based on the daily mean outside temperature
 209 at 2 m. \bar{D} is the yearly average heating degree day ($\bar{D} = \frac{1}{N} \sum_{j=1}^N H$). However, gas consumption related to warm
 210 water supply and cooking is largely independent of the outside temperature and therefore a constant offset is
 211 included in the heating degree day factor:

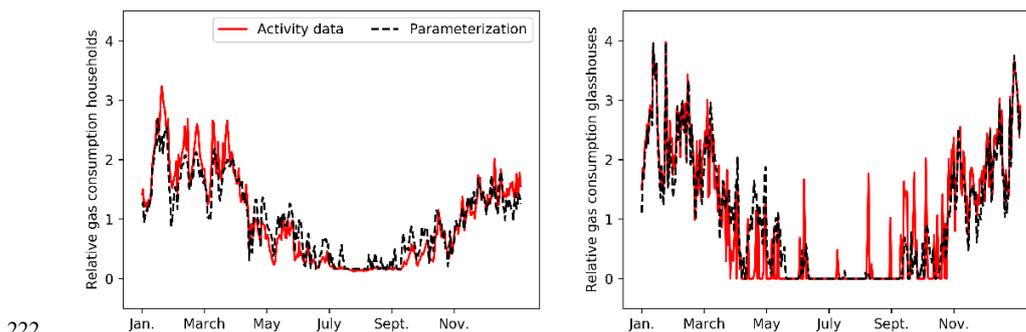
$$212 \quad H_f = H + f \cdot \bar{D} \quad (4)$$

213 where f is the constant offset. We assumed an offset of 20 %, similar to Mues et al. (2014). The time factor can
 214 now be defined as:

$$215 \quad T_t = H_f / \bar{D}_f \quad (5)$$

216 where the average heating degree day accounted for the constant offset $\bar{D}_f = (1 + f)\bar{D}$.

217 We compared the heating degree day method with gas consumption data on a daily basis (Fig. 3). The degree day
 218 function follows the gas consumption data very well, including the higher consumption at the start of the year,
 219 reaching an R^2 of 0.90 ($N=365$). The gas consumption of consumers also has a diurnal pattern with peaks in the
 220 early morning and late afternoon. Therefore, a diurnal profile can be estimated based on typical working hours.
 221 For hourly data R^2 is 0.80 ($N=8760$, not shown).



222
 223 **Figure 3. Daily time profiles for households (left) and glasshouses (right). Solid red lines are based on true activity data,**
 224 **whereas dashed black lines are parameterizations based on the degree day function.**



225 For the energy consumption of glasshouses there is no true activity data available. Instead, we use modelled daily
226 energy consumption for a typical Dutch glasshouse cultivating tomatoes (courtesy of Bas Knoll, TNO) as the
227 ‘truth’ (activity data). This time profile is calculated for typical meteorological conditions, such that the order of
228 magnitude and the peaks are representative for an average year. There is almost no energy consumption during the
229 summer, which indicates that there is no constant offset. So, we use Eq. (3) to determine the emission factor.
230 Moreover, we use a lower temperature threshold of 15 °C to get a better fit with the observed energy consumption.
231 The estimated function compares well with the activity data (Fig. 3) with an R^2 of 0.85 (N=365).

232 The diurnal cycle of glasshouse emissions is likely to be different from that of household emissions. Yet we lack
233 data to establish a diurnal cycle. We therefore use the same diurnal profile as for households, although this is likely
234 to be incorrect.

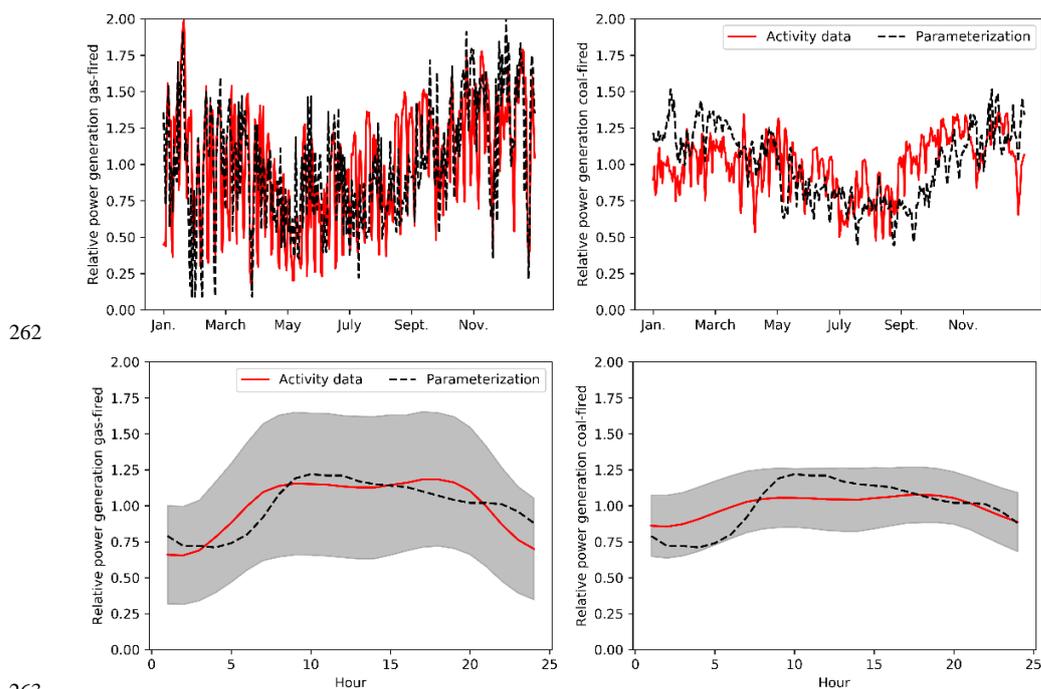
235 Power plants can use different fuels such as hard coal, natural gas or biomass. In the Netherlands coal-fired and
236 gas-fired power plants account for 80–85 % of the total energy production. The remainder comes mainly from
237 wind energy (5–6 %) and biomass burning (5–6 %). Power generation data are reported by the European Network
238 of Transmission System Operators for Electricity (ENTSO-E), which has detailed data available for the whole of
239 Europe. Coal-fired power plants are currently the main source of energy and their generation is relatively stable
240 compared to other sources. It does, however, show a seasonal cycle with less energy production during the summer
241 months. Gas-fired power plants have a larger temporal variability as they are mainly used as back-up for peak
242 hours, depending also on the amount of renewable energy that is available.

243 We use Eq. (5) to estimate the time profiles of both coal- and gas-fired power plants. Linear regression analysis
244 shows that the coal-fired power generation is correlated with degree days ($R^2 = 0.17$). In this case we use a large
245 constant offset of 80 % and a threshold of 25 °C which were chosen to best match the actual power generation
246 data. The offset is much larger than for households because there is always a basic energy demand from the
247 industry. In contrast, the gas-fired power plants are (negatively) correlated with the wind speed ($R^2 = 0.13$) and
248 incoming solar radiation ($R^2 = 0.10$), indicating the need for gas-fired power generation in the absence of
249 renewable sources. Therefore, we replace the temperature used to calculate H_f in Eq. (4) with the multiplication of
250 wind speed and incoming solar radiation:

$$251 \quad H = \max(10 - \bar{u}, 0) \cdot \max(150 - \bar{R}, 0) \quad (6)$$

252 where u is the wind speed (m s^{-1}) and R the incoming solar radiation (J cm^{-2}). Here we use a constant offset of 10
253 % and a threshold of 10 m s^{-1} and 150 J cm^{-2} .

254 The diurnal cycles for power plants can be based on socio-economic factors. For example, the energy demand
255 peaks early in the morning when people get ready to go to work and at the end of the afternoon when they get
256 home. We find this pattern in the actual power generation data, with coal-fired power plants being less variable
257 during the day than gas-fired power plants. The fixed profile from the European MACC-III emission inventory
258 (Denier van der Gon et al., 2011; Kuenen et al., 2014) matches reasonably well with gas-fired power plant profiles,
259 but it is less applicable for coal-fired power plants (Fig. 4). Overall, the estimated profiles for gas-fired power
260 plants (hourly data) have an R^2 of 0.32 (N=8784) when compared to the activity data. For coal-fired power plants
261 this is 0.21 (N=8784).



263

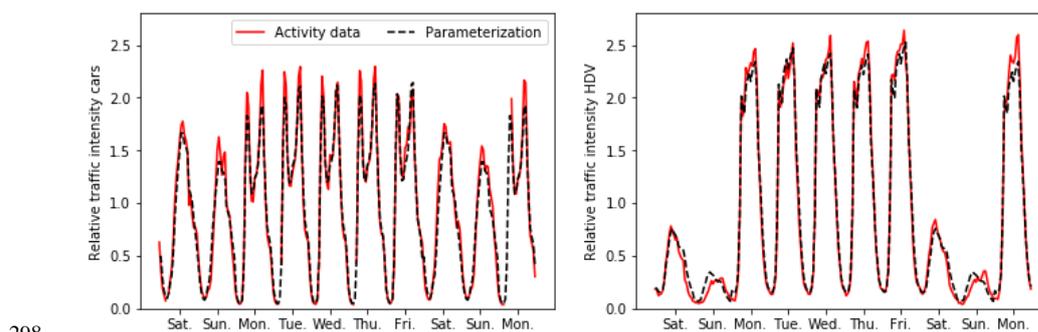
264 **Figure 4. (top row) Daily time profiles for gas-fired (left) and coal-fired (right) power plants. Solid red lines are based**
265 **on true activity data, whereas dashed black lines are parameterizations based on observed temperature (coal) and wind**
266 **speed/radiation (gas). (bottom row) Average diurnal cycle for gas-fired (left) and coal-fired (right) power plants. Solid**
267 **red lines are based on true activity data, whereas dashed black lines are fixed profiles from the MACC inventory (Denier**
268 **van der Gon et al., 2011; Kuenen et al., 2014). Shading gives the 1 σ variability of the diurnal cycle based on activity**
269 **data.**

270 The industrial sector consists of a wide range of activities, of which some are semi-continuous and only interrupted
271 by maintenance stops while others follow working hours. This makes it very difficult to predict the temporal
272 variability, especially for the overall sector. Since the largest CO₂ emissions are related to refineries and heavy
273 industry we will focus on these activities. We find a seasonal cycle in the reported industrial activity, with a small
274 decline during the summer and Christmas holidays. However, the variations are very small (max. 1 %). Therefore,
275 we assume constant emissions.

276 Road transport emissions can vary between different road and vehicle types (Mues et al., 2014), but are also
277 strongly dependent on environmental, socio-economic and driving conditions (such as the amount of stops, free-
278 flow versus stagnant conditions, and engine temperature). Traffic count data are often used to create average time
279 profiles for road traffic emissions, although with traffic counts we are unable to account for environmental and
280 driving conditions. Traffic counts for the Netherlands are made available by the Nationale Databank
281 Wegverkeersgegevens (NDW) and similar data is available in many developed countries. We differentiate between
282 two vehicle types (passenger cars + motorcycles (hereafter referred to as cars) and light duty + heavy duty vehicles
283 (hereafter referred to as HDV)) and three road types (highway, main road, urban road). We selected all available
284 locations for 2014 within or close to Rotterdam that distinguish 3-5 vehicle lengths and filtered for a minimum
285 data coverage of 75 %. This leaves us with 25 highway, 6 main road and 13 urban road locations. From this data
286 we make average time profiles (daily, weekly and monthly) per road and vehicle type, as is often done to



287 disaggregate road traffic emissions. Note that this method excludes any spatial variations (e.g. highways leading
288 towards the city vs. the beach), except for differentiating between road types.
289 Generally, HDV show a larger spread due to the low counts during the weekend (Fig. 5). Car counts on weekdays
290 show a morning and evening rush hour and they go down in between. In contrast, HDV counts peak throughout
291 the day and only go down after the evening rush hour. Moreover, the diurnal cycles are different during the
292 weekend than on weekdays. These patterns can be explained from socio-economic factors. Current time profiles
293 are often based on cars and are unable to correctly represent the temporal variability of HDV. This also affects the
294 spatial distribution of emissions and therefore we create average diurnal, weekly and seasonal profiles separately
295 for cars and HDV, for different road types and considering the day of the week. The comparison of true traffic
296 counts and averaged traffic counts results in R^2 values between 0.83 and 0.95 for hourly data for the whole year
297 (N between 2665 and 6471).



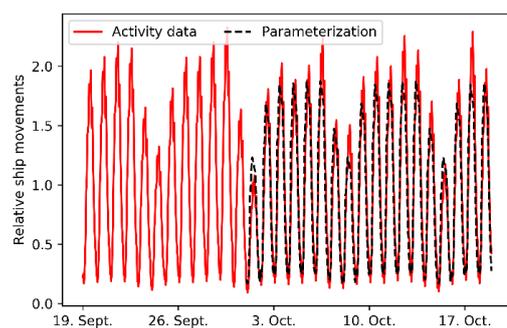
298
299 **Figure 5. Time profiles of passenger cars (left) and heavy-duty vehicles (right) road transport on highways for ten**
300 **randomly chosen days in March. Solid red lines are based on true activity data, whereas dashed black lines are**
301 **parameterizations based on averaged traffic counts for Rotterdam.**

302 Shipping emissions are dependent on the type of fuel used and whether ships apply slow-steaming. Additionally,
303 during loading and unloading ships still emit CO₂ and other pollutants, even though they are not moving. Such
304 information is currently not available, so instead we use information about the arrival and departure of ships in the
305 port of Rotterdam to make a time series of ship movements. Note that this only applies to large vessels that
306 transport goods and passengers and that the time profile will look quite different for recreational shipping.
307 However, large ships account for approximately 80 % of the total shipping emissions in the area of interest. Since
308 we lack information about other type of shipping movements, we will only account for large ships in the time
309 profiles.

310 We collected ship movements for one month (daily data) and an average diurnal profile. The diurnal cycle shows
311 a peak throughout the day, which corresponds well with the HDV road transport emission patterns on highways.
312 The reason for this is that HDV road transport is related to shipping movements, as HDV takes care of part of the
313 good transport further inland after the goods have arrived by ship. We also find a clear weekly pattern with less
314 ship movements during the weekend, although the decrease is less than for HDV road transport. This is likely
315 because large ships, such as entering the port of Rotterdam, continue travelling during the weekend. Therefore,
316 the weekly pattern resembles more that of car road transport on highways. Thus, we can estimate ship movements
317 by using the temporal profiles of HDV and cars on highways. This method is specifically tested for Rotterdam and
318 different patterns might be visible elsewhere. We also use HDV patterns for the seasonal variability, and final



319 parameterized and reported activity in this method reach an R^2 value of 0.89 for a period of 18 days with hourly
320 data ($N=432$) as shown in Fig 6.



321
322 **Figure 6. Daily time profiles for shipping.** Solid red line is based on true activity data, whereas dashed black line is a
323 parameterization based on traffic counts of heavy-duty vehicles (diurnal cycle) and cars (day-to-day variations) on
324 highways.

325 2.1.3 Step 3: Spatial disaggregation.

326 National total sectorial emissions need to be distributed into spatially explicit emissions for our study domain. The
327 spatial disaggregation of emissions has received quite some attention already from inventory builders. Existing
328 emission inventories can be used to describe the spatial disaggregation, if available for the region at high
329 resolution.

330 If not, simple default proxies for the spatial distribution are population density and the presence of roads or
331 waterways (e.g. OpenStreetMap). For example, main roads and urban roads are busiest in densely populated areas
332 and we assume emissions on main and urban roads are correlated with population density. Highways are used for
333 transport between cities and therefore emissions take place outside densely populated areas as well. Nevertheless,
334 highway transport is usually to and from densely populated areas, such that most emissions will take place close
335 to cities. We can therefore relate these emissions with the population density in the area of interest (in this case
336 Rijnmond) relative to the rest of the country, which places the same amount of the country-level emissions in our
337 case study domain as the gridded inventory. Additionally, the location of large power plants or industrial plants is
338 often known (for example from E-PRTR (Pollutant Release and Transfer Register), which can be used directly.
339 Although such information allows us to possibly construct a detailed fossil fuel model in data-sparse regions in
340 the future, in this study we focus first on the more easily implementable and less-developed parameterization of
341 temporal activity in different sectors (step 2) to assess whether this approach is promising enough for future
342 extension.

343 2.1.4 Step 4: Uncertainty analysis

344 The emission model we have constructed in steps 1–3 contains several parameters per source sector: activity,
345 emission factor, spatial proxy and time profile. For the analysis we only consider the emission factors and time
346 profiles, as we assume activity data and the spatial distribution to be well-known. As input for step 1 in the dynamic
347 emission model we use generalized parameters which we take from the IPCC, EEA and other organizations. These
348 databases also provide an uncertainty range, which we use in a final step to create a covariance matrix. The
349 covariance matrix describes the Gaussian uncertainty of these parameters (diagonal values) and error correlations
350 between parameters (off-diagonal values). From the covariance matrix we create an ensemble of parameters



351 (N=500) that represents their joint distributions, and we use them to calculate an ensemble of emissions. In this
352 Monte Carlo simulation, we transform some Gaussian parameters into log-normal distributions to account for non-
353 negativity, or to account for distributions with a very long tail (mainly emission ratios, which can become high in
354 specific cases where no emission reduction measures are taken). Appendix A summarizes the used parameter
355 values and uncertainties (including the shape of the distributions) and shows an example of the covariance matrix.
356 In a final step, we select the most important parameters which are either very uncertain or have a large impact on
357 the total emissions. This leaves us with the 44 parameters that we optimize in a set of data assimilation experiments,
358 described next. In Sect. 3.1 we report uncertainties in % (1σ) for normal distributions (CO_2) or as a 90 %
359 confidence interval (CI) for lognormal distribution (co-emitted species).

360 2.2 Data assimilation to estimate fossil fuel sources

361 The goal of data assimilation is to find a state at which the system is in optimal agreement with observations. In
362 this work, the observations we want to explore are the mole fractions of CO_2 and its co-emitted species while the
363 state of the system is the underlying spatiotemporal distribution of fossil fuel emissions. Such configurations are
364 sometimes referred to as “FFDAS” (fossil fuel data assimilation systems) applications, with a number of examples
365 in recent literature (Rayner et al., 2010; Asefi-Najafabady et al., 2014; Basu et al., 2016; Graven et al., 2018).
366 Given the sparsity of approaches explored so far, the dynamic emission model with its parameter driven emissions
367 we present here could lend itself well for application in an FFDAS, and this is what we explore through a set of
368 experiments with our own data assimilation methodology.

369 In this study we use the CarbonTracker Data Assimilation Shell (CTDAS) (v1.0) described in detail in Van der
370 Laan-Luijkx et al. (2017). Briefly, the CTDAS system is a flexible implementation of a square-root Ensemble
371 Kalman Filter (Whitaker and Hamill, 2002), which also allows lagged windows (i.e. smoothing instead of
372 filtering). The Ensemble Kalman Filter optimizes the cost function for unknown variables in the state vector x
373 using information from observations (y^0 with covariance R) and a prior estimate of the state vector (x^b with
374 covariance P).

$$375 J(x) = (y^0 - \mathcal{H}(x))^T R^{-1} (y^0 - \mathcal{H}(x)) + (x - x^b)^T P^{-1} (x - x^b) \quad (7)$$

376 In this function, \mathcal{H} is the observation operator that returns simulated mole fractions given the state vector. R and
377 P determine how much weight is given to the observations and prior estimate, respectively.

378 The optimized state vector (indicated with superscript a , whereas b refers to the prior estimates) which minimizes
379 the cost function is

$$380 x^a = x_t^b + K(y_t^0 - \mathcal{H}(x_t^b)) \quad (8)$$

381 and its covariance is

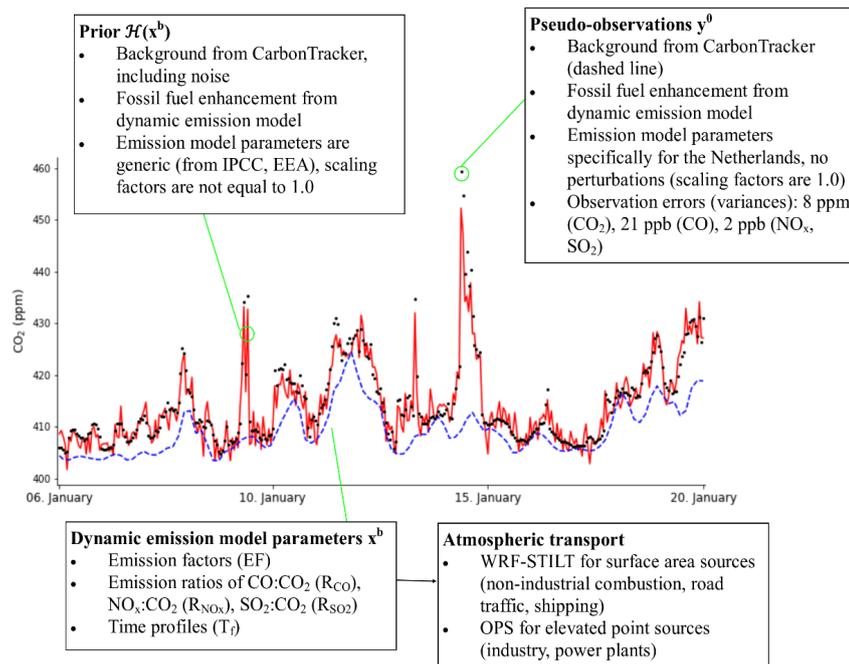
$$382 P_t^a = (I - KH)P_t^b \quad (9)$$

383 Here, H is the linearized observation operator and K is the Kalman gain matrix:

$$384 K = (P_t^b H^T)(H P_t^b H^T + R)^{-1} \quad (10)$$



385 The solutions of Eq. (8) and Eq. (9) are calculated as in Peters et al. (2005) using an ensemble of 80 members. The
386 choice for the ensemble size was based on the typical dimensions of our inverse problem, which has $N=1960$
387 observations and $M=44$ unknowns for the base run.



388
389 **Figure 7. Time series of pseudo-observations and prior CO_2 mole fractions and a summary of how these time series**
390 **were created.**

391 We have adapted CT DAS for smaller scale studies by replacing the typical observation operator \mathcal{H} , which is the
392 global TM5 transport model (Huijnen et al., 2010), with a combination of WRF-STILT footprints and the OPS
393 plume model, building on the methods described in Super et al., (2017a) and He et al. (2018). Moreover, we have
394 added our dynamic emission model to the observation operator so that we can sample its parameter distribution in
395 atmospheric mole fraction space. More details about the individual parts of this system are provided below and
396 are summarized in Fig. 7.

397 2.2.1 Observation operator

398 The observation operator translates the 44 parameters in the dynamic emission model first into emissions (through
399 Eq. (1) and Eq. (2)) and then into atmospheric mole fractions. The transport modelling consists of two parts. The
400 first part, the Weather Research and Forecasting-Stochastic Time-Inverted Lagrangian Transport (WRF-STILT,
401 (Nehrkorn et al., 2010) model, is used for surface emissions that are representative of large areas (i.e., not a point
402 source). STILT is a Lagrangian particle dispersion model that describes the footprint of a single measurement by
403 dispersing particles back in time (Gerbig et al., 2003; Lin et al., 2003). With this footprint the surface influence of
404 emissions on a single observation can be described. An advantage of this method is that it allows the pre-
405 calculation of linear atmospheric transport, which makes this part of the observation operator less computationally



406 demanding than running an ensemble of a full atmospheric transport model (like WRF with chemistry). The total
407 domain covered with WRF-STILT is 77 x 88 km (Fig. 1) and includes most of the Randstad.

408 To generate a footprint, 75 particles are released at the observation site at the start of the back-trajectory and
409 followed back in time. Given that the variability in hourly observations at an urban location is dominated by local
410 signals, we construct back-trajectories spanning 6 hours. This is based on the domain size, which could be covered
411 within 6 hours for typical wind speeds of 4 m s⁻¹. Within this time frame emissions can become well-mixed
412 throughout the boundary layer under normal daytime mixing conditions, such that emissions outside this range
413 can be represented by a boundary inflow. Footprints are generated for each hour within the back-trajectory to
414 account for hourly variations in the emissions. We drive STILT with meteorology from the WRF model (v3.5.1).
415 The WRF model was set up with two nested domains (15x15 and 3x3 km² horizontal resolution) and the STILT
416 footprints have a 1x1 km² resolution over the entire domain.

417 The second part of the transport modelling is a plume model. In a previous study we have shown that point source
418 (stack) emissions should be modelled with a plume model to better represent the limited dimensions of the stack
419 plume (Super et al., 2017a). Similarly, Vogel et al. (2013) have shown that the surface influence calculated by
420 STILT can lead to large model errors for stack emissions. Therefore, we include the OPS (Operational Priority
421 Substances, short-term version) plume model in our framework to model the transport and dispersion of stack
422 emissions (Van Jaarsveld, 2004; Sauter et al., 2016). OPS provides hourly concentrations at pre-defined receptor
423 points, which represent our measurement sites. The model keeps track of a plume trajectory, considering time-
424 varying transport over longer distances (e.g. changes in wind direction and dispersion). If for a time step a specific
425 plume affects the receptor, a Gaussian plume formulation is used to calculate the mole fraction caused by that
426 source based on the true travel distance along the trajectory. We drive the model with the same WRF meteorology
427 as STILT. Only primary meteorological variables (temperature, relative humidity, wind direction, wind speed,
428 precipitation, global radiation) are prescribed, secondary variables (e.g. boundary layer height, friction velocity)
429 are calculated by OPS itself and can differ from WRF.

430 Similar to the WRF-STILT model, we assume an influence time of 6 hours on our observations. However, in this
431 case we run the OPS model forward from -6 hours to the time of observation. We apply the OPS model only to
432 point source emissions within the Rijnmond area, as we found in a previous study that a plume model only has an
433 added value less than 10–15 km downwind from the stack (Super et al., 2017a). Point sources at more than 10–15
434 km from the observation site can be sufficiently represented with a Eulerian model. The OPS model input includes
435 detailed information about the exact stack height and heat content of the plume.

436 In addition to the fossil fuel contribution we also include background mole fractions for CO₂ and CO. NO_x and
437 SO₂ are short-lived and therefore the variations in the background are relatively small compared to the fossil fuel
438 signals. The CO₂ background is taken from the 3-D mole fractions of CarbonTracker Europe (Peters et al., 2010)
439 and also accounts for biogenic fluxes. The resolution of these CO₂ fields is 1x1° and we select the grid box that is
440 situated over Rotterdam. The 3-hourly data are linearly interpolated to get hourly background mole fractions that
441 are added to the fossil fuel signals calculated by the transport models. We use the strong wintertime correlation
442 between CO₂ and CO mole fractions ($r = 0.73$) to calculate CO background conditions from the CO₂ background.
443 This is not very accurate, but for the purpose of this OSSE it provides us with a decent estimate of the variability
444 in background mole fractions.

445 2.2.2 State vector



446 We populated the state vector with a selection of the most important parameters of the dynamic emission model,
447 based on their impact on the total emission uncertainty described in the results (Sect. 3.1). However, we
448 hypothesize that emission model parameters that are not part of the state vector are nevertheless uncertain and may
449 affect the results. Therefore, we include a total of 44 scaling factors in our state vector (x^b), and each scaling factor
450 is linearly related to a parameter from the dynamic emission model. The uncertainty in these parameters
451 (covariance matrix P) is derived from the Monte Carlo simulations described in Sect. 2.1, with the spread in the
452 emission model parameter values provided by the same databases of the IPCC and EEA. These uncertainty values
453 can also be found in Appendix A.

454 For this study we selected an arbitrary two-week period in January 2014 (6–20 January). Note that during the
455 summer the importance of source sectors might be different, e.g. there will be less heating from households.
456 Nevertheless, this period is sufficient to test the applicability of our DA system. We loop over the 14 days in our
457 study period, resulting in one posterior state vector for each day. We initialize our state vector for every new day
458 using the posterior values and posterior uncertainties from the previous day. Because the footprints we generated
459 extend backwards for six hours, the state vector for each day is effectively only constrained by the observations
460 from that same day, and hence we did not use a Kalman-smoother approach in this work in contrast to other
461 CTDAS applications.

462 Although this is a data-rich region, we use generic values for the prior emission model parameters which we take
463 from the IPCC, EEA and other organisations (Appendix A). These values are typically valid for a large region
464 (e.g. Europe) and not necessarily the best estimate for our regional case study. The reason that we use these values
465 is that they can provide a first estimate of the emissions in data-scarce regions where inverse modelling might add
466 most to our knowledge. With this set-up we can examine how well we can constrain the true emissions starting
467 with this generic, and widely available, information.

468 One major challenge in this study is to attribute the mismatch between the observed and modelled mole fractions
469 to a specific sector, as a CO₂ observation alone provides no details on the origin of the CO₂. Therefore, we include
470 three tracers (CO, NO_x and SO₂) that are co-emitted with CO₂ during fossil fuel combustion in a ratio (referred to
471 as R_{CO} , R_{NO_x} and R_{SO_2}) that is specific for each source sector (Fig. 2). Their (pseudo-)observations can inform us
472 about the source of the mismatch, but through their emission ratio to CO₂ they also constrain the magnitude of
473 CO₂ emissions in the emission model. The ratios R_{CO} , R_{NO_x} and R_{SO_2} used for this conversion to CO₂ emissions is
474 not fixed: for each of the co-emitted species we included them in the state vector. This recognizes that emission
475 ratios are highly variable and uncertain but play an important role in source attribution.

476 2.2.3 Pseudo-observations

477 In this work we create observing system simulation experiments (OSSEs), which use pseudo-observations instead
478 of true observations. The advantage of using pseudo-observations is that we can accurately examine the abilities
479 of our new approach without having to account yet for (often dominant) atmospheric transport errors.

480 The pseudo-observations used to optimize the emission model parameters are created using the same observation
481 operator as described above. The dynamic emission model is used to create realistic emissions with a high
482 spatiotemporal resolution. Yet in contrast to the prior, we use specific local (Dutch) values for the emission model
483 parameters. These parameters are considered to be the truth and are therefore not scaled (scaling factors are 1.0).
484 We found that these local parameter values are always within the uncertainty range of the general (prior) values,



485 so that the true solution is part of the distribution explored within the prior. This is confirmed in an experiment
486 with a small model-data mismatch and no noise on the background, which reproduces the true parameters very
487 well (not shown).

488 The resulting emissions are used in combination with the background mole fractions and transport calculated by
489 WRF-STILT and the OPS model to create pseudo-observations at the locations shown in Fig. 1. For the pseudo-
490 observations the original background time series are used, whereas in the inversion random noise is added to the
491 background mole fractions with a standard deviation of 2 ppm for CO₂. We assume no contribution from biogenic
492 CO₂ to the excess CO₂ over the background, which means that any biogenic contribution to CO₂ within our
493 footprint is the same as in the inflow from outside our domain, thus cancelling in the subtraction of the background
494 CO₂.

495 One simulated time series is illustrated in Fig. 7. The monitoring network consists of seven sites that are scattered
496 over the city of Rotterdam and the port. All sites exist in the national CO₂ or air quality measurement networks,
497 although not all species used in the inversion are observed at all locations. We only use the daytime (12–16 h LT)
498 observations to constrain our emissions. This is normally done to favour well-mixed conditions when simulated
499 transport is more reliable, and we want to mimic this limitation. We assume all instruments have an inlet at 10m
500 above ground level. In reality this is lower for several sites, but during the well-mixed daytime conditions the
501 difference is minimal.

502 The covariance matrix R describes the observation error. It accounts for errors related to instrumentation, but also
503 representativeness errors due to model transport, interpolation, and parameterization used in the dynamic emission
504 model. Although in principle such errors can be excluded in an OSSE, we prefer to use realistic estimates of these
505 errors to allow for the random errors that we applied to the prescribed boundary inflow, as well as to account for
506 some parameters in the emission model that are not optimized even though they contained uncertainty in the
507 pseudo-data creation. We base the R matrix on the calculated errors and variability caused by these specific
508 differences, and we end up with variances of 2.5 ppm (CO₂), 8 ppb (CO), 3 ppb (NO_x) and 1 ppb (SO₂).

509 2.3 Data Assimilation Experiments

510 We perform various experiments to examine the sensitivity of the system to different set-ups and sources of error.
511 The experiments are discussed here, and the detailed set-up of the inversions is summarized in Table 2. The base
512 run is labelled “Base”.

513 1) State vector definition: We start with a comparison of two different state vectors. For this purpose, we compare
514 the base run with an inversion (Short_state) which only includes the 21 most important parameters as identified in
515 the sensitivity analysis. This test allows us to examine the impact of erroneous, non-optimized emission model
516 parameters on the emission estimates. The results are discussed in Sect. 3.2.

517 2) Source attribution: Next we compare two monitoring network configurations which differ in the number of
518 tracers used. We perform an inversion with CO₂ as the only tracer (CO₂_only) and one with the full range of tracers
519 (Base) to assess the added value of including co-emitted species for source attribution. These tests address the
520 question whether co-emitted species can be used for source attribution. The results are discussed in Sect. 3.2.

521 3) Propagation: The third experiment is used to examine the effect of propagation of posterior values and
522 uncertainties on the final emission estimates. We compare the base run to a run that has no propagation
523 (No_propagation and CO₂_only_no_propagation) but instead starts from the same prior mean and uncertainty on



524 each of our 14 days considered. The runs without would allow the parameter values to change over time. The
525 results are discussed in Sect. 3.3.

526 **Table 2. Overview of the inversions: which tracers are included, the length of the state vector and whether posterior**
527 **values and uncertainties are propagated.**

Inversion name	Tracers	State vector length (per day)	Propagation to the next day
Base	All	44	Yes
Short_state	All	21	Yes
No_propagation	All	44	No
CO ₂ _only	CO ₂	44	Yes
CO ₂ _only_no_propagation	CO ₂	44	No

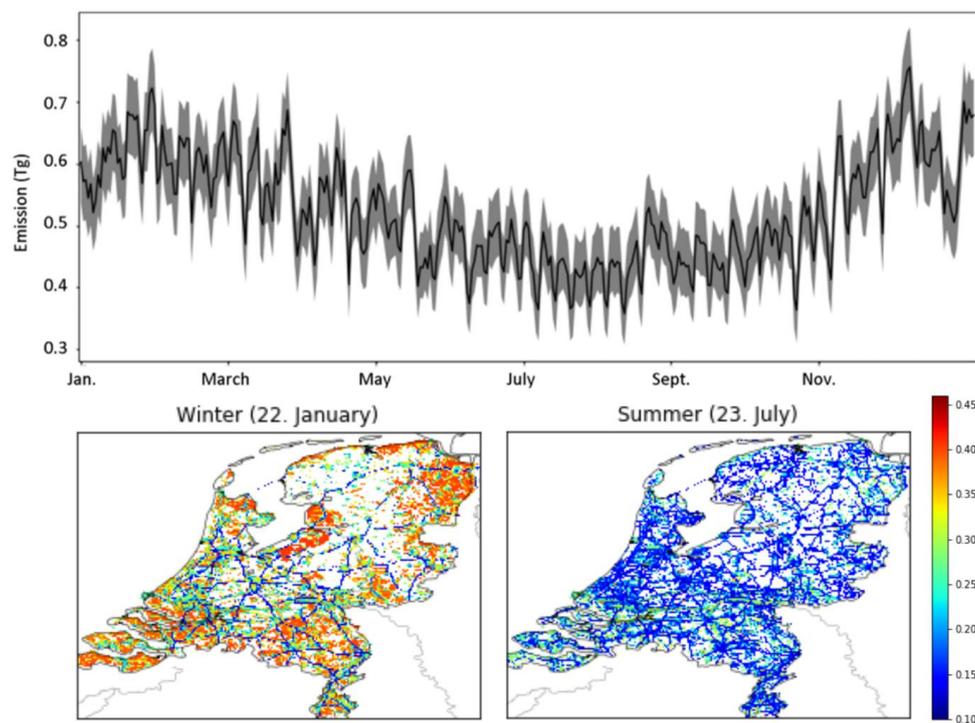
528 3 Results

529 Before demonstrating the use of our dynamic emission model in an inverse framework, we demonstrate its
530 application as a simple but versatile method to generate hourly gridded emissions for multiple species with full
531 covariances.

532 3.1 Dynamic emissions and their uncertainty

533 The total yearly emission of CO₂ for the Netherlands calculated with the dynamic emission model is 180 Tg CO₂
534 with an uncertainty of 15 % (1-sigma Gaussian based on 500 members of a Monte Carlo simulation). This matches
535 the total of the Dutch national emission inventory for 2014 by design (step 1), but the uncertainty on the latter was
536 estimated with a similar Monte Carlo simulation to be only 1 % for CO₂ in 2004 (Ramírez et al., 2006). This
537 smaller uncertainty is fully due to the use of country-specific emission factors with a much smaller range than we
538 derived from the IEA and IPCC inventories. Spatial disaggregation (step 2) does not affect the uncertainty of the
539 domain aggregated annual fluxes, and the time profiles (step 3) have no impact on the yearly total emissions. For
540 CO, NO_x and SO₂ the uncertainties in the dynamic emission model are much larger, with medians (CI's) of 6.5x10⁸
541 (1.3x10⁸–6.8x10⁹) kg CO yr⁻¹, 5.0x10⁸ (1.2x10⁸–5.1x10⁹) kg NO_x yr⁻¹, and 1.3x10⁸ (5.1x10⁶–2.2x10¹⁰) kg SO₂ yr⁻¹.
542 These ranges result from uncertainties in the assumed ratios of their release per unit of CO₂ emitted.

543 Below the annual scale, time profiles have an impact on the uncertainties as well. The daily emissions of the
544 Netherlands depend on the day and the season (Fig. 8) and range from 0.36 to 0.76 Tg CO₂ day⁻¹. The time series
545 shows a seasonal cycle with lower emissions during the summer. There is a clear weekly cycle with reduced
546 emissions during the weekend. The uncertainty in the total daily emission varies between 8 and 15 %, which is
547 similar to or lower than the uncertainty in the yearly total emissions. The explanation for these relatively low
548 uncertainties is that many uncertainties are temporally uncorrelated and their impacts on individual days partially
549 cancel out. Moreover, the largest sectors (coal-fired power plants and industry) already have a large uncertainty
550 and adding more uncertainty through the time profiles has little impact. Nevertheless, the uncertainties introduced
551 through the time profiles cause an uncertainty in daily CO₂ emissions of about 7 %, if the other uncertainties are
552 excluded from the analyses.



553

554 **Figure 8. (top) Time series of daily CO₂ emissions (in Tg CO₂ day⁻¹) and their uncertainty. Given is the interquartile**
555 **range (shaded area) and the median (line) from the ensemble. (bottom) Map of annual mean relative uncertainty of**
556 **emissions for the top 25 % pixels with the largest emissions, during a winter month (dominated by household gas- and**
557 **electricity use) and a summer month (electricity and road-traffic dominated).**

558

Differences in the relative contribution of different sectors are evident when looking at the map of uncertainties across the Netherlands (Fig. 8), reflecting both the most uncertain parameters, but also the dominant source sectors.

559

Winter emissions for example are dominated by household gas-usage, while industrial and traffic emissions give rise to uncertainty year-round at a 10–30 % level. We further identified the most important parameters per source sector with a Monte Carlo simulation per source sector (Fig. 9). Results shows that the road traffic and shipping sectors contain the smallest relative uncertainties, although the time profile for shipping causes an uncertainty of

560

about 7 % in the total shipping emissions. The industrial emissions are most uncertain, and this is almost exclusively due to the emission factor, which causes an uncertainty of 41 % in the total industrial emissions.

561

Similarly, the power plant emissions have a large relative uncertainty due to the uncertain emission factor of coal-fired power plants (19 %). Also, for households and glasshouses the emission factor is uncertain (14 % and 26 %, respectively), but here the time profiles also have a large impact (10 % and 16 %, respectively).

562

563

564

565

566

567

568

569

570

571

572

573

574

575

576

577

578

579

580

581

582

583

584

585

586

587

588

589

590

591

592

593

594

595

596

597

598

599

600

601

602

603

604

605

606

607

608

609

610

611

612

613

614

615

616

617

618

619

620

621

622

623

624

625

626

627

628

629

630

631

632

633

634

635

636

637

638

639

640

641

642

643

644

645

646

647

648

649

650

651

652

653

654

655

656

657

658

659

660

661

662

663

664

665

666

667

668

669

670

671

672

673

674

675

676

677

678

679

680

681

682

683

684

685

686

687

688

689

690

691

692

693

694

695

696

697

698

699

700

701

702

703

704

705

706

707

708

709

710

711

712

713

714

715

716

717

718

719

720

721

722

723

724

725

726

727

728

729

730

731

732

733

734

735

736

737

738

739

740

741

742

743

744

745

746

747

748

749

750

751

752

753

754

755

756

757

758

759

760

761

762

763

764

765

766

767

768

769

770

771

772

773

774

775

776

777

778

779

780

781

782

783

784

785

786

787

788

789

790

791

792

793

794

795

796

797

798

799

800

801

802

803

804

805

806

807

808

809

810

811

812

813

814

815

816

817

818

819

820

821

822

823

824

825

826

827

828

829

830

831

832

833

834

835

836

837

838

839

840

841

842

843

844

845

846

847

848

849

850

851

852

853

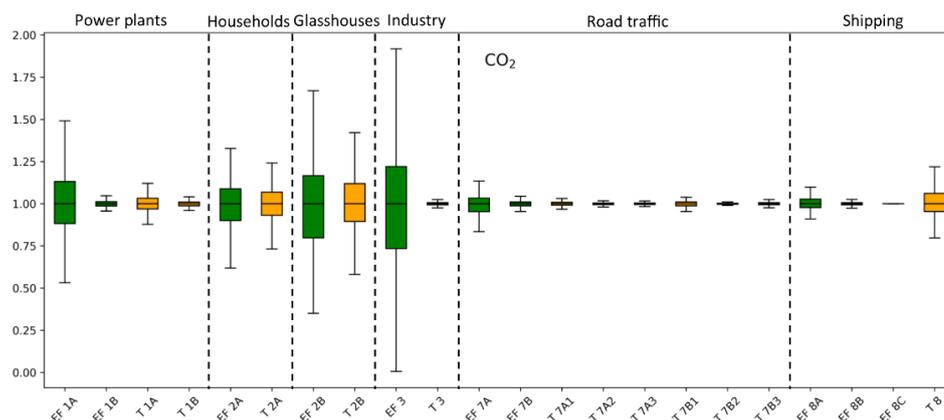
854

855

856

857

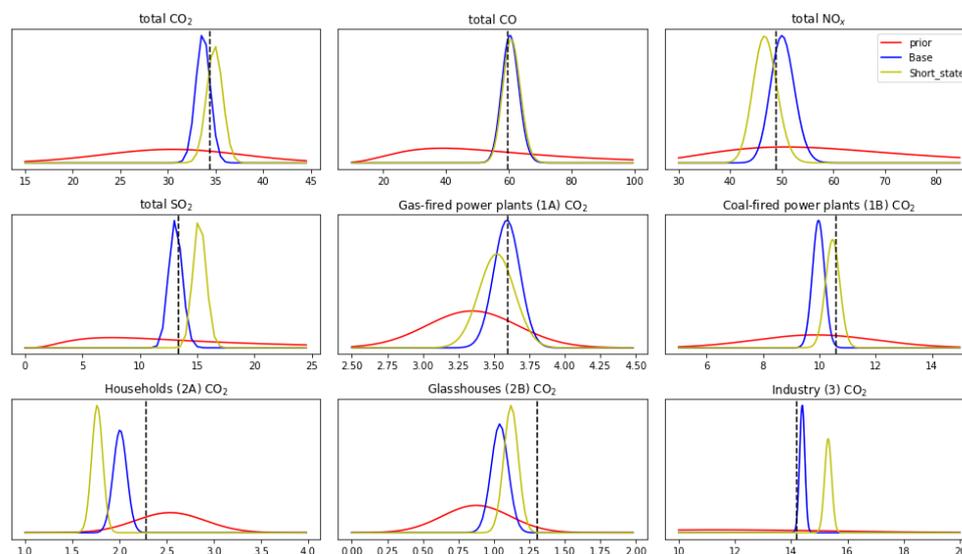
858



569
 570 **Figure 9.** Box plots showing the uncertainty in the CO₂ emissions from power plants (1A+1B), households (2A),
 571 glasshouses (2B), industry (3), road traffic (7A+7B) and shipping (8A+8B+8C) caused by individual parameters
 572 affecting that sector. Uncertainty is represented as the spread in daily (normalized) emissions from each ensemble
 573 member (N=500) over a full year (N=365). EF refers to an emission factor (green bars) and T to a time profile (orange
 574 bars). (Sub)sectors are indicated with their short names as summarized in Table 1. Note that the time profiles of road
 575 traffic emissions are specified per road type (1 = highway, 2 = main road, 3 = urban road). Minor parameters that have
 576 very small impacts on CO₂ emissions are not shown here (23 out of 44).

577 3.2 Optimizing dynamic emissions

578 In the base inverse modelling setup, our system is able to improve the mean estimate and reduce the uncertainty
 579 on total CO₂, CO, NO_x, and SO₂ emissions. Figure 10 shows the probability density function of these estimated
 580 total emissions, compared to the prior (using parameters derived from IPCC/EEA) and the truth (created with
 581 country-specific parameter values). Interestingly, the posterior result deteriorates slightly when using a shortened
 582 state vector in which 11 parameters of “minor” influence (such as the SO₂:CO₂ ratio of household emissions) are
 583 not optimized from their incorrect prior. This is caused by sporadic atmospheric signals that are dominated by
 584 household emissions, even if these emissions only contribute a small fraction to the total emissions. These signals
 585 are then used to update the emission factor, while the emission ratios are also incorrect.



586

587

588

589

590

Figure 10. Probability density functions of emissions per species or per source category (for CO₂ in units of Tg (CO₂) or Gg (CO, NO_x, SO₂). The truth is shown as a vertical dashed line, typically well-matched by the mean of the posterior in blue. Using a shortened state vector (yellow) deteriorates the total non-CO₂ emissions substantially and leads to misattribution of CO₂ emissions in minor categories such as 2A (households).

591

592

593

594

595

596

597

598

599

600

601

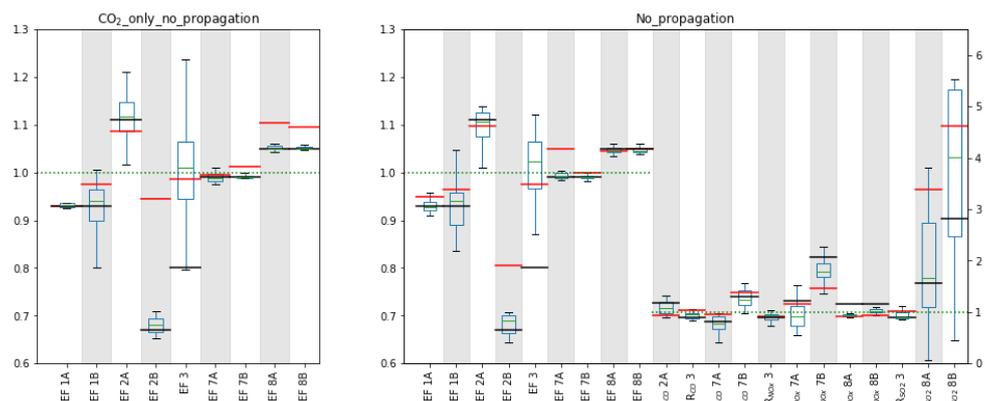
602

603

604

605

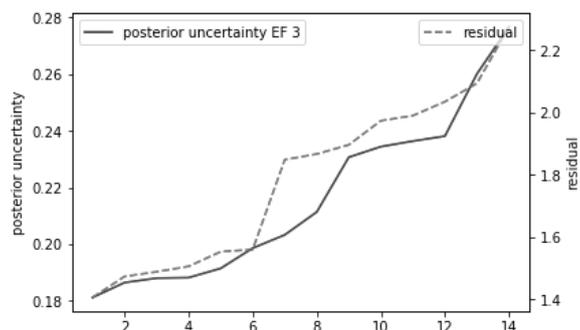
With CO₂ as the only tracer in the inversion we find that we can still estimate total CO₂ emissions quite well (truth-minus-optimized = 0.03 Tg CO₂ yr⁻¹), but we lose the capacity to attribute emissions to specific sectors. Instead, mainly the emission factor of the largest single source being industry (EF3) is optimized. We illustrate this in Fig. 11, using the No_propagation run. The large spread across the 14 individual days indicates that the emission factor jumps around within a large prior uncertainty distribution and is not well-constrained on each day. Some of the other emission factors show almost no deviation from the prior and little variability. Given the constraints posed by CO₂ observations alone, and the limited number of parameters that change the simulated CO₂, optimizing EF3 improves the results at the lowest costs. Introducing the co-emitted species allows the system to identify the source of a residual, and attribute it to the right parameters if sufficient sensitivity is present. This is especially true for those sectors that have relatively small emissions and/or uncertainties, like 2B and 1A. This is corroborated by the posterior covariance matrices (See Appendix B) which show a reduction in parameter correlations for those parameters (i.e., a better mathematical separation of the estimates) when all tracers are included in the estimate. For other parameters the median values are further from the truth than the prior (e.g. for R_{SO2} 8), which indicates that there is too little sensitivity to these parameters.



606
 607 **Figure 11. Spread (Q1-Q3) and median values of the parameter scaling factors for the fourteen individual days included**
 608 **in the CO₂_only_no_propagation (left) and No_propagation (right) inversions, and final value of the CO₂_only (left)**
 609 **and base (right) inversion (red lines). The prior values are indicated by the black lines and the truth is indicated with**
 610 **the green dotted lines (value of 1.0). The left y-axis is for the emission factors, the right y-axis for the tracer ratios. The**
 611 **inversion with all tracers shows more variability in the emission factors and larger deviations from the prior values.**

612 3.3 Localization and propagation of information

613 Propagating information on parameter values from one day to the next is often better than using the median of
 614 individual days' estimates as illustrated by the red lines in Fig. 11. Nevertheless, the sporadic detection of plumes
 615 with specific signatures suggests that a form of selection or localization of the strongest signals could reduce noise
 616 and improve the estimate for the No_propagation run. We therefore ranked the 14 daily independent parameter
 617 estimates based on their relative posterior uncertainty and the residuals in an attempt to find the most trustworthy
 618 parameter values. This ranking is done per parameter, so the best estimate of different parameters can be related
 619 to different days. The increase in residual (same for all parameters) and posterior uncertainty (of the industrial
 620 emission factor) is shown in Fig. 12, where the 3–5 highest ranked days have similar characteristics after which
 621 the reliability decreases. On the lower ranked days, atmospheric signals from that particular source sector are too
 622 small (or even absent) to update the parameters related to that source sector. A similar pattern is found for the other
 623 parameters (not shown), with 2–5 days of high sensitivity out of 14.

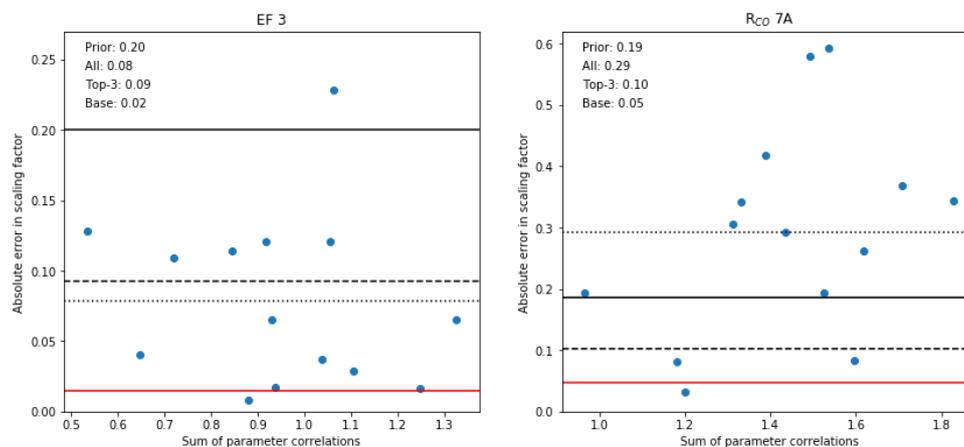


624
 625 **Figure 12. Increase in posterior uncertainty (1σ of unitless scaling factor) in the industrial emission factor (EF 3) and**
 626 **absolute mean residual of CO₂ (in ppm) from highest- to lowest-ranked days.**

627



628 When we use the top-3 averaged parameter values to calculate emissions we find for most sectors that the emission
629 estimate is similar to the base run, albeit with a larger uncertainty, while for a few specific sectors results
630 deteriorate. This suggests that selecting for strong signals can dampen spurious noise, but still does not improve
631 on the base run that includes full propagation of the covariances, hence carrying information on parameter
632 correlations that is partially lost in the No_propagation run.
633 From the posterior covariance matrices we can confirm our selection of “good” days, as these typically show
634 relatively weak correlations between parameters. For the industrial sector (emission factor, R_{NOx} , R_{SO2}) these are
635 typically weak on most days, and indeed the mean over the entire period already gives a robust estimate of the true
636 parameter value (Fig. 13). The parameters with the strongest correlations are R_{CO} of households and road traffic,
637 and their mean values tend to be dominated by a few outliers. Selecting days on which the posterior parameter
638 correlations are weak (i.e. the atmospheric signal clearly contains information about this specific parameter) results
639 in a large improvement compared to the prior or a 14-day average. Moreover, these results show a similar or better
640 performance as the top-3 selection based on Fig. 12 (0.08 for EF3 and 0.18 for R_{CO} 7A, not shown), and are closer
641 to the base run.



642 **Figure 13.** Scatter plot of the absolute error in the scaling factor of the industrial emission factor (EF 3) and R_{CO} of road
643 traffic (7A) against the sum of the parameter correlations of the same parameters. The correlation coefficients are -0.17
644 and 0.37 respectively. The horizontal lines give the average absolute error in the scaling factor for the prior (full black
645 line), if all 14 days are averaged (dotted line), and based on the 3 days with the smallest parameter correlations (dashed
646 line) and the result for the base run (full red line). The values are also given.
647

648 4 Discussion

649 4.1 Optimizing the dynamic emission model

650 The dynamic emission model has the advantage over static emission fields that its parameters are optimized, giving
651 physical meaning to the results. To reduce the size of the problem, the state vector can be populated with those
652 parameters that are most important and/or uncertain. However, we find that uncertain, non-included parameters
653 can still significantly affect the optimization. Therefore, the size of the state vector should be considered carefully
654 when applying this method. Moreover, we performed an experiment to establish the possibility to optimize the
655 time profiles as part of the state vector. Although we found some improvements, it appears to be difficult to
656 differentiate between the different variables in Eq. (2) that have a linear relationship based purely on the



657 observations. Therefore, the results are not shown and optimizing the temporal dynamics of the emission model
658 requires further work.

659 Additionally, we identified the base run as the simplest method to get good estimates, but we do note that our
660 current propagation scheme does not yet include error growth. That means that eventually the ensemble will
661 converge on a parameter value and discard incoming observational evidence, unless the covariance is inflated to
662 allow new updates. Examples of such a covariance inflation scheme are ample in literature and in principle not
663 difficult to include, but were not yet considered in this work as the time periods covered were still short.

664 Finally, we have demonstrated that tracers are suitable for source attribution. Several previous studies have used
665 co-emitted species as tracer for fossil fuel CO₂ by taking advantage of the specific emission ratio characteristics
666 of each source sector (Lauvaux et al., 2013; Lindenmaier et al., 2014; Turnbull et al., 2015) and came to similar
667 conclusions. Nevertheless, the uncertainty in emission ratios remains a source of error and therefore the
668 optimization of emission ratios with our system is a promising step forward. Using co-emitted species to identify
669 the total fossil fuel contribution to the observed CO₂ signal is more difficult (Turnbull et al., 2006). The reason for
670 this is that there is a large variability in emission ratios between sectors. This makes it difficult to establish an
671 average emission ratio for an urban area, because it depends strongly on the relative contribution of each source
672 sector and may vary over time.

673 **4.2 Radiocarbon and background definition**

674 Therefore, a nice addition to this inversion system would be the inclusion of radiocarbon measurements. The
675 radiocarbon isotope (¹⁴CO₂) can be used to simulate fossil fuel CO₂ records and has been applied successfully in
676 several inverse modelling studies (Turnbull et al., 2006; Levin and Karstens, 2007; Miller et al., 2012; Turnbull et
677 al., 2015; Basu et al., 2016; Wang et al., 2018). The radiocarbon measurements could be used directly in the
678 inversion (as we did with the co-emitted species) or be used to define a fossil fuel CO₂ record in advance (Fischer
679 et al., 2017; Graven et al., 2018). Our urban network detects average fossil fuel CO₂ signals of about 5 ppm with
680 peaks up to 50 ppm. This would result in $\Delta^{14}\text{C}$ signals (the ratio of ¹⁴CO₂ to ¹²CO₂) of around 13 up to 130 per
681 mille, which are certainly detectable with current techniques. However, observations of carbon isotopes are
682 expensive and currently not widely available, so their applicability is still limited. Besides $\Delta^{14}\text{C}$ other isotope
683 signatures and tracers can also provide additional information. For example, ¹³CO₂ and O₂/N₂ can give insight in
684 the dominant sources and sinks or fuel types (Lopez et al., 2013; Van der Laan et al., 2014) and as such be an
685 indicator for the transition from fossil fuels to biofuels. They might also help to separate between the stack
686 emissions of industry and coal- and gas-fired power plants.

687 An additional advantage of including the radiocarbon isotope is that the uncertainty in the background CO₂ can be
688 excluded, i.e. only the fossil fuel record is considered. Here, we choose to ignore the uncertainty in the background,
689 except in the definition of the covariance matrix *R*, and attribute all tracer residuals to the fossil fuel emissions.
690 Yet an incorrect definition of the background causes a large bias in the optimized emissions (Göckede et al., 2010).
691 There are also several other methods to deal with the non-fossil fuel related CO₂ signals. First, the uncertain
692 background can be added to the state vector and be optimized in the inversion. For example, He et al. (2018) have
693 shown that high-altitude aircraft observations are suitable to improve regional biosphere flux estimates by
694 correcting the bias in boundary conditions. Second, a mole fraction gradient over the area of interest can be
695 calculated using an upwind and downwind site such that the boundary inflow plays no role anymore (Turnbull et



696 al., 2015). This method was shown to reduce the impact of boundary inflow, but only when the wind direction is
697 more or less perpendicular to the gradient (Bréon et al., 2015; Staufer et al., 2016). Therefore, this method limits
698 the amount of useful measurements.

699 **4.3 Error correlations**

700 The dynamic emission model also allows us to study the correlations between model parameters, therefore giving
701 more insight in how information can be used in the system and which parameters are more challenging to separate.
702 Previously, Boschetti et al. (2018) have used the presence of error correlations between emissions of different
703 species and found that this reduces the posterior uncertainties for all species. They even show that the uncertainty
704 reduction increases with the correlation and that an incorrect definition of the error correlations may cause a
705 systematic bias in the posterior emission estimate. However, error correlations are only beneficial if the
706 atmospheric observations can distinguish between the correlated parameters. If this is not the case the presence of
707 parameter correlations can result in poorly constrained parameters and/or large posterior uncertainties. This is
708 especially true when parameters are sensitive to parameter correlations, as we show for R_{CO} of road traffic.
709 An important question is then why some emission model parameters are more sensitive to the presence of
710 parameter correlations than others. One hypothesis is that parameters with a lower prior uncertainty are more
711 sensitive to the presence of parameter correlations. The idea behind this is that if we reduce the diagonal value
712 (uncertainty) by a factor of 4 the off-diagonal value (parameter correlation) reduces by a factor of 2. This means
713 that the parameter correlation is relatively stronger if the uncertainty is lower (Boschetti et al., 2018). This
714 hypothesis cannot be confirmed by our results, as we only find a correlation of -0.27 between the prior uncertainty
715 and the sensitivity to parameter correlations (defined as the correlation between the posterior uncertainty and the
716 sum of the parameter correlations). The main difficulty here is that not all parameters can be discerned with the
717 observed atmospheric signals. Although we included the additional co-emitted tracers for source attribution, the
718 emission ratios have a large uncertainty and the system can have difficulties assigning residuals to either the
719 emission ratio or the emission factor. Yet if we calculate an average sensitivity and total posterior uncertainty per
720 sector (by combining the emission factor and emission ratios per sector) we find a correlation coefficient of -0.82.
721 This suggests that this hypothesis might indeed be correct and source sectors with larger parameter uncertainties
722 are less sensitive to the presence of parameter correlations.

723 **4.4 Atmospheric transport model errors**

724 In addition to the experiments described in Sect. 2.3 we conducted an experiment that focused on the role of
725 transport model errors by using observed meteorology to drive the OPS model in the inversion. Like many authors
726 before us (McKain et al., 2012; Brioude et al., 2013; Lauvaux et al., 2013; Bréon et al., 2015; Boon et al., 2016)
727 we found a large impact on the performance of our system and once again confirmed the need for accurate transport
728 models. This experiment is not further shown in this work because of its redundancy with previous conclusions.
729 Nevertheless, we performed this experiment to examine whether transport errors are important when the state
730 vector consists of parameters that are valid for the entire domain. Random errors, such as errors in the wind
731 direction, are unlikely to affect the optimized emissions much when averaged over a longer time period and
732 domain. This was shown by Deng et al. (2017), who found little variation in the average CO_2 emission for
733 Indianapolis using different configurations of WRF to calculate the transport. However, they did find an impact



734 on the spatial distribution of the emissions. This becomes important when optimizing a specific source sector that
735 is clustered in one place, such as the glasshouses. We found that the glasshouse sector is only correctly optimized
736 with a specific wind direction. If the modelled wind direction is wrong the residuals would thus not be attributed
737 to the glasshouse sector as it is not in the modelled footprint of the measurement site. As such, we conclude that
738 the footprint definition has an impact on the optimized parameters, despite that the parameters have no spatial
739 distribution. Similarly, Broquet et al. (2018) mention that the location and structure of a simulated urban plume
740 might differ significantly from the true plume characteristics due to errors in the simulated wind speed and wind
741 direction.

742 Systematic errors, whether in the modelled transport or in the observations, are more difficult to solve as they do
743 not cancel out when simulating a longer period, and this can lead to biased emission estimates (Meirink et al.,
744 2008; Su et al., 2011). Several methods have been suggested to overcome problems with an incorrect description
745 of atmospheric transport, such as using an ensemble of atmospheric transport model simulations (Angevine et al.,
746 2014) or the assimilation of meteorological observations (Lauvaux et al., 2013). The latter showed lower biases in
747 buoyancy and mean horizontal wind speed. Another method that is often used is the selection of well-mixed
748 afternoon hours to exclude stable conditions under which pollutant dispersion is often poorly represented (Lauvaux
749 et al., 2013; Bréon et al., 2015; Boon et al., 2016). Such data selection however leads to a bias in the estimated
750 emissions when the diurnal cycle is not correctly accounted for (Super et al., 2019).

751 Here, we also applied a daytime selection criterion to mimic this situation. However, we found that night time
752 hours could be very useful to constrain our emissions. In our DA system we use residual fossil fuel enhancements
753 over a background (prior - true mole fraction enhancement) to constrain the fossil fuel fluxes. The larger the
754 residual, the more information can be gained from it since the impact of the observation error (R matrix) is
755 relatively small. If, for example, the industrial emission factor is underestimated by 10 %, the residual industrial
756 enhancement (given a linear relationship between the emission factor and the total emission from this sector) will
757 be 10 % of the pseudo-observed mole fraction. This means that a large signal from the industry is needed to reach
758 a residual that is larger than the observation error (σ is 1.6 ppm for CO_2). Looking at the time series of pseudo-
759 observations we find that such large signals mostly occur during night time or in the early morning. Therefore, the
760 inversion could benefit strongly from an improved description of night time boundary layers and stable conditions,
761 so that the large night time enhancements can be used to constrain the fossil fuel fluxes.

762 5 Conclusions

763 The aim of this study was to examine how well our DA system can quantify urban CO_2 emissions per source
764 sector. Since the prior consists of a dynamic fossil fuel emission model the model parameters are optimized rather
765 than the emissions themselves. The parameters are related to specific source sectors and to attribute residuals to
766 these sectors measurements of additional tracers (CO , NO_x and SO_2) are included in the inversions. We tested this
767 system to examine its ability to overcome some major limitations in current urban-scale inversions: source
768 attribution, definition of the prior and its uncertainties, and the sensitivity to errors in atmospheric transport.

769 We find that inverse modelling at the urban scale is feasible when the observations contain a lot of information
770 about the different source sectors. When only CO_2 mole fractions are used in the inversion the total CO_2 emission
771 are well-constrained. Additional tracers are an important addition to the inversion framework in order to discern
772 the information belonging to specific source sectors and emission model parameters. However, even more tracers



773 might be needed to fully capture the heterogeneity of the emission landscape. Moreover, we argue that a dynamic
774 emission model has some major advantages over regular emission maps, allowing us to constrain physically
775 relevant parameters even in the absence of good prior information.
776 Nevertheless, quite some challenges remain. Transport modelling at this small scale needs to be improved to be
777 able to use real urban observations, as under current conditions the transport error strongly dominates the results.
778 Especially improving the description of night time boundary layers could be beneficial, because large atmospheric
779 signals mostly occur during the period. For the future, additional advances need to be made to include satellite
780 observations in the inverse modelling framework. The advantage of satellite data is that it covers data-sparse
781 regions and with a larger view it can differentiate between the urban dome with high pollution levels and the
782 cleaner rural areas, which is a nice addition to in situ measurements.

783 **Code and data availability**

784 The availability of the CTDAS (v1.0) code is described in a previous publication (Van der Laan-Luijkx et al.,
785 2017), which forms the basis of the system described in this paper. Minor changes have been made to include the
786 dynamic emission model. Revised code and the additional module used to describe the dynamic emission model
787 and the creation of pseudo-observations is included as Supplement, as is a script used for the emission uncertainty
788 analysis (Monte Carlo simulation). Input data for the dynamic emission model are taken from open, online
789 databases and are summarized in Appendix A, including their data sources. Example input files for CTDAS and
790 the OPS model are also included as Supplement.



791 **Appendix A**

792 **Table A1. Overview of all parameters in the dynamic emission model, their unit, function type, expected value and**
 793 **uncertainty (range).**

Parameter	(Sub)sector	Unit	Function type	Expected value	Uncertainty
Emission factor^(a)	Coal-fired power plants ^(c)	kg PJ ⁻¹	Normal	1.01E8	23 %
	Gas-fired power plants ^(c)	kg PJ ⁻¹	normal	5.61E7	10 %
	Households ^(c)	kg PJ ⁻¹	normal	5.89E7	14 %
	Glasshouses ^(c)	kg PJ ⁻¹	normal	5.61E7	25 %
	Industry ^(d)	kg PJ ⁻¹	normal	7.66E7	40 %
	Road traffic cars ^(e)	kg PJ ⁻¹	normal	7.24E7	10 %
	Road traffic HDV ^(e)	kg PJ ⁻¹	normal	7.33E7	5 %
	Ocean shipping ^(f)	kg PJ ⁻¹	normal	7.76E7	5 %
	Inland shipping ^(f)	kg PJ ⁻¹	normal	7.30E7	5 %
	Recreational shipping ^(f)	kg PJ ⁻¹	normal	7.10E7	5 %
Emission ratio CO:CO₂	Coal-fired power plants ^(c)	kg kg ⁻¹	lognormal	1.29E-4	8.7E-7–2.9E-4
	Gas-fired power plants ^(c)	kg kg ⁻¹	lognormal	8.47E-4	3.4E-4–2.5E-3
	Households ^(c)	kg kg ⁻¹	lognormal	3.88E-3	8.3E-4–9.6E-3
	Glasshouses ^(c)	kg kg ⁻¹	lognormal	5.40E-4	3.1E-5–7.7E-4
	Industry ^(d)	kg kg ⁻¹	normal	2.06E-3	40 %
	Road traffic cars ^(e)	kg kg ⁻¹	lognormal	1.32E-2	8.0E-5–6.5E-2
	Road traffic HDV ^(e)	kg kg ⁻¹	lognormal	2.22E-3	9.3E-5–1.3E-2
	Ocean shipping ^(f)	kg kg ⁻¹	normal	2.32E-3	30 %
	Inland shipping ^(f)	kg kg ⁻¹	normal	3.42E-3	30 %
	Recreational shipping ^(f)	kg kg ⁻¹	normal	2.96E-1	30 %
Emission ratio NO_x:CO₂	Coal-fired power plants ^(c)	kg kg ⁻¹	lognormal	5.94E-4	3.0E-4–9.4E-4
	Gas-fired power plants ^(c)	kg kg ⁻¹	lognormal	2.00E-3	2.6E-4–3.7E-3
	Households ^(c)	kg kg ⁻¹	lognormal	1.50E-3	4.8E-4–3.3E-3
	Glasshouses ^(c)	kg kg ⁻¹	lognormal	1.63E-3	5.0E-4–3.5E-3
	Industry ^(d)	kg kg ⁻¹	normal	6.56E-4	40 %
	Road traffic cars ^(e)	kg kg ⁻¹	lognormal	1.76E-3	9.0E-5–7.5E-3
Road traffic HDV ^(e)	kg kg ⁻¹	lognormal	1.11E-2	3.3E-4–3.7E-2	



	Ocean shipping ^(f)	kg kg ⁻¹	normal	2.32E-2	30 %	
	Inland shipping ^(f)	kg kg ⁻¹	normal	1.37E-2	30 %	
	Recreational shipping ^(f)	kg kg ⁻¹	normal	1.97E-3	30 %	
Emission ratio SO₂:CO₂	Coal-fired power plants ^(e)	kg kg ⁻¹	lognormal	1.66E-4	2.9E-5–4.4E-4	
	Gas-fired power plants ^(e)	kg kg ⁻¹	lognormal	5.01E-6	2.9E-6–7.2E-6	
	Households ^(e)	kg kg ⁻¹	lognormal	2.21E-5	1.4E-5–6.7E-5	
	Glasshouses ^(e)	kg kg ⁻¹	lognormal	8.91E-6	5.2E-6–1.3E-5	
	Industry ^(d)	kg kg ⁻¹	normal	4.28E-4	40 %	
	Road traffic cars ^(g)	kg kg ⁻¹	normal	1.01E-6	100 %	
	Road traffic HDV ^(g)	kg kg ⁻¹	normal	8.16E-7	100 %	
	Ocean shipping ^(f)	kg kg ⁻¹	lognormal	6.18E-3	3.3E-4–2.0E-2	
	Inland shipping ^(f)	kg kg ⁻¹	lognormal	6.57E-3	3.5E-4–3.0E-2	
	Recreational shipping ^(f)	kg kg ⁻¹	lognormal	3.14E-4	1.1E-4–7.0E-4	
Hourly time factor^(h)	Coal-fired power plants	-	normal	1	28 %	
	Gas-fired power plants	-	normal	1	43 %	
	Industry	-	normal	1	5 %	
	Households	-	normal	1	43 %	
	Glasshouses	-	normal	1	74 %	
	Road traffic cars highway	-	normal	1	18 %	
	Road traffic cars main road	-	normal	1	18 %	
	Road traffic cars urban road	-	normal	1	18 %	
	Road traffic HDV highway	-	normal	1	41 %	
	Road traffic HDV main road	-	normal	1	18 %	
	Road traffic HDV urban road	-	normal	1	48 %	
	Total shipping	-	normal	1	31 %	
	Energy consumption per activity data⁽ⁱ⁾	Total power plants	PJ/mln €	-	8.22E-4	-
		Households	PJ/dd ^(b)	-	0.199	-
Glasshouses		PJ/dd ^(b)	-	0.061	-	
Industry		PJ/mln €	-	7.05E-4	-	
Road traffic cars		PJ/mln €	-	3.98E-4	-	



	Road traffic HDV	PJ/mln €	-	2.01E-4	-
	Total shipping	PJ/mln €	-	1.51E-4	-
Fraction of total energy consumption per subsector^(j)	Total power plants: coal	-	-	0.62	-
	Total power plants: gas	-	-	0.38	-
	Road traffic cars: highway	-	-	0.47	-
	Road traffic cars: main road	-	-	0.28	-
	Road traffic cars: urban road	-	-	0.25	-
	Road traffic HDV: highway	-	-	0.56	-
	Road traffic HDV: main road	-	-	0.24	-
	Road traffic HDV: urban road	-	-	0.20	-
	Total shipping: ocean	-	-	0.79	-
	Total shipping: inland	-	-	0.20	-
	Total shipping: recreational	-	-	0.01	-

794 ^(a) Emission factor for coal-fired and gas-fired power plants include uncertainty due to variations in fuel type, including burning
 795 of biomass (5 % uncertainty). For households assume 8 % wood combustion based on CO₂ emission values (*Vernieuwd*
 796 *emissiemodel houtkachels*, by B.I. Jansen (TNO, 2016)), the remainder is natural gas (with 10 % uncertainty). For glasshouses
 797 assume only natural gas combustion, including 20 % additional uncertainty due to use of cogeneration plants. For road traffic
 798 cars assume 69 % gasoline, 29 % diesel and 2 % LPG (with 5 % uncertainty); for road traffic HDV assume 100 % diesel.

799 ^(b) dd = degree day

800 ^(c) Expected value and uncertainty based on IPCC Emission Factor Database (EFDB) using 2006 IPCC guidelines

801 ^(d) Expected value based on Emissieregistratie (emission) and CBS (energy consumption); uncertainty based on expert
 802 judgement

803 ^(e) Expected value and uncertainty based on the EMEP/EEA air pollutant emission inventory guidebook 2016

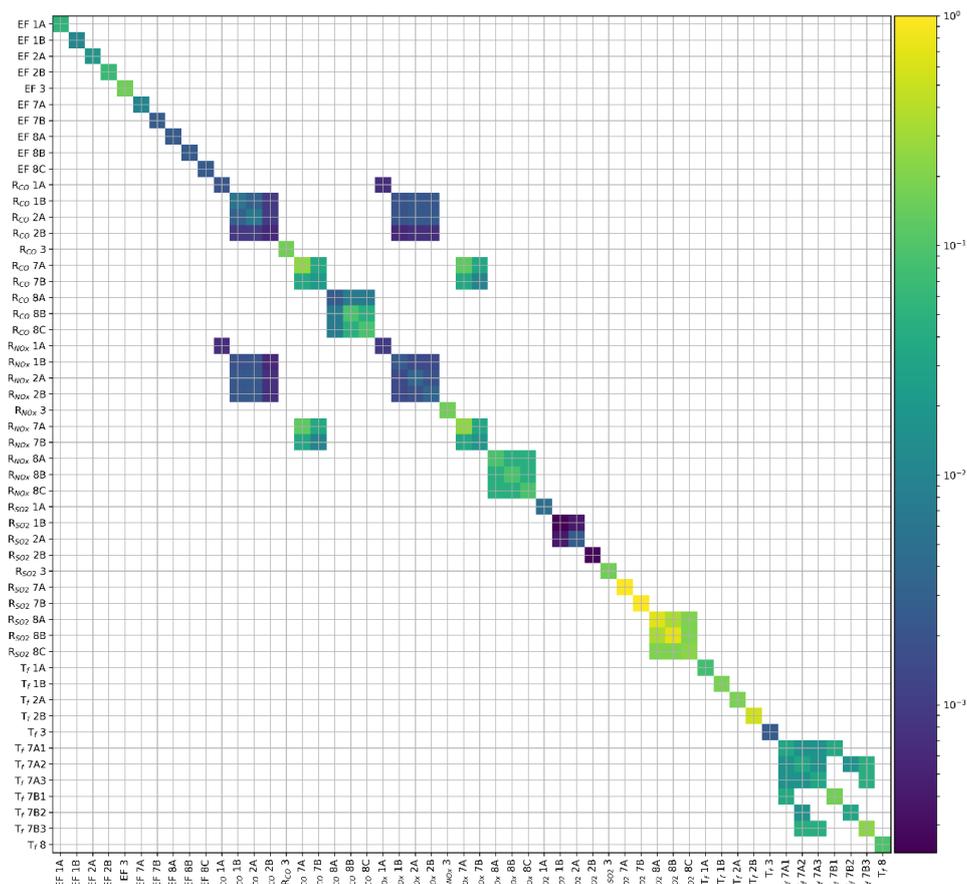
804 ^(f) Expected value and uncertainty based on *CO₂, CH₄, and N₂O emissions from transportation-water-borne navigation*, by Paul
 805 Jun, Michael Gillenwater, and Wiley Barbour (Good Practice Guidance and Uncertainty Management in National Greenhouse
 806 Gas Inventories)

807 ^(g) Expected value based on Air Pollutant Emission Factor Library (Finish Environment Institute); uncertainty based on expert
 808 judgement

809 ^(h) Uncertainties based on comparison activity data-based time profiles and estimated time profiles from environmental/socio-
 810 economic factors

811 ⁽ⁱ⁾ Expected value based on CBS (energy consumption, GDP) and KNMI (degree day sum)

812 ^(j) Expected value based on Emissieregistratie

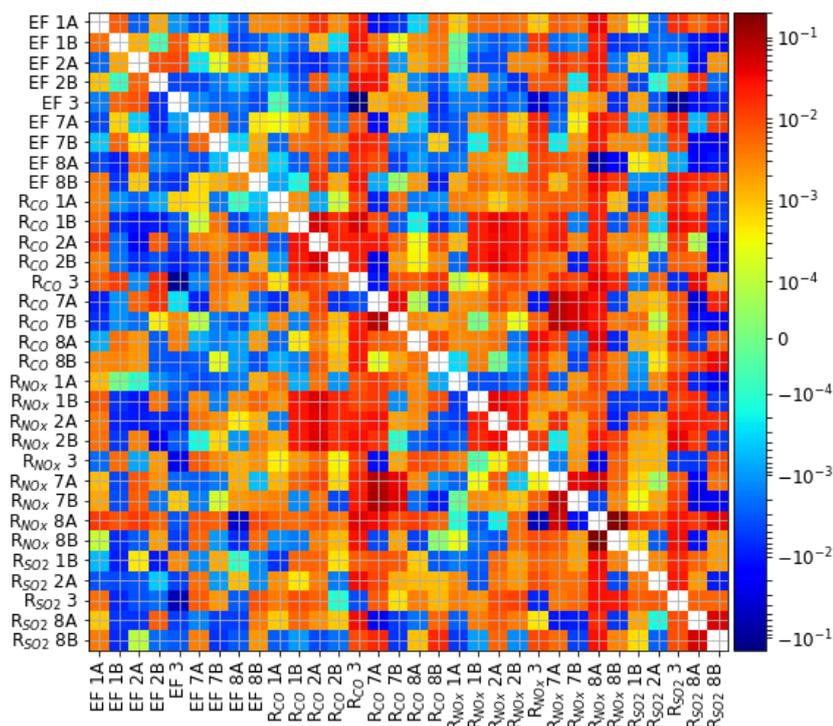


813

814 **Figure A1.** Covariance matrix for all parameters in the dynamic emission model. For all covariances we assume a
 815 correlation coefficient of 0.5. (Sub)sectors are indicated with their short names as summarized in Table 1. Note that the
 816 time profiles of road traffic emissions are specified per road type (1 = highway, 2 = main road, 3 = urban road).



817 **Appendix B**



818
819 **Figure B1.** Matrix showing the difference in correlation coefficient (r) between the CO_2 _only_no_propagation and
820 No_propagation run averaged for all 14 days, where positive differences indicate reduced parameter correlations when
821 all tracers are included (No_propagation). (Sub)sectors are indicated with their short names as summarized in Table 1.
822 For some parameters a strong reduction in parameter correlations is shown, indicating that with all tracers that
823 parameter can be more easily separated from others, for example the emission factors of industry and coal-fired power
824 plants (EF3 and EF1B).

825 **Author contribution**

826 The initial ideas are developed by WP, IS, HACDvdG and MKvdM. IS and SNCD developed the dynamic
827 emission model. IS and WP are responsible for setting up the inverse modelling experiments and prepared the
828 manuscript with contributions from all co-authors.

829 **Competing interests**

830 The authors declare that they have no conflict of interest.

831 **Acknowledgments**

832 This study was supported by the VERIFY project, funded by the European Union's Horizon 2020 research and
833 innovation programme under grant agreement No 776810; partly funded by EIT Climate-KIC project Carbocount-



834 CITY (APIN0029 2015-3.1-029 P040-04) and the EIT Climate-KIC Fellows programme (ARED0004 2013-1.1-
835 008 P017-0x).

836 References

- 837 Andres, R. J., Boden, T. A., and Higdon, D.: A new evaluation of the uncertainty associated with CDIAC estimates
838 of fossil fuel carbon dioxide emission, *Tellus B Chem. Phys. Meteorol.*, 66, 23616, 10.3402/tellusb.v66.23616,
839 2014.
- 840 Andres, R. J., Boden, T. A., and Higdon, D. M.: Gridded uncertainty in fossil fuel carbon dioxide emission maps,
841 a CDIAC example, *Atmos. Chem. Phys.*, 16, 14979-14995, 10.5194/acp-16-14979-2016, 2016.
- 842 Angevine, W. M., Brioude, J., McKeen, S., and Holloway, J. S.: Uncertainty in Lagrangian pollutant transport
843 simulations due to meteorological uncertainty from a mesoscale WRF ensemble, *Geosci. Model Dev.*, 7, 2817-
844 2829, 10.5194/gmd-7-2817-2014, 2014.
- 845 Asefi-Najafabady, S., Rayner, P. J., Gurney, K. R., McRobert, A., Song, Y., Coltin, K., Huang, J., Elvidge, C.,
846 and Baugh, K.: A multiyear, global gridded fossil fuel CO₂ emission data product: Evaluation and analysis of
847 results, *J. Geophys. Res.-Atmos.*, 119, 10,213-210,231, 10.1002/2013jd021296, 2014.
- 848 Basu, S., Miller, J. B., and Lehman, S.: Separation of biospheric and fossil fuel fluxes of CO₂ by atmospheric
849 inversion of CO₂ and ¹⁴CO₂ measurements: Observation System Simulations, *Atmos. Chem. Phys.*, 16, 5665-5683,
850 10.5194/acp-16-5665-2016, 2016.
- 851 Boon, A., Broquet, G., Clifford, D. J., Chevallier, F., Butterfield, D. M., Pison, I., Ramonet, M., Paris, J. D., and
852 Ciais, P.: Analysis of the potential of near-ground measurements of CO₂ and CH₄ in London, UK, for the
853 monitoring of city-scale emissions using an atmospheric transport model, *Atmos. Chem. Phys.*, 16, 6735-6756,
854 10.5194/acp-16-6735-2016, 2016.
- 855 Boschetti, F., Thouret, V., Maenhout, G. J., Totsche, K. U., Marshall, J., and Gerbig, C.: Multi-species inversion
856 and IAGOS airborne data for a better constraint of continental-scale fluxes, *Atmos. Chem. Phys.*, 18, 9225-9241,
857 10.5194/acp-18-9225-2018, 2018.
- 858 Bréon, F. M., Broquet, G., Puygrenier, V., Chevallier, F., Xueref-Remy, I., Ramonet, M., Dieudonné, E., Lopez,
859 M., Schmidt, M., Perrussel, O., and Ciais, P.: An attempt at estimating Paris area CO₂ emissions from atmospheric
860 concentration measurements, *Atmos. Chem. Phys.*, 15, 1707-1724, 10.5194/acp-15-1707-2015, 2015.
- 861 Brioude, J., Angevine, W. M., Ahmadov, R., Kim, S. W., Evan, S., McKeen, S. A., Hsie, E. Y., Frost, G. J.,
862 Neuman, J. A., Pollack, I. B., Peischl, J., Ryerson, T. B., Holloway, J., Brown, S. S., Nowak, J. B., Roberts, J. M.,
863 Wofsy, S. C., Santoni, G. W., Oda, T., and Trainer, M.: Top-down estimate of surface flux in the Los Angeles
864 Basin using a mesoscale inverse modeling technique: Assessing anthropogenic emissions of CO, NO_x and CO₂
865 and their impacts, *Atmos. Chem. Phys.*, 13, 3661-3677, 10.5194/acp-13-3661-2013, 2013.
- 866 Brophy, K., Graven, H., Manning, A. J., White, E., Arnold, T., Fischer, M. L., Jeong, S., Cui, X., and Rigby, M.:
867 Characterizing uncertainties in atmospheric inversions of fossil fuel CO₂ emissions in California, *Atmos. Chem.*
868 *Phys. Discuss.*, 2018, 1-44, 10.5194/acp-2018-473, 2018.
- 869 Broquet, G., Bréon, F. M., Renault, E., Buchwitz, M., Reuter, M., Bovensmann, H., Chevallier, F., Wu, L., and
870 Ciais, P.: The potential of satellite spectro-imagery for monitoring CO₂ emissions from large cities, *Atmos. Meas.*
871 *Tech.*, 11, 681-708, 10.5194/amt-11-681-2018, 2018.



- 872 Chevallier, F., Viovy, N., Reichstein, M., and Ciais, P.: On the assignment of prior errors in Bayesian inversions
873 of CO₂ surface fluxes, *Geophys. Res. Lett.*, 33, doi:10.1029/2006GL026496, 2006.
- 874 Ciais, P., Paris, J. D., Marland, G., Peylin, P., Piao, S. L., Levin, I., Pregger, T., Scholz, Y., Friedrich, R., Rivier,
875 L., Houwelling, S., and Schulze, E. D.: The European carbon balance. Part 1: Fossil fuel emissions, *Glob. Change*
876 *Biol.*, 16, 1395-1408, 10.1111/j.1365-2486.2009.02098.x, 2010.
- 877 Deng, A., Lauvaux, T., Davis, K. J., Gaudet, B. J., Miles, N., Richardson, S. J., Wu, K., Sarmiento, D. P., Hardesty,
878 R. M., Bonin, T. A., Brewer, W. A., and Gurney, K. R.: Toward reduced transport errors in a high resolution urban
879 CO₂ inversion system, *Elementa (Wash. D.C.)*, 5, 10.1525/elementa.133, 2017.
- 880 Denier van der Gon, H. A. C., Hendriks, C., Kuenen, J., Segers, A., and Visschedijk, A.: Description of current
881 temporal emission patterns and sensitivity of predicted AQ for temporal emission patterns, TNO, Utrecht, 2011.
- 882 Djuricin, S., Pataki, D. E., and Xu, X.: A comparison of tracer methods for quantifying CO₂ sources in an urban
883 region, *J. Geophys. Res.-Atmos.*, 115, 1-13, 10.1029/2009JD012236, 2010.
- 884 Fauser, P., Sørensen, P. B., Nielsen, M., Winther, M., Plejdrup, M. S., Hoffmann, L., Gyldenkerne, S., Mikkelsen,
885 M. H., Albrektsen, R., Lyck, E., Thomsen, M., Hjelgaard, K., and Nielsen, O.-K.: Monte Carlo (Tier 2) uncertainty
886 analysis of Danish Greenhouse gas emission inventory, *Greenhouse Gas Measurement and Management*, 1, 145-
887 160, 10.1080/20430779.2011.621949, 2011.
- 888 Fischer, M. L., Parazoo, N., Brophy, K., Cui, X., Jeong, S., Liu, J., Keeling, R., Taylor, T. E., Gurney, K., Oda,
889 T., and Graven, H.: Simulating estimation of California fossil fuel and biosphere carbon dioxide exchanges
890 combining in situ tower and satellite column observations, *J. Geophys. Res.-Atmos.*, 122, 3653-3671,
891 10.1002/2016jd025617, 2017.
- 892 Gerbig, C., Lin, J. C., Wofsy, S. C., Daube, B. C., Andrews, A. E., Stephens, B. B., Bakwin, P. S., and Grainger,
893 C. A.: Toward constraining regional-scale fluxes of CO₂ with atmospheric observations over a continent: 2.
894 Analysis of COBRA data using a receptor-oriented framework, *J. Geophys. Res.-Atmos.*, 108,
895 10.1029/2003JD003770, 2003.
- 896 Göckede, M., Turner, D. P., Michalak, A. M., Vickers, D., and Law, B. E.: Sensitivity of a subregional scale
897 atmospheric inverse CO₂ modeling framework to boundary conditions, *J. Geophys. Res.-Atmos.*, 115,
898 10.1029/2010JD014443, 2010.
- 899 Graven, H., Fischer, M. L., Lueker, T., Jeong, S., Guilderson, T. P., Keeling, R. F., Bambha, R., Brophy, K.,
900 Callahan, W., Cui, X., Frankenberg, C., Gurney, K. R., LaFranchi, B. W., Lehman, S. J., Michelsen, H., Miller, J.
901 B., Newman, S., Paplawsky, W., Parazoo, N. C., Sloop, C., and Walker, S. J.: Assessing fossil fuel CO₂ emissions
902 in California using atmospheric observations and models, *Environ. Res. Lett.*, 13, 065007, 10.1088/1748-
903 9326/aabd43, 2018.
- 904 Gurney, K. R., Patarasuk, R., Liang, J., Song, Y., O'Keefe, D., Rao, P., Whetstone, J. R., Duren, R. M., Eldering,
905 A., and Miller, C.: The Hestia fossil fuel CO₂ emissions data product for the Los Angeles Megacity (Hestia-LA),
906 *Earth Syst. Sci. Data Discuss.*, 2019, 1-38, 10.5194/essd-2018-162, 2019.
- 907 He, W., Van der Velde, I. R., Andrews, A. E., Sweeney, C., Miller, J., Tans, P., Van der Laan-Luijkx, I. T.,
908 Nehrkorn, T., Mountain, M., Ju, W., Peters, W., and Chen, H.: CTDAS-Lagrange v1.0: a high-resolution data
909 assimilation system for regional carbon dioxide observations, *Geosci. Model Dev.*, 11, 3515-3536, 10.5194/gmd-
910 11-3515-2018, 2018.



- 911 Hogue, S., Marland, E., Andres, R. J., Marland, G., and Woodard, D.: Uncertainty in gridded CO₂ emissions
912 estimates, *Earth's Future*, 4, 225-239, doi:10.1002/2015EF000343, 2016.
- 913 Huijnen, V., Williams, J., van Weele, M., van Noije, T., Krol, M., Dentener, F., Segers, A., Houweling, S., Peters,
914 W., de Laat, J., Boersma, F., Bergamaschi, P., van Velthoven, P., Le Sager, P., Eskes, H., Alkemade, F., Scheele,
915 R., Nédélec, P., and Pätz, H. W.: The global chemistry transport model TM5: description and evaluation of the
916 tropospheric chemistry version 3.0, *Geosci. Model Dev.*, 3, 445-473, 10.5194/gmd-3-445-2010, 2010.
- 917 Hutchins, M. G., Colby, J. D., Marland, G., and Marland, E.: A comparison of five high-resolution spatially-
918 explicit, fossil-fuel, carbon dioxide emission inventories for the United States, *Mitig. Adapt. Strat. Gl.*, 22, 947-
919 972, 10.1007/s11027-016-9709-9, 2017.
- 920 IEA: World Energy Outlook 2008, International Energy Agency, Paris, 2008.
- 921 Kuenen, J. J. P., Visschedijk, A. J. H., Jozwicka, M., and Denier van der Gon, H. A. C.: TNO-MACC-II emission
922 inventory; A multi-year (2003-2009) consistent high-resolution European emission inventory for air quality
923 modelling, *Atmos. Chem. Phys.*, 14, 10963-10976, 10.5194/acp-14-10963-2014, 2014.
- 924 LaFranchi, B. W., Pétron, G., Miller, J. B., Lehman, S. J., Andrews, A. E., Dlugokencky, E. J., Hall, B., Miller, B.
925 R., Montzka, S. A., Neff, W., Novelli, P. C., Sweeney, C., Turnbull, J. C., Wolfe, D. E., Tans, P. P., Gurney, K.
926 R., and Guilderson, T. P.: Constraints on emissions of carbon monoxide, methane, and a suite of hydrocarbons in
927 the Colorado front range using observations of ¹⁴CO₂, *Atmos. Chem. Phys.*, 13, 11101-11120, 10.5194/acp-13-
928 11101-2013, 2013.
- 929 Lauvaux, T., Miles, N. L., Richardson, S. J., Deng, A., Stauffer, D. R., Davis, K. J., Jacobson, G., Rella, C.,
930 Calonder, G. P., and Decola, P. L.: Urban emissions of CO₂ from Davos, Switzerland: The first real-time
931 monitoring system using an atmospheric inversion technique, *J. Appl. Meteorol. Clim.*, 52, 2654-2668,
932 10.1175/JAMC-D-13-038.1, 2013.
- 933 Lauvaux, T., Miles, N. L., Deng, A., Richardson, S. J., Cambaliza, M. O., Davis, K. J., Gaudet, B., Gurney, K. R.,
934 Huang, J., O'Keefe, D., Song, Y., Karion, A., Oda, T., Patarasuk, R., Razlivanov, I., Sarmiento, D., Shepson, P.,
935 Sweeney, C., Turnbull, J., and Wu, K.: High-resolution atmospheric inversion of urban CO₂ emissions during the
936 dormant season of the Indianapolis Flux Experiment (INFLUX), *J. Geophys. Res.-Atmos.*, 121, 5213-5236,
937 10.1002/2015jd024473, 2016.
- 938 Levin, I., and Karstens, U.: Inferring high-resolution fossil fuel CO₂ records at continental sites from combined
939 ¹⁴CO₂ and CO observations, *Tellus B Chem. Phys. Meteorol.*, 59, 245-250, 10.1111/j.1600-0889.2006.00244.x,
940 2007.
- 941 Liander: Innovatie & Diensten: Open data: <https://www.liander.nl/over-liander/innovatie/open-data/data>, last
942 access: Feb 2018.
- 943 Lin, J. C., Gerbig, C., Wofsy, S. C., Andrews, A. E., Daube, B. C., Davis, K. J., and Grainger, C. A. C.: A near-
944 field tool for simulating the upstream influence of atmospheric observations: The Stochastic Time-Inverted
945 Lagrangian Transport (STILT) model, *J. Geophys. Res.-Atmos.*, 108, 10.1029/2002jd003161, 2003.
- 946 Lindenmaier, R., Dubey, M. K., Henderson, B. G., Butterfield, Z. T., Herman, J. R., Rahn, T., and Lee, S. H.:
947 Multiscale observations of CO₂, ¹³CO₂, and pollutants at Four Corners for emission verification and attribution,
948 in: *P. Natl. Acad. Sci. USA*, 2014, 8386-8391, 2014.
- 949 Lopez, M., Schmidt, M., Delmotte, M., Colomb, A., Gros, V., Janssen, C., Lehman, S. J., Mondelain, D., Perrussel,
950 O., Ramonet, M., Xueref-Remy, I., and Bousquet, P.: CO, NO_x and ¹³CO₂ as tracers for fossil fuel CO₂: Results



951 from a pilot study in Paris during winter 2010, *Atmos. Chem. Phys.*, 13, 7343-7358, 10.5194/acp-13-7343-2013,
952 2013.

953 McKain, K., Wofsy, S. C., Nehrkorn, T., Eluszkiewicz, J., Ehleringer, J. R., and Stephens, B. B.: Assessment of
954 ground-based atmospheric observations for verification of greenhouse gas emissions from an urban region, in: *P.*
955 *Natl. Acad. Sci. USA*, 2012, 8423-8428, 2012.

956 Meirink, J. F., Bergamaschi, P., Frankenberg, C., d'Amelio, M. T. S., Dlugokencky, E. J., Gatti, L. V., Houweling,
957 S., Miller, J. B., Röckmann, T., Villani, M. G., and Krol, M. C.: Four-dimensional variational data assimilation
958 for inverse modeling of atmospheric methane emissions: Analysis of SCIAMACHY observations, *J. Geophys.*
959 *Res.-Atmos.*, 113, 10.1029/2007jd009740, 2008.

960 Miller, J. B., Lehman, S. J., Montzka, S. A., Sweeney, C., Miller, B. R., Karion, A., Wolak, C., Dlugokencky, E.
961 J., Southon, J., Turnbull, J. C., and Tans, P. P.: Linking emissions of fossil fuel CO₂ and other anthropogenic trace
962 gases using atmospheric ¹⁴CO₂, *J. Geophys. Res.-Atmos.*, 117, 1-23, 10.1029/2011jd017048, 2012.

963 Monni, S., Syri, S., and Savolainen, I.: Uncertainties in the Finnish greenhouse gas emission inventory, *Environ.*
964 *Sci. Policy*, 7, 87-98, <https://doi.org/10.1016/j.envsci.2004.01.002>, 2004.

965 Mues, A., Kuenen, J., Hendriks, C., Manders, A., Segers, A., Scholz, Y., Hueglin, C., Builtjes, P., and Schaap, M.:
966 Sensitivity of air pollution simulations with LOTOS-EUROS to the temporal distribution of anthropogenic
967 emissions, *Atmos. Chem. Phys.*, 14, 939-955, 10.5194/acp-14-939-2014, 2014.

968 Nakicenovic, N., Alcamo, J., Davis, G., De Vries, B., Fenhann, J., Gaffin, S., Gregory, K., Grübler, A., Jung, T.
969 Y., Kram, T., La Rovere, E. L., Michaelis, L., Mori, S., Morita, T., Pepper, W., Pitcher, H., Price, L., Riahi, K.,
970 Roehrl, A., Rogner, H.-H., Sankovski, A., Schlesinger, M., Shukla, P., Smith, S., Swart, R., Van Rooijen, S.,
971 Victor, N., and Dadi, Z.: IPCC Special Report on Emissions Scenarios, Cambridge University Press, Cambridge,
972 UK, 2000.

973 Nassar, R., Napier-Linton, L., Gurney, K. R., Andres, R. J., Oda, T., Vogel, F. R., and Deng, F.: Improving the
974 temporal and spatial distribution of CO₂ emissions from global fossil fuel emission data sets, *J. Geophys. Res.-*
975 *Atmos.*, 118, 917-933, 10.1029/2012jd018196, 2013.

976 Nathan, B., Lauvaux, T., Turnbull, J., and Gurney, K.: Investigations into the use of multi-species measurements
977 for source apportionment of the Indianapolis fossil fuel CO₂ signal, *Elementa*, 6, 10.1525/elementa.131, 2018.

978 Nehrkorn, T., Eluszkiewicz, J., Wofsy, S. C., Lin, J. C., Gerbig, C., Longo, M., and Freitas, S.: Coupled weather
979 research and forecasting-stochastic time-inverted lagrangian transport (WRF-STILT) model, *Meteorol. Atmos.*
980 *Phys.*, 107, 51-64, 10.1007/s00703-010-0068-x, 2010.

981 Netherlands PRTR: Netherlands Pollutant Release & Transfer Register: <http://www.emissieregistratie.nl/>, last
982 access: Mar 2018.

983 Palmer, P. I., O'Doherty, S., Allen, G., Bower, K., Bösch, H., Chipperfield, M. P., Connors, S., Dhomse, S., Feng,
984 L., Finch, D. P., Gallagher, M. W., Gloor, E., Gonzi, S., Harris, N. R. P., Helfter, C., Humpage, N., Kerridge, B.,
985 Knappett, D., Jones, R. L., Le Breton, M., Lunt, M. F., Manning, A. J., Matthiesen, S., Muller, J. B. A., Mullinger,
986 N., Nemitz, E., O'Shea, S., Parker, R. J., Percival, C. J., Pitt, J., Riddick, S. N., Rigby, M., Sembhi, H., Siddans,
987 R., Skelton, R. L., Smith, P., Sonderfeld, H., Stanley, K., Stavert, A. R., Wenger, A., White, E., Wilson, C., and
988 Young, D.: A measurement-based verification framework for UK greenhouse gas emissions: an overview of the
989 Greenhouse gAs UK and Global Emissions (GAUGE) project, *Atmos. Chem. Phys.*, 18, 11753-11777,
990 10.5194/acp-18-11753-2018, 2018.



- 991 Peters, W., Miller, J. B., Whitaker, J., Denning, A. S., Hirsch, A., Krol, M. C., Zupanski, D., Bruhwiler, L., and
992 Tans, P. P.: An ensemble data assimilation system to estimate CO₂ surface fluxes from atmospheric trace gas
993 observations, *J. Geophys. Res.-Atmos.*, 110, 1-18, 10.1029/2005JD006157, 2005.
- 994 Peters, W., Krol, M. C., Van der Werf, G. R., Houweling, S., Jones, C. D., Hughes, J., Schaefer, K., Masarie, K.
995 A., Jacobson, A. R., Miller, J. B., Cho, C. H., Ramonet, M., Schmidt, M., Ciattaglia, L., Apadula, F., Heltai, D.,
996 Meinhardt, F., Di Sarra, A. G., Piacentino, S., Sferlazzo, D., Aalto, T., Hatakka, J., Ström, J., Haszpra, L., Meijer,
997 H. A. J., Van der Laan, S., Neubert, R. E. M., Jordan, A., Rodó, X., Morguí, J. A., Vermeulen, A. T., Popa, E.,
998 Rozanski, K., Zimnoch, M., Manning, A. C., Leuenberger, M., Uglietti, C., Dolman, A. J., Ciais, P., Heimann, M.,
999 and Tans, P.: Seven years of recent European net terrestrial carbon dioxide exchange constrained by atmospheric
1000 observations, *Glob. Change Biol.*, 16, 1317-1337, 10.1111/j.1365-2486.2009.02078.x, 2010.
- 1001 Ramírez, A. R., De Keizer, C., and Van der Sluijs, J. P.: Monte Carlo analysis of uncertainties in the Netherlands
1002 Greenhouse Gas Emission Inventory for 1990 – 2004, Copernicus Institute for Sustainable Development and
1003 Innovation, Utrecht, 2006.
- 1004 Rao, P., Gurney, K. R., Patarasuk, R., Song, Y., Miller, C. E., Duren, R. M., and Eldering, A.: Spatio-temporal
1005 Variations in on-road CO₂ Emissions in the Los Angeles Megacity, *AIMS Geosciences*, 3, 239-267,
1006 <http://dx.doi.org/10.3934/geosci.2017.2.239>, 2017.
- 1007 Raupach, M. R., Marland, G., Ciais, P., Le Quéré, C., Canadell, J. G., Klepper, G., and Field, C. B.: Global and
1008 regional drivers of accelerating CO₂ emissions, *P. Natl. Acad. Sci. USA*, 104, 10288-10293,
1009 10.1073/pnas.0700609104, 2007.
- 1010 Rayner, P. J., Raupach, M. R., Paget, M., Peylin, P., and Koffi, E.: A new global gridded data set of CO₂ emissions
1011 from fossil fuel combustion: Methodology and evaluation, *J. Geophys. Res.-Atmos.*, 115, 10.1029/2009jd013439,
1012 2010.
- 1013 Sauter, F., Van Zanten, M., Van der Swaluw, E., Aben, J., De Leeuw, F., and Van Jaarsveld, H.: The OPS-model.
1014 Description of OPS 4.5.0, National Institute for Public Health and the Environment (RIVM) Bilthoven, 2016.
- 1015 Staufer, J., Broquet, G., Bréon, F. M., Puygrenier, V., Chevallier, F., Xueref-Rémy, I., Dieudonné, E., Lopez, M.,
1016 Schmidt, M., Ramonet, M., Perrussel, O., Lac, C., Wu, L., and Ciais, P.: The first 1-year-long estimate of the Paris
1017 region fossil fuel CO₂ emissions based on atmospheric inversion, *Atmos. Chem. Phys.*, 16, 14703-14726,
1018 10.5194/acp-16-14703-2016, 2016.
- 1019 Su, H., Yang, Z. L., Niu, G. Y., and Wilson, C. R.: Parameter estimation in ensemble based snow data assimilation:
1020 A synthetic study, *Adv. Water Resour.*, 34, 407-416, 10.1016/j.advwatres.2010.12.002, 2011.
- 1021 Super, I., Dellaert, S.N.C., Visschedijk, A.J.H., and Denier van der Gon, H.A.C.: Uncertainty analysis of a
1022 European high-resolution emission inventory of CO₂ and CO to support inverse modelling and network design,
1023 *Atmos. Chem. Phys. Discuss.*, <https://doi.org/10.5194/acp-2019-696>, in review, 2019.
- 1024 Super, I., Denier van der Gon, H. A. C., Van der Molen, M. K., Sterk, H. A. M., Hensen, A., and Peters, W.: A
1025 multi-model approach to monitor emissions of CO₂ and CO from an urban-industrial complex, *Atmos. Chem.*
1026 *Phys.*, 17, 13297-13316, 10.5194/acp-17-13297-2017, 2017a.
- 1027 Super, I., Denier van der Gon, H. A. C., Visschedijk, A. J. H., Moerman, M. M., Chen, H., Van der Molen, M. K.,
1028 and Peters, W.: Interpreting continuous in-situ observations of carbon dioxide and carbon monoxide in the urban
1029 port area of Rotterdam, *Atmos. Pollut. Res.*, 8, 174–187, 10.1016/j.apr.2016.08.008, 2017b.



- 1030 Terrenoire, E., Bessagnet, B., Rouil, L., Tognet, F., Pirovano, G., Létinois, L., Beauchamp, M., Colette, A., Thunis,
1031 P., Amann, M., and Menut, L.: High-resolution air quality simulation over Europe with the chemistry transport
1032 model CHIMERE, *Geosci. Model Dev.*, 8, 21-42, 10.5194/gmd-8-21-2015, 2015.
- 1033 Tolk, L. F., Meesters, A. G. C. A., Dolman, A. J., and Peters, W.: Modelling representation errors of atmospheric
1034 CO₂ mixing ratios at a regional scale, *Atmos. Chem. Phys.*, 8, 6587-6596, 10.5194/acp-8-6587-2008, 2008.
- 1035 Turnbull, J. C., Miller, J. B., Lehman, S. J., Tans, P. P., Sparks, R. J., and Southon, J.: Comparison of ¹⁴CO₂, CO,
1036 and SF₆ as tracers for recently added fossil fuel CO₂ in the atmosphere and implications for biological CO₂
1037 exchange, *Geophys. Res. Lett.*, 33, 1-5, 10.1029/2005GL024213, 2006.
- 1038 Turnbull, J. C., Sweeney, C., Karion, A., Newberger, T., Lehman, S. J., Tans, P. P., Davis, K. J., Lauvaux, T.,
1039 Miles, N. L., Richardson, S. J., Cambaliza, M. O., Shepson, P. B., Gurney, K., Patarasuk, R., and Razlivanov, I.:
1040 Toward quantification and source sector identification of fossil fuel CO₂ emissions from an urban area: Results
1041 from the INFLUX experiment, *J. Geophys. Res.-Atmos.*, 120, 292-312, 10.1002/2014jd022555, 2015.
- 1042 UNFCCC: Paris Agreement, UNFCCC, 2015.
- 1043 Van der Laan-Luijckx, I. T., Van der Velde, I. R., Van der Veen, E., Tsuruta, A., Stanislawski, K.,
1044 Babenhauerheide, A., Zhang, H. F., Liu, Y., He, W., Chen, H., Masarie, K. A., Krol, M. C., and Peters, W.: The
1045 CarbonTracker Data Assimilation Shell (CTDAS) v1.0: Implementation and global carbon balance 2001–2015,
1046 *Geosci. Model Dev.*, 10, 2785-2800, 10.5194/gmd-10-2785-2017, 2017.
- 1047 Van der Laan, S., Van der Laan-Luijckx, I. T., Zimmermann, L., Conen, F., and Leuenberger, M.: Net CO₂ surface
1048 emissions at Bern, Switzerland inferred from ambient observations of CO₂, δ(O₂/N₂), and ²²²Rn using a customized
1049 radon tracer inversion, *J. Geophys. Res.-Atmos.*, 119, 1580-1591, 10.1002/2013JD020307, 2014.
- 1050 Van Jaarsveld, J. A.: The Operational Priority Substances model. Description and validation of OPS-Pro 4.1,
1051 National Institute for Public Health and the Environment, Bilthoven, 2004.
- 1052 Vogel, F. R., Thiruchittampalam, B., Theloke, J., Kretschmer, R., Gerbig, C., Hammer, S., and Levin, I.: Can we
1053 evaluate a fine-grained emission model using high-resolution atmospheric transport modelling and regional fossil
1054 fuel CO₂ observations?, *Tellus B Chem. Phys. Meteorol.*, 65, 10.3402/tellusb.v65i0.18681, 2013.
- 1055 Wang, Y., Broquet, G., Ciais, P., Chevallier, F., Vogel, F., Wu, L., Yin, Y., Wang, R., and Tao, S.: Potential of
1056 European ¹⁴CO₂ observation network to estimate the fossil fuel CO₂ emissions via atmospheric inversions, *Atmos.*
1057 *Chem. Phys.*, 18, 4229-4250, 10.5194/acp-18-4229-2018, 2018.
- 1058 Whitaker, J. S., and Hamill, T. M.: Ensemble data assimilation without perturbed observations, *Mon. Weather*
1059 *Rev.*, 130, 1913-1924, 10.1175/1520-0493(2002)130<1913:EDAWPO>2.0.CO;2, 2002.
- 1060 Zhou, Y., and Gurney, K. R.: Spatial relationships of sector-specific fossil fuel CO₂ emissions in the United States,
1061 *Global Biogeochem. Cycles*, 25, 10.1029/2010gb003822, 2011.
- 1062
- 1063